XANES evidence for sulphur speciation in Mn-, Ni- and
W-bearing silicate melts

Evans, K.A., O'Neill, H. St. C † Mavrogenes, J.A.†, Keller, N.S., ${\rm Jang,\ L-Y\ \S}\ {\rm Lee,\ J-F\ \S}$

August 14, 2009

^{*}RSES, ANU, ACT 0200, Australia; now at Dept. Applied Geology, Curtin University, GPO Box U1987, Bentley, WA6845, Australia, tel: 0061 8 92664682, fax: 0061 8 92663153, email: k.evans@curtin.edu.au

 $^{^{\}dagger}$ Research School of Earth Sciences, Building 61, Mills Road, Australian National University, Acton 0200, Australia

[†]Woods Hole Oceanographic Institution, Clark South 283, MS24, Woods Hole, Ma. 02543

[§]National Synchrotron Radiation Research Centre, 101 Hsin-Ann Road, Hsinchu Science Park, Hsinchu 30076, Taiwan

1 ABSTRACT

5

- S K edge XANES and Mn, W and Ni XANES and EXAFS spectra of silicate glasses synthesised at 1400° C and 1 bar with compositions in the CaO-MgO-Al₂O₃-SiO₂-S plus MnO, NiO, or WO₃ systems were used to investigate sulphur speciation in silicate glasses.
- S K-edge spectra comprised a composite peak with an edge between 2470 and 2471.4 eV, which was attributed to S²⁻, and a peak of variable height with an edge at 2480.2 to 2480.8 eV, which is consistent with the presence of S⁶⁺. The latter peak was attributed to sample oxidation during sample storage. W-rich samples produced an additional lower energy peak at 2469.8 eV that is tentatively attributed to the existence of S 3p orbitals hybridised with the W 5d states.

Deconvolution of the composite peak reveals that the composite peak for Mn-bearing 15 samples fits well to a model that combines three Lorentzians at 2473.1, 2474.9 and 2476.2 16 eV with an arctan edge step. The composite peak for W-bearing samples fits well to the 17 same combination plus an additional Lorentzian at 2469.8 eV. The ratio of the proportions 18 of the signal accounted for by peaks at 2473.1eV and 2476.2eV correlates with Mn:Ca molar 19 ratios, but not with W:Ca ratios. Spectra from Ni-bearing samples were qualitatively similar 20 but S levels were too low to allow robust quantification of peak components. Some part of 21 the signal accounted for by the 2473.1 eV peak was therefore taken to record the formation 22 of Mn-S melt species, while the 2469.8 peak is interpreted to record the formation of W-S melt species. The 2474.9 eV and 2476.2 eV peaks were taken to be dominated by Ca-S and Mg-S interactions. However, a 1:1 relationship between peak components and specific energy transitions is not proposed. This interpretation is consistent with known features of 26 the lower parts of the conduction band in monosulphide minerals and indicates a similarity 27 between sulphur species in the melts and the monosulphides. S XANES spectra cannot be 28 reproduced by a combination of the spectra of the component element monosulphides. 29

Mn-, W- and Ni- XANES and EXAFS for synthetic glasses without sulphide exsolution did not show any sensitivity to the presence of sulphur, which is unsurprising as S:O ratios were sufficiently low that metals would be mostly co-ordinated by O. Mn EXAFS spectra were consistent with divalent Mn in 5 co-ordinated Mn-O melt species. W spectra were consistent with tetrahedrally co-ordinated hexavalent W, most likely in scheelite-like melt species, and Ni spectra were consistent with [4] co-ordinated divalent Ni. These results indicate lower coordinations for both W and Ni than those inferred by some previous workers. Cation coordination may reflect the proportion of non-bridging oxygens, which is lower in

the Ca-rich and Al-poor samples investigated here than for previous studies.

2 INTRODUCTION

39

A quantitative understanding of the solubility of sulphur in silicate melts is necessary if we are to understand a wide range of processes that include the distribution of commercially 41 valuable and/or geochemically significant chalcophile elements such as copper, platinum, 42 rhenium, osmium and palladium (e.g. Naldrett, 1999; Wang and Zhou, 2006), the effects 43 of volcanic eruptions on the atmosphere (e.g. de Hoog et al., 2004), global sulphur cycling (e.g. Alt et al., 1993; Fischer et al., 1998), and the formation of magmatic-related economic 45 sulphide deposits (e.g Cawthorn, 2005; Lee and Ripley, 1995; Li and Ripley, 2005; Pina et 46 al., 2006). Quantitative understanding and predictions (e.g. O'Neill and Mavrogenes, 2002; 47 Moretti and Ottonello, 2005) are based on the thermodynamic properties of the melts, which depend on the melt structure, which is a combination of long range structure and the short range combinations of anions and cations that are described as melt species. The term melt 50 species is used here to describe the arrangement of the first coordination shell around the 51 element of interest. Melt structure and speciation have historically proved difficult to identify 52 unambiguously. This is because the temperatures at which silicate melts form are difficult 53 to access with experiments, and because magnified atomic vibrations at high temperature 54 induce significant deviations in the spectroscopic signal so that the results of investigations 55 rarely provide a unique or well constrained solution. Quantification of sulphur speciation is 56 particularly challenging because sulphur can occur in a wide range of oxidation states, from -2 to +6, and because sulphur is a low atomic number element present in low concentrations in 58 silicate melts (less than 1wt%, and often less than 0.1 wt%) which restricts the applicability 59 of many spectroscopic techniques. 60

Methods used to derive information on the speciation of sulphur in silicate melts include relationships between the concentrations of sulphur and possible ligands such as Fe (O'Neill and Mavrogenes et al., (2002), Mn (Evans et al., 2008), and values of intensive variables such as fO_2 and fS_2 (e.g. Carroll and Rutherford, 1985; Fincham and Richardson, 1954; Wallace and Carmichael, 1992; Carroll and Webster, 1994; O'Neill and Mavrogenes et al., 2002) and the position of the S peak determined by electron microprobe (e.g Carroll and Rutherford,

1988; Metrich and Clocchiati, 1996). These works have shown that sulphur concentrations are highly dependent on melt composition, particularly Fe and Mn (O'Neill and Mavrogenes et al., 2002; Evans et al., 2008), which suggests that S forms melt species with these elements. 69 Ni partition coefficients between silicate melt and olivine (Li et al., 2003) have been used 70 to infer the existence of a Ni-S melt species. Sulphur concentrations are also dependent 71 on the ratio of sulphur to oxygen fugacities, and the form of this relationship indicates that 72 sulphur dissolves as S^{2-} under reducing conditions similar to those found at mid-ocean ridges 73 and as S^{6+} under conditions comparable to those found in the more oxidised arc volcanic 74 environments (e.g. Metrich and Clocchiati, 1996), and that S^{2-} or SO_4^{2-} replace O^{2-} on the 75 anion sublattice (e.g. Fincham and Richardson, 1954).

Further information on the details of sulphur speciation in silicate melts has been provided 77 by XANES (X-ray Absorption Near edge Structure) studies (e.g. Paris et al., 2001; Bonin-78 Mosbah et al., 2002; Fleet, 2005; Farges et al., 2006; Wilke et al., 2008; Backnaes et al., 2008; 79 Metrich et al., 2009). XANES uses high fluxes of monochromated light from synchrotron 80 light sources to investigate the details of absorption edges of the elements of interest. The 81 spectra consist of measurements of X-ray absorption by a sample as the energy of the X-82 rays is scanned in small steps from around 100eV below to 100 eV above the absorption 83 edge. Variation in absorption in the vicinity of the edge are created by variation in the 84 probabilities of the different electron transitions as a function of the incoming X-ray energy 85 and by multiple scattering effects. These features are ultimately controlled by the immediate atomic environment of the element of interest, rather than medium or long-range ordering 87 features, and XANES spectra have been used to to determine oxidation state, site geometry, 88 and the identity of nearest neighbours in a wide variety of materials (e.g., Li et al., 1994; Li 89 et al., 1995; Farges et al., 1997; Berry et al., 2006; Liu et al., 2006). 90

XANES spectroscopy on sulphur-bearing materials of geological interest have been reviewed by Fleet et al. (2005). The sulphur K edge, which is of interest here, is attributed to the transition of S 1s electrons to unoccupied antibonding orbitals at the bottom of the conduction band. The position of the edge changes as a function of sulphur oxidation state, with up to a 12 eV difference between S²⁻ and S⁶⁺ (e.g. Li et al., 1995), and, to a lesser

extent, as a function of the electronic structure of the lower parts of the conduction bands in the materials examined. Considerable attention has been paid to the XANES spectra of the monosulfides FeS, NiS, CoS, MgS, CaS and MnS (e.g. Farrell and Fleet, 2000; Farrell and 98 Fleet, 2001; Farrell et al., 2002). Results of this work (Fig. 1a) suggest that monosulphides 99 that exhibit a higher degree of covalent bonding, such as FeS and CoS, have a lower energy 100 absorption edge that results from the transition of 1s S electrons into unoccupied S 3p σ^* 101 orbitals hybridised with the metal 3d states. Monosulphides with bonding that has a more 102 ionic character, such as MgS and CaS, have a slightly higher absorption edge energy because 103 the metal 3d states are effectively unavailable for hybridisation in these materials. Spectra 104 for solid solutions between the monosulphide end-members, where accessible, are not linear 105 combinations of the component end-members, which indicates that electronic properties are 106 not a linear function of sample composition. These conclusions are supported by calculated 107 XANES spectra for the monsulphides (Soldatov et al., 2004; Kravtsova et al., 2004). 108

Most studies of XANES for sulphur-bearing silicate melts have been consistent with the 109 presence of S^{2-} in melts from more reduced settings and S^{6+} in melts from more oxidised 110 environments (e.g Paris et al., 2001; Fleet et al., 2005; Metrich et al., 2009). The presence of 111 tetravalent sulphur has also been proposed on the basis of XANES spectra (Metrich et al., 112 2002; Bonin-Mosbah et al., 2002), but it has since been demonstrated that, in most cases, 113 this species is an artifact of the intense microbeam used for those measurements (Wilke et 114 al., 2008; Metrich et al., 2009). Spectra for S²-bearing natural glasses have been found to be similar to linear combinations of the XANES of sulphides of the main cations in the 116 sample, FeS and CaS (Fleet et al., 2005). Formation of CaS_n polyhedral melt species, where 117 n is an unknown stoichiometric coefficient, was also proposed for CaS dissolved in a synthetic 118 sodium aluminosilicate glass (Fleet et al., 2005). Metrich et al. (2009) show qualitatively 119 that the S XANES spectra of synthetic glasses prepared under reducing conditions such that 120 all S is present as S^{2-} are sensitive to the presence or absence of FeO. These observations lead 121 to the hypothesis that S²⁻ in silicate melts is speciated as monosulphide-like melt species of 122 unknown geometry and stoichiometry. 123

The aim of this study is to use XANES spectra of S in Mn-, Ni- and W-bearing syn-

124

thetic silicate glasses to investigate the hypothesis that sulphur forms monosulphide-like 125 melt species in silicate melts. These potential ligands were chosen because of their very different effects on sulphur solubility (Evans et al., 2008). Mn strongly enhances sulphur 127 solubility in a similar way to Fe, but has the advantage that its valence is predicted to be 128 2+ over a wide $f(O)_2$ range, so coupled changes in Mn and S valence during cooling are 129 effectively precluded. Ni at low ppm concentrations causes saturation of the melt with an 130 immiscible NiS phase, and W has little effect on S solubility, which suggests that formation 131 of W-S melt species is limited. Observation of S speciation of melts that contain these three 132 metals should therefore cover a wide range of melt speciation types. XANES and EXAFS 133 (Extended X-ray Absorption Fine Structure) spectra for Mn, Ni and W were also measured 134 to enable a complementary study of the speciation of these metals. 135

3 METHODS

136

138

3.1 Preparation of synthetic glasses

Preparation methods are described in detail by Evans et al. (2008a). However, brief details 140 are provided here. Sample compositions were based on a CMAS (CaO-MgO-Al₂O₃-SiO₂) 141 anorthite-diopside eutectic mix with added CaO. The Ca was added to increase the base level 142 S concentration in the samples, which reduces analytical uncertainties. This mix is referred 143 to here as CMAS. A number of samples were prepared with zero additive metal content 144 to establish the baseline S XANES for the study. Comparison of these samples allows the 145 repeatability of runs at identical, theoretically, imposed fO_2 and fS_2 values to be assessed. 146 Oxides of Mn, Ni and W were added to the dried oxides and carbonates of Ca, Mg, Al and 147 Si in concentrations that varied from 0.1 wt % to 10 wt%. The exception is run B070406, 148 which contained Ni at 10, 15, 20, 25 and 30 wt \%. Sample notation is of the form Bxxxxxxpy, 149

where xxxxxx indicates the date of the run, and y indicates the sample number from that run.

The glasses were synthesised in vertical muffle tube furnaces which have been modified 152 to allow accurately measured gas mixes to flow upwards through the furnace. All runs were 153 equilibrated at 1400°C. fO_2 and fS_2 were controlled by the proportions of CO, CO₂ and 154 SO₂, which were supplied to the furnace by Tylan F2800 mass flow controllers. Values of 155 fO_2 and fS_2 corresponding to the input gas mixes were calculated as described by O'Neill 156 and Mavrogenes (2002). Uncertainties in fO_2 and fS_2 are estimated to be \pm 0.05 log units 157 (c.f. O'Neill and Mavrogenes, 2002). Samples were run for 24 hours. Previous work (O'Neill 158 and Mavrogenes, 2002) suggests that this time is sufficient for equilibration between the glass 159 and the input gases even when immiscible sulphide melts are formed. After 24 hours the runs 160 were quenched by release of the melt beads into water so quench was close to instantaneous. 161 Such a fast quench should be sufficient to preserve the composition of the melt at 1400°C, 162 with respect to sulphur content, and to preserve melt speciation and structure at the melt-163 glass transition. However, it is possible that electron exchange continues down to room 164 temperature. Most of the experiments were performed at an fO_2 of -9.6 ($\Delta QFM=-3.2$) and 165 a log fS_2 of -1.91. These conditions have the advantage that they produce high S contents (> 166 0.2 wt % in most cases) without any risk of blocking the furnace with precipitated elemental 167 S. A small number of experiments were performed at more oxidising conditions but sulphur 168 has previously been shown to be present as S²⁻ in the melt for all the conditions investigated 169 here (O'Neill and Mavrogenes, 2002) except those of B190606 ($\Delta QFM=3.4$). Run conditions 170 for runs that produced data used here are summarised in Table 1. 171

Glass products were split into portions for electron microprobe and synchrotron analysis.

Samples for synchroton analysis were placed into polyethylene containers and stored. The
glass beads were kept whole as far as possible, but some disintegrated somewhat during
quench and the sample splitting process. Samples for probe analysis were set in EPO-FIX
epoxy and polished. These samples were then examined optically for evidence of exsolution
of immiscible sulphide and silicate crystal phases. The epoxy-mounted samples were then
carbon-coated and analysed for major elements plus S, Mn, W and Ni on the Cameca SX100

electron microprobe at the Research School of Earth Sciences at the Australian National 179 University. WDS analysis and a 15 KeV accelerating voltage was used for all elements. 180 Raw counts were converted to element weight percentages using a modified ZAF correction 181 scheme. The major elements were analysed first using a 10 nA, 15 KeV beam with a 10 182 micron radius. Sulphur, Mn, W and Ni were subsequently analysed using a 100 nA, 20 183 micron beam. Details of the sulphur measurement routine, standards, detection limits and 184 uncertainties are provided by Evans et al. (2008a). Time series by O'Neill and Mavrogenes 185 (2002) suggest that the length of experiment utilised here is sufficient for equilibrium to be 186 reached, and the good repeatability of the chemical compositions between runs for this study 187 (Evans et al., 2008) supports this conclusion. However, additional tests for equilibrium were 188 also made in the form of multiple measurements spatially distributed over the glass surface. 189 Samples showing variation in excess of 2\% relative to the mean value were discarded. 190

191 3.2 XAS analysis

Experimental glass products designated for XANES analysis were dried, where necessary, and 192 stored in glass tubes. The material was taken to the 1.5 GeV NSRRC (National Synchrotron 193 Radiation Research Centre) in Hsinchu, Taiwan, which operates in a 300 mA top-up mode. 194 Sulphur K-edge spectra were measured on BL16A (Dann et al., 1998). This beamline uses a 195 water-cooled Si(111) monochromator to provide monochromated X-rays in the energy range 196 2000 - 8000 eV. The monochromator energy resolution in the vicinity of S K-edge (2472) 197 eV) is 0.5 eV. Photon flux is a maximum of 3×10^{11} photons over an area around 4×2 198 mm. K-edge spectra for Mn, W and Ni were measured on BL17C. This beamline is provided 199 with photons by a wiggler with 20cm period magnets, and energy is monochromated with a 200 Si(111) double crystal monochromator in the range 4 - 15 KeV. Photon flux is 10^9 to 10^{10} 201 photons per second over an area approximately 8x3mm, and the resolution at the Mn K, Ni 202 K and W L3 edges is around 1 eV. 203

$_{04}$ 3.2.1 Sulphur analysis

Glasses and standards were prepared for synchrotron analysis by crushing in an agate mortar 205 under acetone. Crushing was employed to increase sample surface area, expose fresh surfaces 206 for analysis, and decrease the likelihood of data artifacts as a result of self-absorption. Epoxy-207 mounted samples could not be used for the XANES because of the high S content of epoxy 208 relative to the samples. Calculated absorption lengths, which is the length over which the 209 proportion of penetrating X-ray drops by a factor of e, were 5-7 microns. The crushed 210 material was applied onto S-free kapton tape with a paintbrush. The agate mortar and 211 paintbrush was cleaned thoroughly after each sample was loaded, then re-used. Sample 212 preparation blanks were undertaken by going through all the preparation steps without any sample in the agate mortar. Cl-XANES spectra measured for the sample preparation blanks (Evans et al., 2008b) did not show a signal so cross-contamination between samples is likely 215 to have been negligible. S-XANES was measured for sample-free tape and was found to 216 be free of any sulphur signal. Four samples at a time were loaded onto the sample holder 217 and introduced into the sample chamber. The sample chamber was purged with helium 218 for at least an hour after the sample introduction to remove air introduced during sample 219 loading. The beam upstream of the sample chamber is contained within a helium-filled tube, 220 to minimise X-ray absorption by the air. S K edge spectra were measured in fluorescence 221 geometry, with the sample at 45 degrees to the beam. The fluorescent signal was detected using the in-house Lytle detector. The energy calibration was checked daily with a Mo foil 223 (L3 edge: 2520 eV). Shifts in the monochromator energy were always less than 0.05 eV. 224

The XANES measurement routine consisted of 1eV, 2 second measurement steps from 2272 to 2442 eV, 0.2 eV, 2 second measurements in the S K edge XANES region from 2442 eV to 2502 eV, and 2 eV, 2 second measurements in the post edge region from 2502 eV to 2847 eV. Passage of the scan over the Cl edge at 2822 eV provided a useful test for contamination from epoxy or other factors as the samples were nominally Cl-free. Scans were repeated three times and averaged to optimise signal quality. Total sample collection time was around three hours for each sample.

Standards investigated included MnS, NiS (millerite), WS₂, NiSO₄, CaSO₄, MnSO₄, FeS 232 and FeS₂. The MnS was sourced from a new container of synthetic MnS from Sigma Aldrich (99\% pure). This container was opened for the first time at the beamline on the day of 234 the standard analysis. Millerite and pyrite were taken from natural samples provided by 235 the mineral collection at the Department of Earth and Marine Sciences at the Australian 236 National University. Other standards were taken from the existing chemical collection at the 237 Research School of Earth Sciences, Australian National University. All standards except the 238 MnS were checked with X-ray diffraction prior to the synchrotron experiment. Results were 239 analysed with Siroquant (Taylor, 1991) and contaminants were not detected. It was not 240 possible to measure spectra for CaS and MgS standards because of the extreme tendency of these phases to oxidise on contact with air, which necessitates glove-box facilities for sample 242 insertion. Such facilities were not available at the NSRRC beamline. Instead, spectra were 243 obtained by digitisation of published spectra from Farrell and Fleet (2000). These spectra 244 were aligned with XANES measured for this work by comparison of acquired FeS spectra 245 with digitised FeS spectra from Farrell and Fleet (2000).

The potential for oxidation during sample preparation was investigated because sulphur is susceptible to preparation-induced oxidation (Fleet, 2005). Two subsamples from four of the glass beads were prepared (a) by grinding under acetone in air and (b) by grinding under acetone in a nitrogen-filled glove box. XANES spectra for the two sub-samples were then compared.

252 3.2.2 Mn, W and Ni analysis

Standards analysed included Mn, Ni, and W metals, MnO, MnO₂, MnS, WS₂, WO₃, CaWO₄
(scheelite), NiS (millerite) and NiO. The metal standards were foils provided by the beamline.

MnO, MnO₂, MnS and WS₂ were 99% pure or better commercial powders from SigmaAldrich. Spectra for the MnO, MnO₂, MnS and WS₂ were taken on the same day that the
containers, previously sealed by the manufacturers, were opened. CaWO₄ was a natural
scheelite; the sample was checked by XRD and found to be >99% pure scheelite. WO₃ and

NiO were >99% pure powders purchased from Sigma-Aldrich. NiS was a natural millerite sample; the identity and purity of the material were confirmed by XRD.

The monochromator energy was calibrated using the metal foils. The first derivative of the Mn K edge, W L3 edge and Ni K edge were set to 6539, 8333 and 10207 eV respectively prior to analysis for each of these metals. All standards except for the pure metals were powdered under acetone in an agate pestle and mortar and applied to kapton tape with a paintbrush. Standard analysis utilised the transmitted signal, to avoid problems with self absorption.

Metal XANES and EXAFS measurements on the samples were performed on either crushed sample applied to tape, as described above for the S K edge measurements, or on the polished sample fragments set in epoxy that were used for the electron microprobe analysis. In the latter case, flexible lead sheet was used to shield the parts of the mount that did not contain the sample under investigation. No systematic difference in sample spectra were noted as a result of the two different sample preparation methods. Sample analysis utilised the fluorescence signal.

Scans across the Mn K-edge were measured for 2 seconds per point in 7 eV steps between 275 200 and 25eV below the edge, for 2 seconds per point in 0.3 eV steps in the XANES region 276 from 25eV below to 50 eV above the edge, and for 4 seconds per point in 0.06 k unit steps in the EXAFS region from 50 to 540eV above the edge. k is the relative wave number of the photo-electron ejected in the absorption process, and is related to energy by $k^2 = 2m_e(E - E_0)/\hbar^2$, where m_e is the mass of an electron, E and E_0 are the energy and edge energy respectively, and \hbar is the reduced Planck's constant.

Scans across the W L3 edge were measured for 2 seconds per point in 7.5 eV steps between 200 and 20eV below the edge, for 2 seconds per point in 0.4 eV steps in the XANES region 20eV below to 40 eV above the edge, and for 2 seconds per point in 0.06 k unit steps in the EXAFS region from 80 to 1000eV above the edge.

Scans across the Ni K edge were measured for 2 seconds per point in 10 eV steps between 200 and 20eV below the edge, for 2 seconds per point in 0.4 eV steps in the XANES region

from 20eV below to 80 eV above the edge, and for 4 seconds per point in 0.06 k unit steps in the EXAFS region from 80 to 1000eV above the edge.

289 3.3 Post Run Analysis and Interpretation

₂₉₀ 3.3.1 S - XANES

XANES spectra for Mn, W, Ni and S were loaded into the Athena (Ravel and Newville, 291 2005) software package, which was used for visualisation and manipulation of the data. The 292 multiple spectra from each glass composition were checked for repeatability, averaged, and 293 normalised to an edge step of one, where the edge step was obtained from the difference 294 between a straight line fit to the pre-edge points and a polynomial fit to the post-edge 295 points extrapolated back to the absorption edge. Spectra for the sulphur standards were 296 corrected for self-absorption using the routine embedded in Athena. The need for self-297 absorption corrections for the synthetic and natural glasses was assessed using the same 298 routine and was found to be negligible. The position of the main edge was taken to be 299 recorded by the position of the maximum first derivative for each spectra. S K edge spectra 300 were compared with those of the sulphide and sulphate standards as a first order attempt 301 to identify speciation via a fingerprinting approach. 302

The peak deconvolution facility embedded in Athena was used to quantify the relation-303 ships between XANES and sample compositions. Initial fitting attempts were used to refine 304 the number of peaks, peak shape and peak values. Subsequent fits with set peak positions 305 were used to solve for unknown peak areas and widths. The final fitting process utilised, 306 for Mn and Ni bearing samples, Lorentzian peaks centred at 2473.1 eV, 2474.9 eV, 2476.2 307 eV and, where such a peak was present, at 2481.4 eV. These peaks were superimposed on a 308 arctan function at 2470.8 eV. W-bearing samples required an additional Lorentzian centred 309 at 2469.8 eV, which is subsequently referred to as the W peak. The width of the arctan 310 function was set to 0.3, and the widths and amplitudes of the other peaks were solved for 311 simultaneously by the program. Fits were stable and converged to the same values for a 312

wide range of starting guesses. The stated uncertainties are propagated from analytical uncertainties based on the noise level of the spectra, and on the goodness of fit. Correlation coefficients were calculated for ratios between peak proportions, normalised to exclude contributions at 2481.4 eV (sulphate signal, see discussion below), and the Ca:additive metal molar ratio.

3.3.2 Mn, W and Ni XANES and EXAFS

XANES for Mn, Ni and W were compared to those of the standards. Particular attention 319 was paid to the Ni pre-edge peak, which as been shown to provide information on Ni co-320 ordination (Farges et al., 2001). The pre-edge peak was isolated by a manual spline fit to the main trend of the edge, and the background removed to isolate the pre-edge peak. 322 The centroid of the pre-edge peak was then measured by fit of a Lorentzian peak to the 323 background subtracted XANES spectra. This process was repeated two or three times to 324 assess the sensitivity of results to the manual fitting process. The difference between the 325 centroid energy of the sample and of NiO (E_{pre-edge.sample}-E_{pre-edge.NiO}) was measured and 326 compared to the calibration of Farges et al., (2001). 327

The majority of the metal spectra were identical, so the best quality spectra for each 328 metal were averaged, the baseline was subtracted and the signal was transformed into k329 space, multiplied by k^2 to increase the signal at high k values, and exported into Artemis 330 (Ravel and Newville, 2005). These data were Fourier-transformed and fit to the EXAFS 331 equation (e.g. Fulton et al., 1996). QFS (Quick First Shell) models were adopted for the 332 EXAFS fits; these are appropriate given the probable lack of long range order or structure 333 around the metal-bearing species in the melts. Geometries and initial coordinations for the 334 models were taken from literature sources. Mn models were based on those of McKeown 335 et al. (2003) with octahedrally coordinated oxygens at 2.07 Å. W models were attempted 336 for both octahedrally and tetrahedrally coordinated WO_3 and WO_4^{2-} species respectively, 337 with oxygens at 2.00 Å in the octahedrally coordinated model (Poirier et al., 2005b) and at 338 1.78 Å in the tetrahedrally co-ordinated model (Hazen et al., 1985). Ni was modelled on

the assumption of formation of octahedral melt species with nearest neighbour oxygens at 2.00 Å; tetrahedral QFS models were also tried and the final results were not sensitive to the initial geometry specification.

The Fourier-transformed spectra were initially fit for four parameters: n, the number of nearest neighbours, E_0 , the energy of the edge, σ^2 , the mean square relative displacement, which measures the disorder present in the signal, and ΔR , the difference between the starting guess and modelled distance for i-O where i represents the metal of interest. These four parameters control the shape of the Fourier-transformed signal. To a first approximation, peak size increases with n, peak position is related to E_0 and ΔR , and peak breadth increases with σ^2 .

The fit value of n can represent an average co-ordination, so, for example, a value of 5 350 could represent a mixture between tetrahedrally and octahedrally co-ordinated cations with 351 similar bond lengths and comparable melt species geometry. Fits were tried that utilised an 352 additional fit parameter, c3, the third cumulant, which measures the asymmetry in the signal 353 disorder, but fits were no better, uncertainties did not decrease, and values of c3 were within 354 error of zero. For these reasons the fits documented here used a set c3 of zero. Estimates 355 for numbers of nearest neighbours require specification of S_0^2 , the passive electron reduction 356 factor, which is multiplied by the number of neighbours in the EXAFS equation (see Fulton 357 et al. (1996) for details). S_0^2 for Mn was taken to be 0.75 (after McKeown et al., 2003), 358 and to be 1 for W and Ni. The value of 1 is probably an overestimate as S_0^2 generally falls 359 between 0.8 and 1, so numbers of nearest neighbours for W and Ni may be underestimated 360 by up to 20%. 361

The k-range for fitting was taken to run from the first node in the k oscillations outside the XANES region i.e. at around 20-60eV, to the last node before excessive noise destroyed signal quality. This led to k-ranges for the fits from around 3 to around 9 for all Mn, W and Ni. Fits were considered successful when the R-factor for the fit was less than 0.05. The sensitivity of the fit to input parameters such as k-range and k-weighting was also explored and fits that were excessively affected by changes in these parameters were rejected. Uncertainties were taken from the Artemis output and reflect the combination of the effects of statistically

derived uncertainties in the signal and correlations between the fit parameters. Final fits for Ni utilised a set value for the edge energy, which was specified to be the energy at the maximum first derivative of the edge, because correlations between the fit parameters led to a relatively unconstrained solution when the edge energy was allowed to vary freely. The best fit value for the edge energy should not be significantly different to that of the centre of the edge, so this assumption is justified. Fits for Mn and W allowed E₀ to vary and its proximity to the measured edge was used as a complementary measure of fit quality.

<end small type>

377

378

376

4 RESULTS

9 4.1 Characteristics of synthetic silicate glasses

The appearance and composition of the glasses are described in detail in Evans et al. (2008a). 380 However, essential details are summarised briefly here. Most of the samples formed translu-381 cent clear or yellow-coloured glass beads up to 5mm in diameter and 0.05 to 0.15g in weight. 382 A number of samples contained bubbles and/or black specks. In most cases the black specks 383 occurred on the margins of the sample, or showed dendritic form in fractures but appeared 384 to be a surface feature, probably produced on quench, and of insufficient size for analysis. 385 However, in some Ni-bearing samples specks were distributed through the glass and appeared 386 to define flow patterns; probe analysis (Evans et al. 2008a, not shown here) revealed that 387 this apparently exsolved material contained Ni and S with minor Fe. XANES features of these samples are discussed below. Multiple microprobe analysis produced good repeatabil-389 ity in the results for over 90% of the samples analysed, which supports the proposal that 390 equilibrium was attained in most cases. A small number of the glasses produced high to-391 tals, up to 102 wt%. These totals are associated with slightly higher Si contents (Table 2), 392 and are therefore attributed to issues with Si analysis which have been noted previously on 393 this machine. Concentrations of other elements such as Mn, W, Ni and S were checked by 394

comparison against a secondary standard, and shown to be reliable, so the implications and conclusions of this study are unaffected.

Runs without added Mn, W, Ni produced a dissolved S concentration for the CMAS 397 composition of 0.199 ± 0.026 wt % (19 analyses - see Evans et al., 2008 for details). Sulphur 398 content was positively correlated with Mn concentration for the data from the Mn-bearing 399 experiments (Table 2a); sulphur contents were lower at the same Mn content for the more 400 oxidising experimental conditions. S content is effectively independent of W content (Table 401 2b) and in runs where $\log fO_2$ was increased above -9.6, the S content drops below detection 402 limits for all samples with appreciable W (Evans et al., 2008). Such samples are not discussed 403 here. Ni and S are negatively correlated, with an inverse relationship between S and Ni (Table 404 2c). This is consistent with equilibration of the melt with a Ni sulphide phase, which was 405 observed to have exsolved from the melt in many of the samples (Evans et al., 2008). Low 406 Fe concentrations are observed in a number of samples. This is attributed to contamination 407 during sample preparation or synthesis. The presence of such low quantities of Fe does not 408 affect the conclusions drawn here. 409

410 4.2 Sulphur K-edge XANES

411 4.2.1 Standards

S K edge XANES for the standards show similar characteristics to those noted in the liter-412 ature (e.g. Li et al., 1995; Fleet, 2005). Edge energy increases with sulphur oxidation state 413 and with degree of ionicity of bonding at any given oxidation state (Figs 1b,2). The lowest 414 energies are found for FeS and NiS with energies of 2468.7 and 2469 eV respectively. The 415 edge for WS₂, which also contains S as S²⁻, occurs at slightly higher energy (2470.3 eV), 416 and is coincident with the edge for FeS₂, which contains S⁻. The edge energy for MnS is 417 2470.7 eV, significantly higher than for FeS, which is interesting given the similar effects 418 that Mn²⁺ and Fe²⁺ have on S solubility. The non-transition element monosulfides, CaS 419 and MgS, have higher energies again, with edges at 2471.2 and 2471.8 eV respectively. Edge 420

energy for the sulphates occurs 10-12 eV higher than for the sulphides. The identity of the cation in a sulphate compound has much less effect on the edge energy than for sulphides (Fig. 3), with the edges for FeSO₄ and MnSO₄ at 2480.2, and that for NiSO₄ at 2480.8 eV. 423 The reason for this insensitivity is that O is the nearest neighbour to S in sulphates, so the 424 cation has a much weaker effect on the local electronic environment than it does in the case 425 of the sulphides. Note that tabulated energies (Tables 1 and 2) are edge energies rather than 426 peak energies, which are tabulated in some literature sources (e.g. Li et al., 1995). Edge 427 energies are used here because they are less sensitive to self absorption artifacts than peak 428 energies. 429

The spectra for NiS, FeS and WS₂ also show a small peak at around 2481eV (Fig. 2). This
is in the same position as the sulphate peak, and is attributed to minor standard oxidation
during sample storage that was not detected by the pre-run XRD analysis. It is unlikely
to be due to oxidation during sample preparation because the MnS standard, which is very
prone to oxidation, does not show such features in spite of identical preparation methods.

4.2.2 Synthetic silicate glasses

S K edge spectra for synthetic silicate glasses without sulphide exsolution were significantly different to the spectra of any of the sulphide or sulphate standards (Fig. 4). Well resolved spectra with sufficient information for further quantitative analysis were obtained for samples with S contents greater than 0.18 wt%. Less well resolved spectra with limited information were obtained for samples with S contents between 0.02 and 0.18 wt%.

The main features exhibited by the silicate glass XANES were (1) a composite peak with an edge, for the majority of samples, between 2470 and 2471.5 eV (Fig. 4, Table 2). Mn-, W- and Ni-bearing glasses all produced this peak and its edge energy was not systematically related to sulphur or metal content, or to run conditions. Uncertainties in the edge position, based on measurements of the edge energy from the three spectra for each sample, are of the order of 0.3 eV. Interpretation of the composite peak is discussed further below. Ni-bearing samples from one run, B010806, exhibited edge energies slightly lower than the norm for this

composite peak (Table 2c) but the duplicate of this run, B140906, did not produce the lower edge energies. B010806 samples contained black specks inferred to be exsolved NiS while B140906 did not. However, S contents in the Ni-bearing samples were low, so the spectra 450 are noisy and difficult to interpret unambiguously; (2) a low energy peak at around 2469 451 eV that was produced only by the W-bearing samples with W contents greater than 1 wt% 452 (runs B141106 and B220806). Run B150906, which contained W contents less than 1 wt \%, 453 produced spectra without a discernible 2469 eV peak, referred to hereafter as the W peak; 454 (3) A peak in the same place as that for the sulphate standards with an edge at 2480.8 eV. 455 The size of this peak relative to that of the composite peak was not related in a systematic 456 way to sample composition, preparation methods or experimental run conditions. 457

58 4.2.3 Spectral Deconvolution

Good fits (all χ^2 < 2, Table 3, Fig. 5) were obtained for the CMAS, CMAS+Mn, and 459 CMAS+W glasses with sulphur contents greater than 0.18 wt%. It was not possible to 460 obtain unique solutions to the fits for the Ni-bearing glasses because the sulphur contents 461 were low and the signal:noise ratio was too low to uniquely constrain the size of the four 462 peaks. The proportional area of each of the peaks were calculated (Table 3). Two samples 463 with proportional areas for the sulphate peak at 2481.4 eV greater than 20% (B310706p5 and 464 B151106p1) were excluded from further analysis (see discussion) because post-run oxidation 465 is sufficient to have seriously affected the distribution of sulphur species in the sample. 466

Results (Table 3, Table 4, Fig. 6) showed that there was a good correlation between the proportional areas of the 2473.1 eV peak and the molar Ca:Mn ratio (r=-0.84, Fig. 6a) and between the 2474.9 eV peak and the molar Ca:Mn ratio (r=-0.84, Fig. 6b). These values are significant at the 95% level. Significant correlations were also noted between the Mn concentration and the proportional areas of peaks at 2473.1 eV and 2476.2 eV, and the sulphur concentration and the proportional areas of peaks at 2473.1 eV and 2476.2 eV (Table 4). Mn-bearing glasses from two different runs lie on the same trend. No such correlation was noted between peak proportions and the Ca:W ratio or other compositional parameters

in the W-bearing samples (Fig. 6c, d). The correlation coefficient for the 2473.1 eV peak and the molar Ca:W ratio was -0.43, while that for the 2474.9 eV peak and the molar Ca:W ratio was 0.06. These values are not consistent with a statistically significant correlation. The area of the W peak did not relate systematically to W content, but the peak was poorly resolved so systematic relationships may not have been detectable by this method.

480 4.3 Mn, W and Ni K-edge XANES and EXAFS

81 4.3.1 Standards

Standard spectra are consistent with those described in the literature (e.g. McKeown et al. 482 (2003); O'Neill et al., (2008)). MnO (Fig. 7a) shows a pronounced white line and a small 483 pre-edge peak, while MnO₂ shows a larger pre-edge peak but a much less intense white line. 484 The spectra for MnS has a lower edge energy, and lacks any pre-edge peak. W spectra are 485 sensitive to the oxidation state of W, with a lower edge energy for tetravalent W in WS₂, 486 and higher values for hexavalent W in WO₃ and CaWO₄. Tetrahedrally co-ordinated W 487 in CaWO₄ shows a post-edge bump (Fig. 7b), which is absent in the spectra derived from 488 octahedrally co-ordinated W in WO₃. Ni spectra also show increasing edge energy with 489 increasing oxidation state (Fig. 7c). NiO and NiS spectra are distinctly different, with a 490 pronounced pre-edge peak present in the NiO spectra. 491

492 4.3.2 Synthetic silicate glasses

Mn XANES spectra from the glasses resemble that of MnO (Figs 7a and 8a), and the edge position closely matches that of MnO standards measured by McKeown et al. (2003). This implies that Mn is present as the divalent state in the silicate glass. There is no discernible variation in the XANES spectra as a function of Mn:S ratio in the melt (Fig. 8a) or the $\log fO_2/fS_2$ of experiment. The Fourier transform of the EXAFS signal shows a single peak. The shape, location and size of this peak provides information on the nearest neighbours

L3 edge XANES spectra for the W-bearing glasses (Fig. 8b) are all very similar to those 506 of O'Neill et al. (2008), who demonstrated from solubility measurements that W dissolves 507 in silicate melts in the 6+ oxidation state at the conditions relevant to this study. The 508 spectra exhibit an intense white line (A) and a post-edge bump (B) before the EXAFS oscillations begin. The post-edge bump is similar but less intense than that shown by 510 tetrahedrally co-ordinated W in Na₂WO₄ (Fig. 7b) and scheelite, CaWO₄ (Farges et al., 511 2006). There is no systematic change in spectra as a function of the oxidation state of the 512 experiment, W:S ratios or absolute W and S contents. Linear combination fitting of the 513 spectra to the two standards using the facility in Athena indicates that a contribution from 514 the tetrahedrally coordinated component alone reproduces the spectra best; addition of an 515 octahedrally coordinated component does not improve the fit. There is no evidence for a 516 WS₂ component. The EXAFS for the W-bearing samples fit well to a model based on a 517 tetrahedral arrangement of four oxygens at 1.75 Å from a central W atom (Table 5, Fig. 518 9b). EXAFS analysis in which the W-bearing spectra were fit to a WO₃ model, with six 519 oxygens at an average of 2.00 Å distance were unsuccessful. Further attempts in which the 520 oxygens were allowed to vary in distance as described by Poirier et al. (2005b) did not 521 improve the fit quality. These results confirm the suggestion of O'Neill et al. (2008) that 522 W occurs in the 6+ state in silicate melts, and are consistent with the presence of W in the 523 silicate melts investigated here as WO₄ tetrahedra. 524

Ni K edge XANES spectra for the Ni-bearing glasses are qualitatively similar to those reported by Farges et al. (2001) for divalent Ni in [4], [5] and [6] coordination and superficially resembles that of the NiO standard (Fig. 7c, 8c). The structure is slightly simpler than that for NiO, however; post-edge complexities are absent in the glass spectra, which is consistent with a resonance cause for the features; a lack of long range structure in the melts will dampen the resonances. There is no systematic change in spectral shape as the Ni:S ratio changes from much greater than 1 in sample B190906p5 to much less than 1, e.g. in sample B231205p5 (Fig. 8c). Nor do the spectra depend on the oxgen and sulphur fugacities at which the glasses formed (not shown).

The difference between centres of the pre-edge peak and that of the NiO standard are 0.4 534 to 0.6 eV (Table 6). This is consistent with [4] coordinated Ni in the silicate glass according 535 to the calibration of Farges et al. (2001). However, absolute energies of the pre-edge peaks 536 are slightly but systematically different to those recorded by Farges et al. (2001). The pre-537 edge peaks from this study are centred at 8332.5 to 8332.7 eV, relative to the pre-edge peak of NiO, which is fit by a Lorentzian centred at 8333.15 eV. The results described by Farges et al. (2001) for [4] coordinated glasses are based on peaks centered around energies of around 540 8330.8 eV relative to NiO at 8331.25 eV. The source of the discrepancy is unclear, as both 541 studies are based on calibration of the Ni metal edge at 8333 eV, and non-linear calibration 542 effects would not be expected to be significant so close to the calibration point.

Fits to the EXAFS yield 3.8 ± 0.9 nearest neighbour oxygens at 1.96 ± 0.01 Å (Table 5, Fig. 9c), which is consistent with the coordination inferred from the pre-edge peak. Other fit parameters (Table 5) are also consistent with results reported by Farges et al. (2001). The possibility that Ni in 4 coordination has oxygens at two distinct Ni-O distances was also investigated but the data was insufficient to distinguish between this model and the single Ni-O distance model, so the simpler model was preferred.

5 DISCUSSION

5.1 Origin of peak at 2481 eV in nominally S^{2-} -bearing glasses

550

The similarity of the position of the peak at 2481 eV to the main peak in the sulphate 552 standards is consistent with a S^{6+} source for the 2481 eV peak. The peak is therefore 553 referred to subsequently as the sulphate peak. The proportion of the signal associated with 554 this peak, which was obtained from the spectra deconvolution described above, ranged from zero, i.e. the peak was not detectable, to 68% (Table 3). The proportion of the sulphate peak 556 was greater than 5% for six out of the 21 samples for which robust spectral deconvolution 557 was possible (Table 3). The proportion of the sulphate peak was not systematically related 558 to sample sulphur content (Fig. 10a) or the duration of sample storage (Fig. 10b), although the sample with the highest sulphur content, which also had the longest storage time (112) 560 days), had the highest proportion of sulphate peak. The presence of the sulphate peak is a 561 concern, because the experiments were performed in the sulphide stability field so primary 562 sulphate is not expected. However there are a number of unintended causes for the peak 563 that should be considered. 564

Incorrect gas mixes or inadequate gas flows during the experiment could have resulted in samples more oxidised than intended. However this is an unlikely cause as such factors would have affected all six samples in a run, rather than just one or two, as is the case observed. For example, in run B310706, sulphate peak proportions ranged from to 2% to 68%. Additionally, if sulphate were the stable redox state in the melt then sulphur solubilities would have been far smaller than those measured.

Oxidation during quench is another possibility. It has been shown that S⁶⁺ can be produced by cooling of natural Fe-bearing S²⁻-bearing glasses via exchange with the Fe²⁺-Fe³⁺ redox couple (Metrich et al., 2009). However, all the runs were quenched identically, but only some samples from some runs show oxidation, as discussed for run B310706 above. Additionally, Mn and Ni are unlikely to reduce to a valence less than +2 in order to produce sulphate. W could reduce to a lower valence but the proportion of the sulphate peak in

the W-bearing samples is as low or lower than in the Mn and Ni-bearing samples, so this is unlikely to be a major factor.

Oxidation could have occurred during sample storage. The silicate glasses are highly 579 stressed due to their rapid quenching, and are metastable under room pressure and temper-580 ature, and often crumbled to produce a large surface area that would have facilitated reaction 581 subsequent to sample production. The samples were not stored with a dessicant, and the 582 presence of water vapour would have further catalysed sample degradation and sulphate pro-583 duction. This proposal is supported by the higher, though variable, extent of oxidation in 584 samples from run B310706, which was stored for the longest time prior to sample preparation 585 and synchrotron analysis (Fig. 10a). 586

Sample preparation is known to induce oxidation (e.g. Fleet, 2005) under some circum-587 stances. It is for this reason that four samples were prepared in two ways, with one half of the sample prepared in acetone in air as normal, and the other half prepared in a glovebox 589 filled with nitrogen. There was no systematic increase in the sulphate peak as a result of the 590 preparation protocol. In two cases the nitrogen-prepared portion exhibited a higher sulphate 591 peak (e.g. Fig. 10c), in one the portion prepared under air exhibited the higher sulphate 592 signal (Fig. 10d), and in the final case the two signals were effectively identical (Fig. 4c). 593 It is therefore unlikely that the sample preparation caused the oxidation. These results in-594 dicate that oxidised portions of the samples are heterogeneously distributed. A further test 595 of sample preparation was made by preparation of one half of a sample with acetone under 596 air, and the other dry, but again under air. The spectra from these two splits of the same 597 sample were effectively identical. Additionally, preparation of fresh MnS, which is highly 598 susceptible to oxidation, by grinding under acetone did not produce S-XANES spectra with 599 a sulphate peak (Fig. 2). Contamination introduced during sample preparation is also pos-600 sible. However, this cause is unlikely because the blanks of kapton tape showed no sulphur 601 signal at all, and preparation blanks for Cl-XANES loaded with the same sample protocols 602 (Evans et al., 2008b) showed no cross contamination between samples. 603

Beam-induced oxidation to produce tetravalent sulphur has also been observed (Wilkes et al., 2008; Metrich et al., 2009) for intense microbeams such as that employed at the ESRF.

The beam at the NSRRC is much less powerful, consecutive scans on the same sample do not show any increase in the size of the sulphate peak relative to the composite peak, and there is no peak for tetravalent sulphur. It therefore seems unlikely that the S⁶⁺ was produced by beam-sample interactions.

To summarise, the sporadic nature of the oxidation and the lack of systematic oxidation 610 in any particular run argue against experimental or quench-induced oxidation. Comparison 611 of multiple preparation techniques, repeat analyses, and preparation blanks largely eliminate 612 the possibility that oxidation was caused during sample preparation and analysis. The most 613 likely cause of the oxidation is reaction of the glasses during sample storage. Such oxidation 614 does, in extreme cases, affect the shape of the composite peak (e.g. Fig. 10c,d). For this 615 reason, samples with a sulphate peak proportion greater than 20% were excluded from further analysis. A possible concern is that sulphur oxidation could affect the metal EXAFS results. 617 This is unlikely because metal: S ratios are sufficiently high that even complete oxidation of 618 the sulphur would not significantly change the proportion of metal bonded to the dominant 619 oxygen ligand. 620

5.2 Evidence for tetravalent sulphur

Sulphur XANES spectra are consistent with the presence of S²⁻ and S⁶⁺. There is no evidence for S⁴⁺ in the form of an absorption peak between sulphide and sulphate, as proposed by Metrich et al. (2002). This is as expected; Wilke et al. (2008) and Metrich et al., (2009), recently showed that S⁴⁺ can be produced by the intense photon flux experienced by samples under a synchrotron microbeam. The experiments performed for this study used a mm-size beam and lower photon fluxes than those of Metrich et al. (2002) so beam-induced changes in sample oxidation state were avoided.

629 5.3 Glass composition effects on the composite peak

The correlations between the proportions of the different components of the composite sulphur peak and the Ca:Mn ratio (Fig. 6a, b) provide evidence that the components of the
composite peak relate to melt composition in a systematic way. The simplest interpretation
would be that the lowest energy peak (2473.1 eV) is caused by Mn-S interactions in the
melt, while the higher energy peaks (2474.9 eV and 2476.2 eV) are related to Mg-S and
Ca-S interactions, and the W-peak at 2469.8 eV records W-S interactions.

Such an interpretation would, however, be over simplistic, because (a) the 2473.1 eV peak 636 is not zero in Mn-free samples and (b) the sulphide standard spectra display composite peaks 637 much broader than any of the component peaks found by the deconvolution process, so over-638 lap would be expected to occur between, for example, the Mn:S and Ca:S peak contributions. 639 Peaks in the glass are unlikely to be narrower than those in the sulphides because glasses are 640 much less ordered than crystalline materials, with a consequently broadened XANES signal. 641 It is therefore likely that the deconvolution process is unable to resolve the true complexity 642 of the XANES, and that the four or five peaks that it is possible to resolve, while useful, 643 cannot be assigned on a 1:1 basis to any particular interaction.

A more reasonable proposal is that parts of the signal accounted for by the 2473.1 eV and 2469.8 eV peaks can be attributed to Mn-S and W-S interactions respectively, while parts of the signal accounted for by the 2474.9 eV and 2476.2 eV peaks are caused by Ca-S and Mg-S interactions. Similar results, on qualitative level, were found by Metrich et al. (2009), for CMAS+FeO glass compositions. However, Metrich et al. (2009) only studied Fe-bearing melts with FeO greater than 10 wt%, and so were unable to correlate peak area with metal contents as shown in Fig. 6.

The observations can be explained in terms of the electronic configuration of alkali earth and transition metal sulphides. Transitions to the lowest unoccupied states occur at lower energies for sulphides of the more covalently bonded materials (Mn, W) and at higher energies for sulphides of the more ionically bonded materials (Mg, Ca). For Mn, the transition is likely to be to hybridised S 2p and metal unoccupied 3d orbitals (Farrell et al., 2002), while

for Ca and Mg, the low lying d orbitals are less available for hybridisation which causes
the edge energy to increase. The energy of the peak-related transition for a hypothetical
W- and S-bearing melt species is harder to determine because the valency of the W is not
known and the electronic structure of the conduction band is less well studied. However,
the position of a signal from W-S interactions at a lower energy than a signal from Mn-S
interactions is consistent with the relative positions of the main edge in the WS₂ and MnS
spectra, although the magnitude of the difference is larger for the glass spectra.

Electronic configuration is also related to the conduction properties of the equivalent 664 crystalline sulphides. The more covalent compounds such as MnS and FeS have a smaller 665 band-gap than the compounds with a greater ionic character such as MgS and CaS, and 666 it is the band gap that is reflected by the slightly higher energy of the transition for the alkali earths. However, the peak energies for the glass samples are different to those for the 668 equivalent component sulphides. This is an expected consequence of the different orbital-669 orbital interactions that occur in a silicate melt where the structure is more relaxed. Less 670 complicated spectra may record the lower signal:noise ratio attainable for the lower S con-671 centrations in the glasses, but may also relate to a lack of mid- to long-range ordering and 672 consequent resonances in the glass structure. 673

₆₇₄ 5.4 Features specific to the W-bearing glasses

The low energy peak centred at 2469.8 eV in the S XANES spectra for the CMAS+W melts 675 (e.g. Fig. 4a) is difficult to attribute to a particular transition because the research published 676 on XAS and the electronic configuration of W-S compounds (Dartigeas et al., 1996; Prouzet 677 et al., 2003) is limited and does not address this issue directly. However, the transitions 678 that are most likely to be available to S 1s core electrons in W-S melt species are 3p S σ * 679 orbitals hybridised with unoccupied levels in the W 5d orbitals. These orbitals are sparsely 680 occupied in hexavalent tungsten so it is probable that a transition to these orbitals might be 681 energetically favourable compared to the transitions discussed above for CaS and MgS melt 682 species. Regardless of the precise cause of the peak, its existence suggests that S and W do 683

interact in glasses, contrary to conclusions that would be drawn from studies of S solubility in melts, which indicate that S concentrations in silicate melts are insensitive to W content (Evans et al., 2008). However, the observation cannot be taken as unequivocal evidence for W-S melt species because it is possible that the glass may not faithfully preserve all melt features.

Additional support for the notion that W may form W-S species in melts is provided by 689 consideration of the nature of W bonding in melts. Tungstate species are characterised by 690 a high average bond-valence; four co-ordinated hexavalent tungsten has a bond-valence of 691 1.5. This means that tungsten is much more likely to attach to network modifiers, because 692 network modifiers, which have lower charge than the network formers Al and Si, can charge 693 balance the high valence W more easily. This observation explains why hexavalent tungsten 694 generally occurs in non-silicate minerals in nature, and the strong dependence of hexavalent 695 tungsten solubility on Ca concentration in melts (O'Neill et al., 2008). It may also indicate 696 that W forms species that combine network modifying cations and sulphur. However further 697 work would be required to investigate this possibility. 698

5.5 Linear combinations of sulphide spectra

710

Theoretical spectra for mixtures of the standards were calculated to explore the similarity of 700 spectra to those for mixtures of the component sulphides. Mixed spectra were calculated as 701 a molar proportion-weighted sum of the spectra for the individual standards (Fig. 11). Rela-702 tive weights were taken from expressions for S solubility as a function of metal concentration 703 by O'Neill and Mavrogenes (2002) and Evans et al. (2008a). Validity of the comparison 704 relies on the assumption that the intensity of the S K- α signal is proportional to the S 705 concentration and is not affected by the individual environment of the S. This assumption 706 has been shown to be valid over a range of sulphur contents for metal sulfide solid-solutions 707 (Farrell and Fleet, 2000; 2001; Farrell et al., 2002) and low concentrations of sulphur in coals 708 (Huffman et al., 1991). 709

XANES spectra of the silicate glasses are not similar to those constructed by linear

combination of spectra from the component sulphides (Fig. 11). This is in contrast to the 711 results of Fleet (2005), who found that XANES of S in synthetic and natural glasses could be produced by a simple combination of the sulphides of the cations present in the samples. The 713 discrepancy might be attributable to (a) lower signal:noise of the Fleet samples, which had 714 lower sulphur contents than those for which results are presented here; or (b) the combination 715 of digitised spectra for CaS and MgS standards from one beamline with measured spectra 716 from another. The CSRF has a resolution of about 0.9 eV at the S K edge (Farrell et al., 717 2002) and resolution at the NSRRC is around 0.5 eV. Thus features may appear sharper in 718 data from the NSRRC, as shown by a comparison between data for FeS in this work and in 719 Fleet (2005) (not shown). However, there may be other factors that influence the differences 720 between the spectra. Such factors have not been brought to light by the results of this study, 721 and remain mysterious. 722

5.6 Evidence for formation of metal-S melt species from metal XAS

Mn XANES spectra are very similar to those of Mn-bearing borosilicate waste glasses (McKeown et al., 2003), for which divalent Mn co-ordinated with O was inferred. Mn spectra show 725 no evidence of changes to Mn speciation with increases in S:Mn ratio, and the spectra are 726 consistent with O as the nearest neighbour to divalent Mn. Influences of S in the first coordi-727 nation shell would be expected to manifest in the form of a reduction in the edge energy and loss of the pre-peak feature at 6540.5 eV. Similarly, spectra from CMAS+Ni and CMAS+W 729 glasses show no evidence for a change in spectral features with changing sulphur content. It 730 should be possible to detect the presence of significant quantities of a Ni-S species because 731 the NiS spectra lacks the pre-edge peak shown by NiO and has a lower edge energy. Ex-732 pected changes for W would be much less because the W L3 edge is less sensitive to changes 733 in local environment than the K edges for Mn and Ni. 734

The lack of evidence for formation of metal-S melt species from the metal XAS is, at first sight, contradictory with evidence from the S XAS that supports metal-S interactions. However, the apparent contradiction can be reconciled if the potential proportions of metal-S and metal-O species are considered. The maximum proportion of the additive metals Mn,
W and Ni that could form melt species with sulphur, p_{max} , was calculated from

$$p_{\text{max}} = \frac{n_{\text{S}} - n_{\text{S,CMAS}}}{n_{\text{metal}}} \tag{1}$$

where n_i is the number of moles of melt component i per unit mass. $n_{S,CMAS}$ is the number of moles of S in the end-member CMAS composition. Values could only be calculated reliably for the experiments performed at a $\log_{10} f O_2$ of -9.6 and $\log_{10} f S_2$ of -1.91 as these were the conditions at which the S content of the CMAS composition was best constrained. Values for the Mn-bearing glasses were between 10 and 20%, those for the W-bearing glasses were lower at 1 to 2 %, and those for the Ni-bearing glasses were meaningless, because the loss of the immiscible sulphide phase led to a drop in S content to below the CMAS value.

The absolute maximum ratio of metal:S to metal:O species is therefore 0.2. At this level,
the relatively subtle effects on the metal XAS caused by the presence of metal:S species
would be difficult to distinguish, as changes in the size of the pre-edge feature of up to 20%
are within the uncertainties of the measurement technique. Thus, the metal XAS provides
information on the dominant metal complexing agent, which is oxygen, but lesser quantities
of metal-S melt species are not precluded.

5.7 Glass composition effects on metal-O coordination

Previous work has inferred that W coordination in silicate melts can be octahedral (Poirier et al., 2005a,b; Farges et al., 2006) or tetrahedral (Farges et al., 2006; O'Neill et al., 2008). EXAFS and XANES results from this work are consistent with dominantly tetrahedrally coordinated W. Work that found octahedral coordination used Si and Al-free melts (Poirier et al., 2005b), whereas the lower coordinations were found for Si- and Al-bearing glasses (Farges et al., 2006). A similar phenomena is noted for the CMAS+Ni melts. Ca-dominated melts are shown by Farges et al. (2001) to contain Ni in [5] or greater coordination whereas this study found pre-edge peaks and EXAFS analysis to indicate somewhat lower coordination of [4] to [5]. A significant difference between the melts in the two studies is the higher Si

and Al content in the CMAS composition, with a significant reduction in the proportion of oxygen that is non-bridging.

A possible explanation for these observations is that a reduction in the availability of non-bridging oxygens, which occurs when a significant proportion of oxygen is involved in the polymeric Si-Al framework of the melt, produces a lower coordination for W and Ni.

6 CONCLUSIONS

769

Sulphur spectra for CMAS, CMAS+MnO, CMAS+NiO and CMAS+WO₃ are broadly similar, consisting of a broad composite peak plus a variable sulphate signal that is attributed to sulphur oxidation during sample storage. CMAS+W exhibits an additional peak at lower energy that is tentatively attributed to transition between S 1s electrons into a hybridised W 5d - S 3p σ^* anti-bonding orbital. There is no evidence for quadravalent sulphur at the low pressures and relatively reduced conditions explored here.

Decomposition of the S K-edge XANES into four component Lorentzian peaks plus an 776 edge step reveals a systematic relationship between peak area and sample composition for 777 the CMAS+Mn melts. Mn:Ca ratios are correlated positively with the proportion of the 778 lower energy peak (2473.1 eV) and negatively with the proportion of the peak at 2474.9 779 eV. The correlations are significant at the 95% level. Such changes were not noted for the 780 CMAS+W and CMAS+Ni spectra. However, low sulphur contents in CMAS+Ni glasses 781 resulted in low resolution spectra which hinder recognition of changes such as those noted 782 for CMAS+Mn. 783

Changes in the composite peak shape are consistent with known characteristics of metalS bonding. Metals that show a greater covalent character in bonding with S, such as Mn,
have lower energy hybridised orbitals available, and transitions into these orbitals produce a
contribution to the lower energy peak component at 2473.1 eV. Metals that show metal-S
bonds with a more ionic character, such as Mg and Ca, have higher energies for the S Kedge because the hybridised orbitals are less available. Thus a large proportion of the signal

accounted for by the 2474.9 eV and 2476.2 eV peaks in the composite peak is attributed to transition of 1s S electrons into S-Ca or S-Mg hybridised orbitals. A 1:1 association between the deconvoluted peaks and the transitions is unlikely because the true complexity of the spectra is likely to have been beyond the resolving power of the deconvolution exercise. Nevertheless, the results provide good evidence for formation of Mn-S and W-S melt species in silicate glasses.

Measured spectra were demonstrated not to be any plausible combination of the component sulphide spectra. The difference between the component sulphide and glass spectra
demonstrates that the electronic configuration of S-bearing species in the glasses is quantitatively different to that in the sulphides, although the same processes and elemental properties
can be used to provide a qualitative explanation of the observed phenomena.

XANES and EXAFS analysis of Ni, W and Mn can be explained by formation of metal-O 801 melt species and there is no evidence for metal-S bonding, although such evidence would 802 be hard to distinguish, given the low S:O ratio and calculated maximum contributions of 803 metal-S species to the spectra. Mn is likely to be divalent and be coordinated by 4 to 5 804 oxygens at a distance of 2.06 Å, consistent with the results of previous work. W is likely 805 to be tetrahedrally coordinated with oxygens at a distance of around 1.75 Å, in groups 806 similar to those found in scheelite. Ni EXAFS and pre-edge peak properties are consistent 807 with 4 coordinated Ni-O melt species with Ni-O distances of 1.96 Å. Ni and W both show 808 coordinations lower than those found for previous works that utilised melt compositions with 809 a lower proportion of bridging oxygens. This observation leads to the proposal that metal:O 810 coordination is a function of the ratio of bridging to non-bridging oxygens. 811

Acknowledgments. Thanks to Dave Clark, Dean Scott and Bob Rapp and staff at the NSRRC, who provided technical input into this project. This work was performed with support from the Australian Synchrotron Research Program (ASRP), which is funded by the Commonwealth of Australia under the Major National Research Facilities Program. Charlie Mandeville and two anonymous reviewers are thanked for perceptive and thought-provoking reviews that improved the paper.

7 REFERENCES

818

- Alt, J. C., Shanks, W. C., and Jackson, M. C. (1993). Cycling of sulfur in subduction
 zones the geochemistry of sulfur in the Mariana-island arc and back-arc trough.

 Earth. Planet. Sci. Lett. 119, 477-494.
- Berry, A. J., Hack, A. C., Mavrogenes, J. A., Newville, M., and Sutton, S. R. (2006).

 A XANES study of Cu speciation in high-temperature brines using synthetic fluid inclusions. Am. Mineral. 91, 1773-1782.
- Bonnin-Mosbah, M., Metrich, N., Susini, J., Salome, M., Massare, D., and Menez, B. (2002). Micro X-ray absorption near edge structure at the sulfur and iron K-edges in natural silicate glasses. *Spectrochim. Acta, Part B* **57**, 711-725.
- Carroll, M. R. and Rutherford, M. J. (1985). Sulfide and sulfate saturation in hydrous silicate melts. J. Geophys. Res. **90**, C601-C612.
- Carroll, M. R. and Rutherford, M. J. (1988). Sulfur speciation in hydrous experimental glasses of varying oxidation-state results from measure wavelength shifts of sulfur X-rays. Am. Min. **73**, 845-849.
- Carroll, M. R. and Webster, J. D. (1994). Solubilities of sulfur, noble-gases, nitrogen, chlorine, and fluorine in magmas. *Vol. Mag.* **30**, 231-279.
- Cawthorn, R. G. (2005). Contrasting sulphide contents of the Bushveld and Sudbury

 Igneous Complexes. *Min. Dep.* **40**, 1-12.
- Dann, T. E., Chung, S. C., Huang, L. J., Juang, J. M., Chen, C. I., and Tsang, K. L. (1998). A high-performance double-crystal monochromator soft X-ray beamline. *J. Synch. Rad.* 5, 664-666.
- Dartigeas, K., Gonbeau, D., and PfisterGuillouzo, G. (1996). Core and valence spectra of
 TaS2 and WS2 Experimental and theoretical studies. *J. Chem. Soc.*, Faraday Trans.

 92, 4561-4566.

- de Hoog, J. C. M., Hattori, K. H., and Hoblitt, R. P. (2004). Oxidized sulfur-rich mafic magma at Mount Pinatubo, Philippines. *Contrib. Mineral. Petrol.* **146**, 750-761.
- Evans, K. A., Mavrogenes, J. A., and O'Neill, H. S. J. (2008a). Sulphur solubility and sulphide immiscibility in silicate melts as a function of the concentration of manganese, nickel, tungsten and copper at 1 atmosphere and 1400 C. Chem. Geol. 255, 236-249.
- Evans, K. A., Mavrogenes, J. A., O'Neill, H. S., Keller, N. S., and Jang, L. Y. (2008b).

 A preliminary investigation of chlorine XANES in silicate glasses. *Geochemistry Geo-*physics Geosystems 9.
- Farges, F., Brown, G. E., Petit, P. E., and Munoz, M. (2001). Transition elements in waterbearing silicate glasses/melts. Part I. A high-resolution and anharmonic analysis of Ni coordination environments in crystals, glasses, and melts. *Geochim. Cosmochim.* Acta 65, 1665-1678.
- Farges, F., Brown, G. E., and Rehr, J. J. (1997). Ti K-edge XANES studies of Ti coordination and disorder in oxide compounds: Comparison between theory and experiment.

 Phys. Rev. B: Condens. Matter 56, 1809-1819.
- Farges, F., Linnen, R. L., and Brown, G. E. (2006). Redox and speciation of tin in hydrous silicate glasses: A comparison with Nb, Ta, Mo and W. Can. Mineral. 44, 795-810.
- Farrell, S. P. and Fleet, M. E. (1999). Evolution of local electronic structure in cubic Mg1-xFexS by SK-edge XANES spectroscopy. *Solid State Commun.* **113**, 69-72.
- Farrell, S. P. and Fleet, M. E. (2001). Sulfur K-edge XANES study of local electronic structure in ternary monosulfide solid solution [(Fe, Co, Ni)(0.923)S]. *Phys. Chem. Miner.* **28**, 17-27.
- Farrell, S. P., Fleet, M. E., Stekhin, I. E., Kravtsova, A., Soldatov, A. V., and Liu, X. Y. (2002). Evolution of local electronic structure in alabandite and niningerite solid solutions [(Mn,Fe)S, (Mg,Mn)S, (Mg,Fe)S] using sulfur K- and L-edge XANES spectroscopy. Am. Mineral. 87, 1321-1332.

- Fincham, C. J. B. and Richardson, F. D. (1954). The behaviour of sulphur in silicate and aluminate melts. *Proc. R. Soc. London, Ser. A* **223**, 40-62.
- Fischer, T. P., Giggenbach, W. F., Sano, Y., and Williams, S. N. (1998). Fluxes and sources of volatiles discharged from Kudryavy, a subduction zone volcano, Kurile Islands. Earth. Planet. Sci. Lett. 160, 81-96.
- Fleet, M. E. (2005). Xanes spectroscopy of sulfur in earth materials. *Can. Mineral.* **43**, 1811-1838.
- Fleet, M. E., Liu, X. Y., Harmer, S. L., and King, P. L. (2005). Sulfur K-edge xanes spectroscopy: Chemical state and content of sulfur in silicate glasses. *Can. Mineral.*43, 1605-1618.
- Fulton, J. L., Pfund, D. M., Wallen, S. L., Newville, M., Stern, E. A., and Ma, Y. (1996).

 Rubidium ion hydration in ambient and supercritical water. *J. Chem. Phys.* **105**,

 2161-2166.
- Hazen, R. M., Finger, L. W., and Mariathasan, J. W. E. (1985). High-pressure crystalchemistry of scheelite-type tungstates and molybdates. *J. Phys. Chem. Solids* **46**, 253-263.
- Huffman, G. P., Mitra, S., Huggins, F. E., Shah, N., Vaidya, S., and Lu, F. L. (1991).
 Quantitative-Analysis of All Major Forms of Sulfur in Coal by X-Ray Absorption
 Fine-Structure Spectroscopy. Energy Fuels 5, 574-581.
- Kravtsova, A. N., Stekhin, I. E., Soldatov, A. V., Liu, X., and Fleet, M. E. (2004). Electronic structure of MS (M=Ca,Mg,Fe,Mn): X-ray absorption analysis. *Phys. Rev. B:*Condens. Matter **69**.
- Lee, I. S. and Ripley, E. M. (1995). Genesis of Cu-Ni sulfide mineralization in the South Kawishiwi intrusion, Spruce Road area, Duluth complex, Minnesota. *Can. Mineral.* 33, 723-743.

- Li, C., Ripley, E. M., and Mathez, E. A. (2003). The effect of S on the partitioning of Ni between olivine and silicate melt in MORB. *Chem. Geol.* **201**, 295-306.
- Li, C. S. and Ripley, E. M. (2005). Empirical equations to predict the sulfur content of mafic magmas at sulfide saturation and applications to magmatic sulfide deposits.

 Min. Dep. 40, 218-230.
- Li, D., Bancroft, G. M., Kasrai, M., Fleet, M. E., Feng, X. H., and Tan, K. (1995). S-K-edge and L-edge X-ray absorption spectroscipy of metal sulfides and sulfates - applications in mineralogy and geochemistry. *Can. Mineral.* **33**, 949-960.
- Li, D., Bancroft, G. M., Kasrai, M., Fleet, M. E., Yang, B. X., Feng, X. H., Tan, K., and Peng, M. S. (1994). Sulfur K-edge and L-edge X-ray-absorption spectroscopy of sphalerite, chalcopyrite and stannite. *Phys. Chem. Miner.* **20**, 489-499.
- Liu, W., Etschmann, B., Foran, G., Shelley, M., and Brugger, J. (2006). A XANES study of Zn and Fe(II) chloride complexes in hypersaline brines. *Geochim. Cosmochim. Acta* **70**, A364-A364.
- McKeown, D. A., Kot, W. K., Gan, H., and Pegg, I. L. (2003). X-ray absorption studies of manganese valence and local environment in borosilicate waste glasses. *J. Non-Cryst.* Solids **328**, 71-89.
- Metrich, N., Berry, A. J., O'Neill, H. S. C., and Susini, J. (2009). The oxidation state of sulfur in synthetic and natural glasses determined by X-ray absorption spectroscopy. *Geochim. Cosmochim. Acta* 73, 2382-2399.
- Metrich, N., Bonnin-Mosbah, M., Susini, J., Menez, B., and Galoisy, L. (2002). Presence of sulfite (S-IV) in arc magmas: Implications for volcanic sulfur emissions. *Geophys.*Res. Lett. 29.
- Metrich, N. and Clocchiatti, R. (1996). Sulfur abundance and its speciation in oxidized alkaline melts. *Geochim. Cosmochim. Acta* **60**, 4151-4160.

- Moretti, R. and Ottonello, G. (2005). Solubility and speciation of sulfur in silicate melts:

 The Conjugated Toop-Samis-Flood-Grjotheim (CTSFG) model. Geochim. Cosmochim.
- Acta **69**, 801-823.
- Naldrett, A. J. (1999). World-class Ni-Cu-PGE deposits: key factors in their genesis. *Min.*10 Dep. 34, 227-240.
- Nietubyc, R., Sobczak, E., and Attenkofer, K. E. (2001). X-ray absorption fine structure study of manganese compounds. *J. Alloys Compd.* **328**, 126-131.
- O'Neill, H. S. C., Berry, A. J., and Eggins, S. M. (2008). The solubility and oxidation state of tungsten in silicate melts: implications for the comparative chemistry of W and Mo in planetary differentiation processes. *Chem. Geol.* **255**, 346-359.
- O'Neill, H. S. C. and Mavrogenes, J. A. (2002). The sulfide capacity and the sulfur content at sulfide saturation of silicate melts at 1400 degrees C and 1 bar. J. Pet. 43, 1049-1087.
- Paris, E., Giuli, G., Carroll, M. R., and Davoli, I. (2001). The valence and speciation of sulfur in glasses by X-ray absorption spectroscopy. *Can. Mineral.* **39**, 331-339.
- Pina, R., Lunar, R., Ortega, L., Gervilla, F., Alapieti, T., and Martinez, C. (2006). Petrology and geochemistry of mafic-ultramafic fragments from the Aguablanca Ni-Cu ore breccia, southwest Spain. *Econ. Geol.* **101**, 865-881.
- Poirier, G., Messaddeq, Y., Ribeiro, S. J. L., and Poulain, M. (2005). Structural study of tungstate fluorophosphate glasses by Raman and X-ray absorption spectroscopy. *J.* Solid State Chem. 178, 1533-1538.
- Poirier, G. L., Cassanjes, F. C., Messaddeq, Y., Ribeiro, S. J. L., Michalowicz, A., and Poulain, M. (2005). Local order around tungsten atoms in tungstate fluorophosphate glasses by X-ray absorption spectroscopy. *J. Non-Cryst. Solids* **351**, 3644-3648.
- Prouzet, E., Heising, J., and Kanatzidis, M. G. (2003). Structure of restacked and pillared
 WS2: An X-ray absorption study. Chem. Mater. 15, 412-418.

- Ravel, B. and Newville, M. (2005). ATHENA, ARTEMIS, HEPHAESTUS: data analysis for X-ray absorption spectroscopy using IFEFFIT. J. Synch. Rad. 12, 537-541.
- Soldatov, A. V., Kravtsova, A. N., Fleet, M. E., and Harmer, S. L. (2004). Electronic structure of MeS (Me = Ni, Co, Fe): x-ray absorption analysis. *Journal of Physics-Condensed Matter* **16**, 7545-7556.
- Taylor, J. C. (1991). Computer programs for standardless quantitative analysis of minerals using the full powder diffraction pro?file. *Powder Diffr.* **6**, 2-9.
- Wallace, P. and Carmichael, I. S. E. (1992). Sulfur in basaltic magmas. Geochim. Cosmochim. Acta 56, 1863-1874.
- Wang, C. Y. and Zhou, M. F. (2006). Genesis of the Permian Baimazhai magmatic Ni-Cu-(PGE) sulfide deposit, Yunnan, SW China. *Min. Dep.* **41**, 771-783.
- Wilke, M., Jugo, P. J., Klimm, K., Susini, J., Botcharnikov, R., Kohn, S. C., and Janousch,
 M. (2008). The origin of S⁴⁺ detected in silicate glasses by XANES. Am. Mineral. 93,
 235-240.

959 TABLES

Table 1: Summary of run conditions

Run	Additive	Additive	$Log_{10} fO_2$	$Log_{10} fS_2$	$Log_{10} fSO_2$	$\Delta \mathrm{QFM}$
	metal	conc (wt%)				
B030106	Mn	1-10	-9.60	-1.91	-3.06	-3.2
B310706	Mn	1-10	-9.60	-1.91	-3.06	-3.2
B020806	Mn	1-10	-7.18	-0.84	-0.61	-0.7
B111006	Mn	0-1	-7.59	-2.80	-1.50	-1.2
B151106	Mn	0-1	-9.60	-1.91	-3.06	-3.2
B301205	W	0-1	-9.60	-1.91	-3.06	-3.2
B240306	W	1-10	-9.60	-1.91	-3.06	-3.2
B220806	W	1-10	-9.60	-1.91	-3.06	-3.2
B230806	W	1-10	-7.18	-1.85	-0.61	-0.7
B150906	W	0-1	-9.60	-1.91	-3.06	-3.2
B200906	W	1.5-10	-3.09	-10.85	-0.52	3.3
B141106	W	1-10	-9.60	-1.91	-3.06	-3.2
B231205	Ni	0-1	-9.60	-1.91	-3.06	-3.2
B220306	Ni	1-10	-9.60	-1.91	-3.06	-3.2
B070406	Ni	10-30	-9.60	-1.91	-3.06	-3.2
B120406	Ni	1-10	-7.18	-1.85	-0.61	-0.7
B010806	Ni	0-1	-9.60	-1.91	-3.06	-3.2
B030806	Ni	0-1	-7.18	-1.85	-0.61	-0.7
B140906	Ni	0-1	-9.60	-1.91	-3.06	-3.2
B190906	Ni	1-10	-3.09	-11.88	-1.52	3.3

Sample	Si(wt%)	Al(wt%)	Mg(wt%)	Ca(wt %)	Fe(wt%)	S(wt%)	$\operatorname{Mn}(\operatorname{wt}\%)$	O(wt%)	Total(wt%)	Edge (eV)
B030106p2 ‡	19.93(8)	6.64(4)	5.07(3)	24.54(5)	0.73(1)	0.224(1)	0.381(3)	42.31(1)	99.9(2)	n.a.
B310706p1 *	20.3(1)	(6.98(6))	5.43(5)	25.32(9)	0.018(1)	0.196(3)	bdl	43.2(2)	101.5(5)	2471.3
B310706p2*	19.5(1)	6.91(2)	5.35(7)	24.91(1)	0.04(2)	0.332(3)	2.07(2)	42.8(1)	102.0(2)	2470.4
B310706p3 *	18.9(2)	(6.58(9))	5.17(7)	24.2(2)	0.043(9)	0.480(2)	3.74(5)	42.0(4)	101.1(8)	2470.5
B310706p4*	18.4(1)	6.44(2)	5.06(2)	23.71(7)	0.05(1)	0.726(6)	6.03(2)	42.0(1)	102.4(2)	2470.4
B310706p5	18.2(1)	6.3(1)	4.87(7)	23.24(7)	0.07(3)	0.879(4)	7.15(2)	41.8(1)	102.5(3)	2470.6
B310706p6 *	17.6(2)	6.10(9)	4.76(7)	22.36(6)	0.06(2)	1.117(8)	8.74(3)	41.2(2)	102.0(4)	2470.9
B020806p1	17.6(2)	6.19(3)	4.75(1)	22.62(5)	lpq	0.088(3)	9.008(1)	40.5(2)	100.8(4)	2470.8
B020806p2*	18.40(5)	6.39(2)	4.993(3)	23.76(8)	0.02(2)	0.053(2)	6.10(3)	41.25(9)	101.0(2)	2471.4
B020806p3 *	17.93(5)	6.3(1)	4.92(8)	23.29(3)	0.040(7)	0.12(1)	7.27(2)	40.8(2)	100.7(3)	2470.5
B020806p4	19.4(1)	6.76(5)	5.29(9)	25.07(8)	pql	0.024(2)	2.122(7)	42.3(2)	101.0(3)	2470.5
B020806p5	19.12(5)	6.69(4)	5.09(8)	24.62(6)	lpq	0.039(2)	3.902(7)	42.10(5)	101.6(1)	2471
B020806p6 ‡	18.1(1)	6.3(1)	4.87(7)	23.3(1)	pdl	0.122(4)	7.24(2)	41.0(2)	101.0(4)	2471.2
B111006p2	20.2(5)	7.6(7)	4.9(8)	24.(1)	lpq	0.019(7)	1.02(9)	43.0(3)	101.2(2)	n.a.
B151106p1	21.08(9)	7.08(4)	5.25(2)	22.03(7)	0.048(7)	0.229(2)	0.379(5)	42.93(9)	99.1(1)	2471.2
B151106p4	20.84(5)	(9.95(6)	5.20(2)	21.93(6)	0.05(2)	0.274(5)	1.070(6)	42.7(1)	99.2(2)	2471.4
B151106p5	21.19(4)	7.02(7)	5.22(6)	22.02(6)	0.096(8)	0.219(7)	0.235(3)	42.94(4)	99.0(1)	2471.2
B151106p6	20.89(7)	6.95(3)	5.18(6)	21.82(6)	0.065(4)	0.251(4)	0.840(7)	42.62(9)	98.7(1)	2471.5
*: S XANES depicted	depicted									
‡: Mn XANES depicted	S depicted									

Table 2b: Electron microprobe analyses of W-bearing samples	ctron micre	oprobe anal	lyses of W-b	earing samt	səles					
Sample	Si(wt%)	Al(wt%)	Mg(wt%)	Ca(wt%)	Fe(wt%)	S(wt%)	W(wt%)	O(wt%)	Total(wt%)	Edge (eV)
B301205p3 ‡	19.8(1)	(9.65(6)	5.15(5)	24.38(2)	0.486(7)	0.208(2)	0.740(6)	42.1(2)	99.6(3)	n.a.
B240306p1	20.60(6)	7.32(6)	5.55(1)	15.53(3)	0.06(1)	0.075(4)	7.74(4)	41.97(9)	98.9(2)	n.a.
B220806p2	21.59(1)	7.63(9)	5.80(5)	16.03(5)	0.048(1)	0.064(4)	5.86(4)	43.23(1)	100.3(2)	2469
B220806p3	21.1(1)	7.39(7)	5.69(5)	15.50(7)	0.05(1)	0.068(7)	7.6(1)	42.6(1)	100.0(3)	2468.8
$\mathrm{B}220806\mathrm{p}4$	23.08(5)	8.14(5)	6.14(7)	16.77(2)	0.020(7)	pql	2.23(3)	44.92(9)	101.4(2)	n.a.
$\mathrm{B}220806\mathrm{p}5$	22.8(2)	7.89(3)	(9)60.9	19.73(2)	0.06(2)	0.073(5)	0.223(7)	45.0(2)	101.9(4)	2469
$\mathrm{B}220806\mathrm{p}6$	20.51(9)	7.32(7)	5.6(1)	15.1(2)	0.05(1)	0.071(3)	9.27(5)	42.1(2)	100.1(6)	2470.6
B230806p5 ‡	22.8(2)	8.00(8)	6.2(1)	16.6(1)	pql	pql	2.34(4)	44.4(4)	100.3(9)	n.a.
$B200906p5 \ddagger$	19.2(1)	6.61(8)	5.31(7)	24.46(4)	pql	bdl	2.67(4)	41.7(1)	100.0(3)	n.a.
B150906p1	19.7(2)	7.0(3)	5.5(5)	25.1(3)	pql	0.3(1)	0.70(9)	42.9(2)	101.4(4)	2471.1
m B150906p2	20.1(1)	(9)86.9	5.36(6)	25.1(1)	0.040(5)	0.200(3)	0.28(3)	42.9(2)	101.0(4)	2471.2
m B150906p3	19.7(1)	6.83(5)	5.39(9)	24.95(7)	0.18(2)	0.206(4)	0.89(3)	42.5(2)	100.7(3)	2471.4
m B150906p4	19.9(1)	(9)68.9	5.4(1)	25.14(9)	0.17(2)	0.202(5)	0.49(5)	42.7(2)	101.0(3)	2471.2
m B150906p5	21.16(3)	7.32(1)	5.80(8)	22.80(9)	0.163(7)	0.134(2)	pql	43.7(1)	101.1(3)	2471.4
B141106p1 *	18.4(3)	6.6(2)	5.16(5)	22.0(3)	0.07(2)	0.14(2)	6.4(2)	40.9(7)	100.(2)	2468.9
B141106p2*	19.5(2)	6.8(4)	5.2(3)	24.8(8)	pql	0.10(2)	2.4(2)	42.1(6)	101.(1)	2469.0
$\rm B141106p3$	20.2(7)	6.7(4)	5.1(8)	23.5(8)	0.01(2)	0.05(4)	3.8(1)	42.7(6)	102.(2)	2469.5
$\mathrm{B}141106\mathrm{p}4$	17.57(8)	6.22(4)	5.27(5)	20.55(9)	0.05(3)	0.18(5)	10.1(4)	40.1(2)	100.0(6)	2470.6
*: S XANES depicted	depicted									
‡: W XANES depicted	depicted									
n.a.: not available	lable									

Table 2c: Electron microprobe analyses of	ctron micro $\frac{c:(x+0)}{c:(x+0)}$	probe analy		Ni-bearing samples	SS (707)	(20,000)	N: (2040)	(10+111)	Toto1(22407)	[] (11)
Sample	31(Wt 70)	A1(WU70)	Mg(Wt/70)	Ca(Wt 70)	re(wt/0)	S(WL%)	INI(WL70)	O(Wt%)	Total(Wt70)	rage (ev)
$B231205p1 \ddagger$	19.81(8)	6.64(2)	5.14(3)	24.96(2)	0.374(7)	0.214(2)	0.016(2)	42.2(1)	99.3(2)	n.a.
B231205p5	19.8(2)	6.74(5)	5.15(7)	24.82(3)	0.521(4)	0.214(2)	pql	42.3(2)	99.5(4)	n.a.
B220306p3	19.5(1)	7.22(9)	5.5(1)	24.63(3)	0.10(1)	0.188(5)	0.021(3)	42.3(3)	99.5(7)	n.a.
$\mathrm{B070406p6}$	22.717(8)	8.46(6)	6.5(1)	16.89(8)	0.04(2)	0.058(5)	0.033(2)	44.50(6)	99.2(2)	n.a.
$B120406p6 \ddagger$	18.7(9)	6.9(3)	5.3(2)	24.6(2)	0.02(2)	0.022(3)	0.312(9)	41.(1)	97.(3)	n.a.
B010806p1 *	20.04(8)	7.00(9)	5.50(5)	25.80(9)	0.04(1)	0.180(4)	0.016(4)	43.2(1)	101.8(3)	2470.5
B010806p2	23.73(4)	8.28(4)	6.47(7)	17.26(3)	0.03(1)	pql	0.132(8)	45.64(7)	101.7(1)	n.d.
B010806p4	23.6(1)	7.9(2)	6.4(2)	17.37(9)	0.018(7)	pql	0.091(4)	45.2(2)	100.7(4)	2469
B010806p5	20.1(1)	7.02(1)	5.47(7)	25.53(9)	0.032(9)	0.021(9)	0.10(4)	43.1(2)	101.5(4)	2468.5
B030806p2	20.2(1)	7.04(1)	5.4(1)	25.71(5)	0.05(2)	0.021(8)	0.38(4)	43.2(3)	101.9(6)	n.a.
B030806p5	20.0(1)	(6.98(9)	5.4(2)	25.3(2)	0.02(2)	0.016(2)	0.354(5)	42.7(4)	100.7(9)	n.a.
B140906p4 *	20.2(1)	(2.96(7)	5.3(1)	25.68(6)	0.08(1)	0.203(3)	0.025(6)	43.2(2)	101.6(5)	2471.2
$\mathrm{B}140906\mathrm{p}5$	20.18(9)	7.00(2)	5.28(6)	25.641(8)	0.104(7)	0.199(3)	pql	43.16(8)	101.6(1)	2471.2
B190906p5	19.8(1)	6.84(8)	5.40(4)	24.1(1)	0.020(5)	pql	2.473(7)	42.4(2)	101.0(4)	n.a.
*: S XANES depicted	depicted									
‡: Ni XANES depicted	depicted									
n.a.: not available	lable									

Table 3: Resu	ults of S X ℓ	Table 3: Results of S XANES deconvolution	tion								ı
Sample	Additive	Additive	S (wt %)	Metal:Ca	$\operatorname{Storage}$	P(W)	P(2473.1))	P(2474.9)	P(2476.2)	P(2481.4)	χ^2
		metal (wt %)		(molar)	(Days)						
B310706p1	Mn	0	0.196	0.0000	113	n.a.	0.11(3)	0.34(7)	0.36(5)	0.19(2)	1.548
B310706p2	Mn	2.07	0.332	0.0606	113	n.a.	0.33(3)	0.21(5)	0.44(4)	0.020(6)	0.355
B310706p3	Mn	3.74	0.48	0.1127	113	n.a.	0.26(3)	0.26(6)	0.36(4)	0.11(1)	0.396
$\rm B310706p4$	Mn	6.03	0.726	0.1855	113	n.a.	0.31(3)	0.19(4)	0.44(4)	0.06(1)	1.553
B310706p6	Mn	8.74	1.117	0.2852	113	n.a.	0.03(4)	0.22(7)	0.06(4)	0.68(6)	1.124
B151106p1	Mn	0.379	0.229	0.0126	16	n.a.	0.10(3)	0.22(6)	0.27(5)	0.41(5)	0.283
B151106p1N	Mn	0.378	0.229	0.0125	16	n.a.	0.13(4)	0.45(10)	0.40(7)	0.02(1)	0.284
$\rm B151106p4$	Mn	1.07	0.274	0.0356	16	n.a.	0.19(4)	0.40(9)	0.41(6)	n.a.	0.289
B151106p5	Mn	0.235	0.219	0.0078	15	n.a.	0.15(4)	0.43(9)	0.40(6)	0.03(1)	0.261
$\mathrm{B}151106\mathrm{p}6$	m Mn	0.84	0.251	0.0281	16	n.a.	0.16(4)	0.44(10)	0.40(7)	n.a.	0.24
B150906p1	W	0.7	0.3	0.0203	20	n.a.	0.14(4)	0.36(9)	0.45(7)	0.05(2)	0.784
m B150906p2	W	0.28	0.2	0.0081	20	n.a.	0.15(3)	0.38(8)	0.47(6)	n.a.	1.865
$\mathrm{B}150906\mathrm{p}3$	W	0.89	0.206	0.0260	71	n.a.	0.17(4)	0.34(8)	0.46(6)	0.04(1)	0.636
m B150906p4	W	0.49	0.202	0.0142	71	n.a.	0.22(3)	0.25(6)	0.52(6)	0.01(2)	1.536
$\rm B150906p5$	M	0	0.134	0.0000	71	n.a.	0.16(3)	0.34(8)	0.49(6)	n.a.	0.742
B141106p1	W	6.4	0.14	0.0634	11	0.07(1)	0.11(4)	0.3(1)	0.49(8)	n.a.	1.401
B141106p2a	W	2.4	0.1	0.0211	11	0.09(1)	0.13(3)	0.24(7)	0.46(6)	0.09(2)	0.588
B141106p2b	W	2.4	0.1	0.0223	11	0.08(1)	0.07(3)	0.32(9)	0.43(8)	0.10(2)	0.952
B140906p4	Ni	0.03	0.203	0.0008	71	n.a.	0.15(3)	0.38(8)	0.47(6)	n.a.	0.639
B140906p4D	$ m N_i$	0.03	0.203	0.0008	71	n.a.	0.29(3)	0.18(5)	0.53(5)	n.a.	0.693
B140906p5	Ni	0.03	0.199	0.0008	71	n.a.	0.21(3)	0.33(7)	0.46(5)	0.001(4)	1.494
n.s.: not significant	ificant										

P(xxxx) indicates the proportion of the signal accounted for by the peak at xxxx eV

Table 4: Correlation coefficients between peak areas and sample composition for Mn-bearing samples

		Mn	Ca	S	Mn:Ca
		wt $\%$	wt %	wt $\%$	molar
961	P(2473.1)	0.85	n.s.	0.8	0.84
	P(2474.9)	-0.85	n.s.	-0.81	-0.84
	P(2476.2)	n.s.	n.s.	n.s.	n.s.

P(xxxx) indicates the proportion of the signal accounted for by the peak at xxxx eV

Table 5: Results of EXAFS fitting. Table headings defined in section 3.3

	n	E_0	r(Å)	σ^2	R factor
Mn	5(1)	6540(4)	2.06(2)	0.009(3)	0.007
W	3(1)	10204(3)	1.77(1)	0.001(1)	0.023
Ni	3.8(9)	8345*	1.96(1)	0.007(4)	0.036

^{*:} Energy set to maximum first derivative on edge

Table 6: Ni pre-edge peak centroid energies

	Sample	$E_{\rm pre-edge,sample}\text{-}E_{\rm pre-edge,NiO}$	1σ
		(eV)	(eV)
	B231205p1	0.48	0.07
	B220306p6	0.49	0.05
963	B120406p6	0.56	0.05
	B070406p6	0.41	0.16
	B190906p5	0.49	0.04
	B030806p5	0.5	0.04
	B030806p2	0.45	0.05
	NiO	0	n.a.

8 FIGURE CAPTIONS

- Figure 1: a: Schematic comparison of the transitions that result in the sulphur K edge for compounds with a greater ionic bonding character, such as MgS, and those with a greater covalent bonding character, such as FeS; b: Summary of edge positions, defined as the maximum first derivative, versus oxidation state for standards measured in this study.
- Figure 2. S K edge XANES spectra for sulphide standards. Peak at 2481eV indicated by the dotted line is spurious sulphate caused by oxidation of the samples prior to sample preparation and analysis. Spectra for CaS and MgS are digitised from Farrell and Fleet (2000). All other spectra were measured at the NSRRC.
- Figure 3. S K edge XANES spectra for sulphate standards. All spectra were measured at the NSRRC.
- Figure 4. S K edge XANES spectra for selected synthetic silicate glasses. Peak at 2481eV 976 is spurious sulphate caused by oxidation of the samples prior to sample preparation 977 and analysis. a: Effect of changes in additive metal on XANES for CMAS (sam-978 ple B310706p1), CMAS+Ni (sample B010806p1), CMAS+Mn (sample B310706p2), 979 CMAS + W (samples B141106p1). b: Effect of changes in additive metal content 980 on XANES for CMAS+Mn compositions for B310706p3 (3.7 wt% Mn), B310706p4 981 (6.0 wt% Mn) and B310706p6 (8.7 wt% Mn); c: Effect of changes in sample prepa-982 ration method for sample B140906p4. See also Fig. 10c,d; d: Effect of oxidation 983 state of experiment on S K edge XANES spectra for, from the top down, B020806p2, 984 B310706p6, B020806p3, B310706p2. Both experiments were performed in the sulphide 985 stability field. Samples chosen to illustrate the range of sulphate signal produced by 986 samples from within the same run. 987
- Figure 5. Example of results of S XANES peak deconvolution exercise. Thicker line is the measured XANES. Thin lines are the deconvoluted peak components and fit XANES.

 a: B310706p1 (CMAS); b: B310706p4 (CMAS+Mn); c: B141106p2b (CMAS+W).

Figure 6. Results of S XANES peak deconvolution exercise. Spectra with proportion of signal accounted for by the 2481.4 eV (sulphate-related peak) greater than 0.2 and with low signal:noise are omitted. a: Relationship between the proportion of the signal accounted for by the peak at 2473.1 eV and the molar Mn:Ca ratio; b: Relationship between the proportion of the signal accounted for by the peak at 2474.9 eV and molar Mn:Ca ratio; c: Relationship between the proportion of the signal accounted for by the peak at 2473.1 eV and W:Ca ratio; d: Relationship between the proportion of the signal accounted for by the peak at 2476.9 eV and W:Ca ratio.

Figure 7. Metal XANES spectra for the metal standards. a: Mn standards; b: W standards; c: Ni standards. All metal standards were measured on BL17C at the NSRRC.

Figure 8. Metal XANES spectra for the synthetic glasses. a: Mn K edge spectra in order of decreasing Mn:S ratio; b: W L3 edge spectra in order of decreasing W:S ratio. A indicates the intense white line, B indicates the post-edge feature discussed in the text; c: Ni K edge spectra in order of decreasing Ni:S ratio.

Figure 9. Results of metal EXAFS fitting. R is distance in Angstroms. $|\chi(R)|$ is the magnitude of the Fourier-transformed signal as a function of R. The shape, position and size of the peak provides information on the nearest neighbours to the metal probed by the EXAFS measurements. Solid line is data, dashed line is the fit. a: average of Mn-bearing glasses. b: Average of W-bearing glasses; c: average of Ni-bearing glasses.

Figure 10. Further information on the sullphate peak at 2481.4 eV. a: relationship between sample storage duration and the proportion of the signal accounted for by the 2481.4 eV peak; b: relationship between sample sulphur content and the proportion of the signal accounted for by the 2481.4 eV peak; c: comparison of XANES for the two portions of B140906p1 prepared under air and in a nitrogen-filled glovebox. The extension N indicates preparation in the glovebox; d: comparison of XANES for the two portions of B151106p1 prepared under air and in a nitrogen-filled glovebox. The extension N indicates preparation in the glovebox. Dashed line indicates position of sulphate peak.

Figure 11. Comparison of S K edge XANES spectra with sums of sulphide spectra for a:

B310706p1 (CMAS); b: B310706p6 (CMAS+Mn); c: B141106p1 (CMAS+W).

9 FIGURES

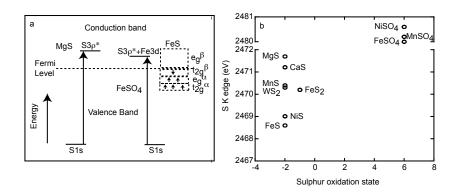


Figure 1:

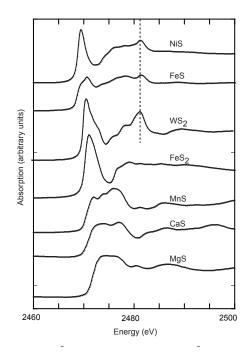


Figure 2:

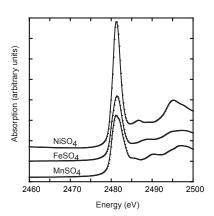


Figure 3:

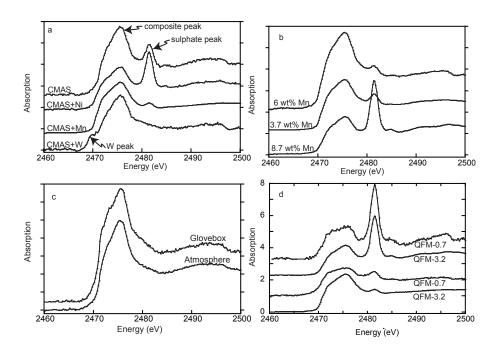


Figure 4:

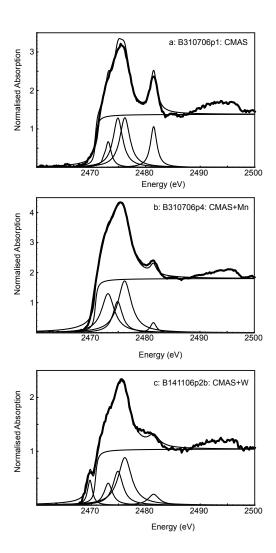


Figure 5:

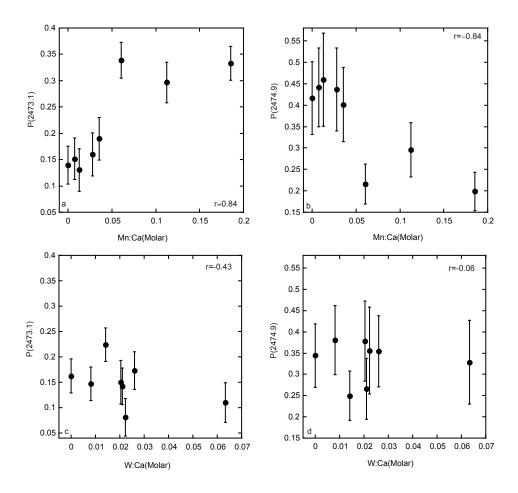


Figure 6:

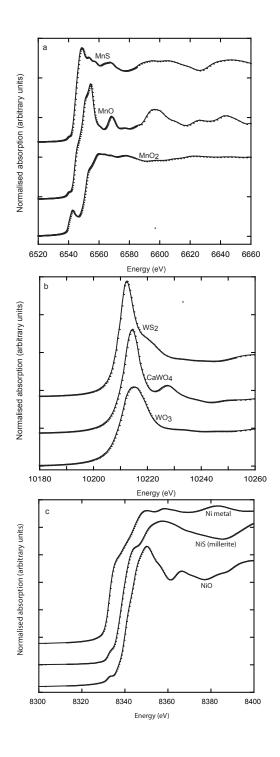


Figure 7:

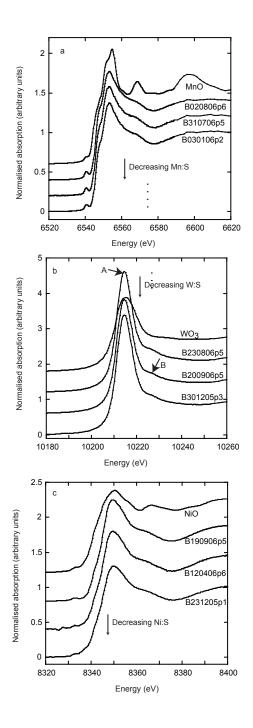


Figure 8:

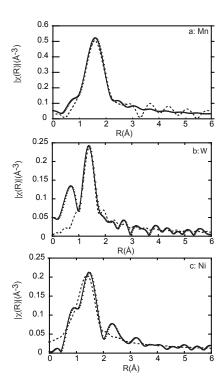


Figure 9:

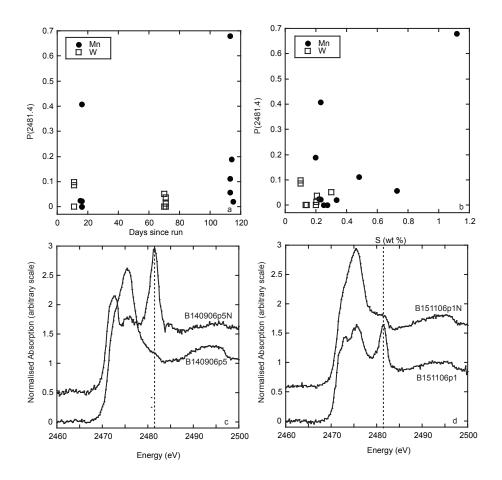


Figure 10:

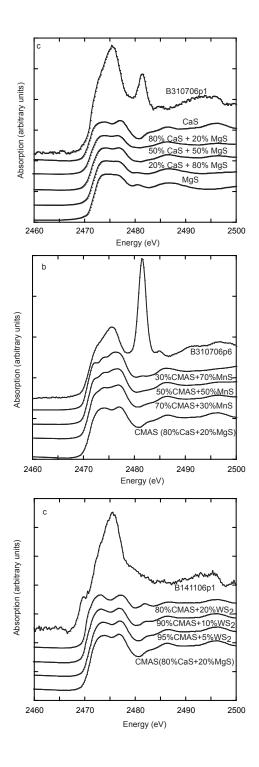


Figure 11: