

1 Understanding vegetation variability and their “hotspots” within Lake
2 Victoria Basin (LVB: 2003-2018)

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6 **Abstract**

7 Lake Victoria’s surface area has recently been shown to have shrunk by 0.3% compared to its
8 1984 value, a decline that has been associated with climatic as well as anthropogenic factors.
9 Climatic factors include, e.g., reduced rainfall, which impacts not only on the lake’s water level
10 but also on the basin’s vegetation that forms the lake’s catchment. Understanding the loca-
11 tions of vegetation changes and the driving forces of such changes, therefore, is of most critical
12 importance to major stakeholders regarding environmental management, policies and planning.
13 For Lake Victoria Basin (LVB; Kenya, Uganda, Tanzania, Rwanda and Burundi), human devel-
14 opment and climatic variability/change have subjected the region to significant changes in its
15 vegetation characteristics whose spatio-temporal patterns are, however, not well understood.
16 To understand this variability in vegetation for the period 2003-2018, this study employs the
17 use of remotely sensed MODIS (Moderate Resolution Imaging Spectroradiometer), CHIRPS
18 (Climate Hazards Group InfraRed Precipitation with station data) precipitation data, Google
19 Earth Pro imagery, Gravity Recovery and Climate Experiment (GRACE)-based Mascon’s to-
20 tal water storage (TWS) products and the statistical PCA (Principal Component Analysis).
21 The study aims at determining (i) “significant hotspots”, i.e. vegetation areas within the LVB
22 largely impacted, and (ii), the extent of which anthropogenic and climatic variability have con-
23 tributed to the “hotspots” formation. The results indicate a total of 8 hotspots; 5 in Uganda
24 and 1 each in Kenya, Tanzania and Rwanda. Google Earth Pro imagery of all the hotspots
25 show the changes in anthropogenic processes as the primary driver for the long-term changes in
26 vegetation characteristics. Conversely, the analysis of PCA and Mascon’s TWS concluded that
27 only the Tanzanian hotspot may have been driven somewhat by climate variability. Climate
28 variability is understood to be the driver in short-term vegetation changes while the long-term
29 effects are driven primarily by human influence.

30 *Keywords:* CHIRPS; Lake Victoria; Vegetation change, NDVI; Climatic variability;
31 Anthropogenic

32 1. Introduction

33 Studies have shown that alterations to the hydrological characteristics of freshwater areas
34 correlate with changes in shoreline vegetation (e.g., (Hudon , 1997; Nilsson and Berggren , 2000;
35 Baldwin et al., 2001; New and Xie , 2008)). Decreases in water-levels can allow vegetation types
36 and other plant species the freedom to regenerate (Hill et al., 1998; Zhang et al., 2017). Con-
37 versely, anthropogenic processes that act to clear vegetation for urbanisation and development
38 can indirectly lead to decreased water-levels of freshwater bodies that lie in proximity to where
39 such stressors occur (Sand-Jensen et al., 2000; Zhang et al., 2017). Such water bodies provide
40 important socio-economic services for the regions in which they are located, hence why they are
41 sourced for extraction. For example, Jeppesen et al. (2012); Zhang et al. (2017); Sand-Jensen et
42 al. (2000) investigate the extent of decline of flora species macrophyte around streams that are
43 in close proximity to cultivated and urbanised regions and conclude that there are significant
44 overall decrease of species richness for each of the streams they surveyed. Yet the converse
45 is also true; anthropogenic activities such as use of fertilizers and industrial wastes also indi-
46 rectly fuel increase in microphytes as demonstrated by Coladello et al., (2020) who studied
47 macrophytes' abundance changes in eutrophicated tropical reservoir of Salto Grande in Brazil.

48 For the Lake Victoria Basin (LVB; 31°39' - 34°53' E and 0°20' - 3° N) that houses the
49 second largest freshwater lake in the world (Mati et al., 2008), Awange et al. (2019a) employed
50 medium-resolution remotely sensed data of Landsat (5, 7 & 8) and Sentinel-2 images for the
51 period 1984-2018 and indicated that the surface area of the lake had shrunk significantly in
52 recent decades at a rate of 5.97 km^2 /year (from manual digitisation) and 5.27 km^2 /year (from
53 Modification of Normalized Difference Water Index (MNDWI)). Four areas that were identified
54 as "hotspots", which had major influences in the overall reduction of Lake Victoria's surface
55 area, i.e., Birinzi, Uganda - 33.9 km^2 ; Mwanza Gulf, Tanzania - 21.5 km^2 ; Emin Pasha Gulf,
56 Tanzania - 14.6 km^2 ; and Winam Gulf, Kenya - 12.8 km^2 (Awange et al., 2019a).

57 To understand this decline, monitoring of its vegetational changes is vital and this predis-
58 poses it to remote sensing techniques due to its sheer size. This is because in-situ recordings
59 over such a large area is impractical and extremely susceptible to inaccuracy, whilst photogram-

60 metric processes would prove to be far too expensive when conducted yearly ([Awange et al.,](#)
61 [2019a](#)). Even with this realization, there is little information pertaining to the changes in areal
62 extent of LVB’s vegetation coverage and their triggers. Studies that have been conducted have
63 focussed largely on variance of NDVI values in the region (e.g., ([Omute et al., 2012](#))) but have
64 not equated the extent of which this contributes to changes in vegetation coverage. For in-
65 stance, a 2014 study conducted in the Winam Gulf (Kenyan section of LVB) utilised coarse
66 resolution 300 m MERIS (Medium Resolution Imaging Spectrometer) and fine resolution 30 m
67 Landsat-7 imagery to analyse the NDVI of aquatic vegetation in the area, and performed accu-
68 racy assessments for each Normalised Difference Vegetation Index (NDVI) extraction method
69 ([Cheruiyot et al., 2014](#)). With the gradual decrease in the surface area of Lake Victoria ([Awange](#)
70 [et al., 2019a](#)), there is reason to suggest that this would be a driving force for changes in the
71 characteristics of the surrounding vegetation within the catchment.

72 Factors that have been known to influence vegetation variability include both climatic and
73 anthropogenic. For example, an NDVI variability study in the USA from 1982 - 1992 confirmed
74 with a 99% confidence that El Niño Southern Oscillation (ENSO) was the primary driver for
75 influencing NDVI interannual variability ([Li & Kafatos, 2000](#)). A South African study ([Richard](#)
76 [& Poccard, 1998](#)) found that seasonal rainfall changes were primary NDVI drivers in areas where
77 the annual rainfall was between 300 and 900 *mm*, or if the total rainfall between the rainy and
78 dry seasons were profound. Within LVB, [Nicholson et al. \(1990\)](#) and [Omute et al. \(2012\)](#)
79 showed the influence of climate on vegetation by looking at the relationship between vegetation
80 and rainfall in the region. Furthermore, climate variability through global teleconnections such
81 as ENSO and Indian Ocean Dipole (IOD) on the one hand, and seasonal trends on the other
82 hand, are also known influencers of short-term variability of vegetation in the LVB region (see
83 e.g., [Park et al., 2020](#); [Zhao et al., 2020](#); [Detsch et al., 2016](#); [Williams & Hanan, 2011](#); [Plisnier](#)
84 [et al., 2000](#)(@; [Agutu et al., 2020](#); [Omute et al., 2012](#); [Awange et al., 2019a](#)).

85 This study extends the work of [Awange et al. \(2019a\)](#), which analysed the physical dynam-
86 ics of Lake Victoria by studying vegetation changes in its entire basin (i.e., LVB), which forms
87 its catchment. The difference between the two studies is that [Awange et al. \(2019a\)](#) aimed
88 at looking at changes on the physical surface of the lake and the triggers therein from 1984
89 to 2018. The current study employs high temporal-resolution MODIS (Moderate Resolution
90 Imaging Spectroradiometer), CHIRPS (Climate Hazards Group InfraRed Precipitation with

91 Station) precipitation, Google Earth Pro imagery and Gravity Recovery and Climate Experi-
92 ment (GRACE)-based Mascon’s total water storage products for the 2003-2018 period, and the
93 statistical PCA (principal component analysis), to determine (i) significant “hotspots”, i.e., the
94 main differences vegetation areas, which negatively impacted within LVB, and (ii), the extent
95 of which anthropogenic and climatic variability have contributed to the “hotspots” formation.
96 In undertaking this study, we seek to understand vegetational changes within LVB, which as
97 catchment, is associated with the Lake’s physical surface changes. Investigating the actual
98 causality however, is out of the scope of the current work.

99 The remainder of the study is organised as follows: In Section 2, study area, the data and
100 methods are presented. Section 3 covers results and discussion before concluding in Section 4.

101 2. Data and methods

102 2.1. Lake Victoria Basin: Background

103 Lake Victoria, the world’s largest tropical lake and second-largest freshwater lake is located
104 in Eastern Africa ($31^{\circ}39' - 34^{\circ}53' \text{ E}$ and $0^{\circ}20' - 3^{\circ} \text{ N}$), with its shoreline covering three countries,
105 Kenya (6%), Uganda (45%) and Tanzania (49%) (Awange and Ong’ang’a , 2005; Mati et al.,
106 2008). Its basin area (Figure 1) is almost three times the size of the lake and extends over
107 the three East African countries together with Rwanda and Burundi. The basin refers to the
108 network of rivers and streams that transport water and nutrients into the lake and is about
109 193000 km^2 (Awange and Ong’ang’a , 2005; Awange et al., 2008; Mati et al., 2008). It therefore
110 facilitates the livelihood and wellbeing of over 40 million people across 5 countries at a density of
111 over 300 people per km^2 , with this population expecting to triple by the year 2050 (Awange et
112 al., 2019a,b; Okotto-Okotto et al., 2018). Water levels for Lake Victoria are primarily dependent
113 on rainfall (approx. 80% of total recharge) (Awange and Ong’ang’a , 2005; Awange et al., 2019a;
114 Kizza et al., 2009). The remaining 20% is the result of discharge from within the catchment
115 area (Awange and Ong’ang’a , 2005; Awange et al., 2008).

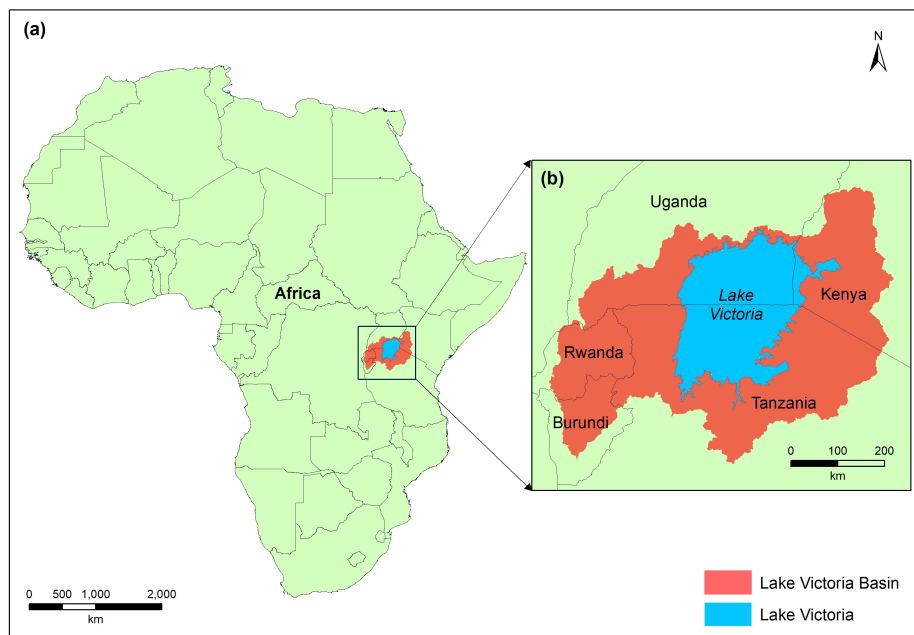


Figure 1: Lake Victoria Basin (LVB), the study area.

116 2.2. Data

117 Atmospherically corrected satellite images and hydroclimate data are employed to analyse
118 the degree of which climate variability and anthropogenic activities has impacted on vegetation
119 changes in the LVB. The data used for the investigation is elaborated upon in Sections 2.2.1–
120 2.2.4 and summarized in Table 1.

121 2.2.1. MODIS

122 MODIS NDVI data products (eMODIS NDVI V6) are available for direct download from the
123 United State Geological Survey (USGS) Earth Explorer database ([https://earthexplorer.
124 usgs.gov/](https://earthexplorer.usgs.gov/)). The eMODIS products includes either 7- or 10-day composite datasets with it's
125 native highest spatial resolution of $250 \times 250 m$ pixels size. The eMODIS NDVI V6 products
126 provide NDVI datasets over entire world. For this study, the 10-day intervals products are
127 download from the USGS website for the period 2003 - 2018 at three years interval, i.e., 2003,
128 2006, 2009, 2012, 2015, and 2018. The month of December is selected as it comes after the
129 short-rainy season of September - November, therefore potentially making it more suitable for
130 identifying NDVI changes (Awange et al., 2019a). A total of 18 images are downloaded for
131 December, i.e., 3 images for each year, with the temporal duration of the images being 10 or 11
132 days, i.e., 1st-10th, 11th-20th and 21st-31st for the evaluated month (December). The imagery
133 itself have coarse spatial resolution that encompasses the entire area of the African continent.
134 These products are regarded as high temporal resolution, therefore, might be a suitable products
135 for identifying areas of spatio-temporal changes in vegetation coverage (Chen et al., 2016). The
136 focus of this study is on the vegetation of the catchment area for the entire LVB, therefore,
137 the areal extent of the study is the lakes's surrounding areas (land area with vegetation cover)
138 resulting in output masking the lake's surface. A subset image created for 2018 MODIS NDVI
139 data using the shapefile of the LVB. All the rest of MODIS imagery are clipped using 2018
140 subset image. During the sub-setting of the images, a snap to pixel technique is used to make
141 sure every pixel for all year represent the same geographical location for capturing genuine
142 changes in NDVI values for the evaluated years.

143 2.2.2. Google Earth Pro

144 Google Earth Pro (GEP) is one of the online platform that uses base images with the option
145 that provides time series image analysis through image slider tool (Saleem & Awange, 2019).

146 This platform provides high to medium resolution satellite images, for instance IKONOS and
147 Landsat, respectively, over the world with variations from one region to another. For this study,
148 the freely available historical Landsat images from GEP are used to visually analyse further
149 (Section 3.4 and Figure 7) the identified NDVI hotspots where vegetation changes occur (see
150 Figure 6). The triggers for these changes could be factors that include anthropogenic activities,
151 e.g., urbanisation, forest and agriculture areas clearing. Furthermore, such visual analysis aid
152 in showing the extent of anthropogenic contributions to vegetation changes.

153 2.2.3. *GSFC Mascons*

154 Mass-concentration (mascon) grids are monitored by the Gravity Recovery and Climate
155 Experiment (GRACE) dual-satellite. The GSFC mascon estimation is processed by a general
156 approach that it models the best-fit trend and annual time-variable gravity signals (Sabaka et
157 al., 2008; Luthcke et al., 2013). It is worth mentioning that although the GSFC solution is
158 comprised of 41,168, 1 arc-degree mascon cells, the original spatial resolution within a region is
159 still 300 km. This means the time series of GSFC for each mascon within the same constraint
160 region (e.g., basins) is highly correlated to the near mascons. In this study, GSFC mascon is
161 used to monitor the equivalent water changes for every month, and compared to rainfall and
162 vegetation changes.

163 2.2.4. *CHIRPS Rainfall*

164 Climate Hazards Group InfraRed Precipitation with Station data (CHIRPS) is a rainfall
165 product that was developed to support the United States Agency for International Development
166 Famine Early Warning Systems Network (FEWS NET) (Funk et al., 2015a). It merges satellite
167 and in-situ observations to perform high spatial ($0.05^\circ \times 0.05^\circ$) resolution and monthly temporal
168 results, e.g., National Oceanic and Atmospheric Administration's (NOAA's) Rainfall Estimate
169 (REF2) (Love et al., 2004), African Rainfall Climatology (Novella and Thiaw, 2013), CHPClim
170 dataset (Funk et al., 2015b) and gauge products (Funk et al., 2015a). Generally, climate change
171 is inferred from climatology data that spans more than 30 years. Given that CHIRPS data are
172 available for 34 years, undertaking a PCA analysis on it can indicate the impacts of climate
173 change through the analysis of its time series. When this is compared with the epochs of
174 vegetation, one can infer on the impacts of climate change/variability on vegetation. The data
175 is available on <https://earlywarning.usgs.gov/fews/datadownloads/Global/CHIRPS>.

Table 1: A summary of data employed in this study.

Description	Sensor	Time-line	Purpose	Data Source
NDVI	MODIS	Dec. 2003-2018	Vegetation changes	https://earthexplorer.usgs.gov/
Google Earth Pro	Landsat	Dec. 2003-2016	Anthropogenic impacts	Application-based
GSFC Mascons	GRACE	January 2003 – July 2016	Water altimetry changes	https://ccar.colorado.edu/grace/gsfsc.html
Infrared rainfall data	CHIRPS	1984-2018	Impacts of climate change	http://chg.geog.ucsb.edu/data/chirps/

176 *2.3. Methods*

177 NDVI change maps are created for inter-annual and intra-annual variations for the 2003,
 178 2006, 2009, 2012, 2015, and 2018 MODIS imagery datasets to identify zones of the catch-
 179 ment area that have undergone extensive decline in vegetation coverage. Principal Component
 180 Analysis (PCA) is computed on rainfall data to determine if any reduction in vegetation is
 181 climate-driven with help of the change maps generated from NDVI data. Google Earth Pro
 182 (GEP) imagery from the time of vegetation decrease for each hotspot is obtained to visualise
 183 if the decline is anthropogenically-driven. There are large amount of variance in NDVI val-
 184 ues among the evaluated years, what can be partly attributed to the low spatial resolution
 185 ($250 \times 250 m$ pixels size) of MODIS imagery. However, high temporal resolution of the data
 186 permit a valid comparison over time for the NDVI value changes. An NDVI threshold of $>$
 187 0.2 is used to determine whether a pixel is considered vegetation or non-vegetation (e.g., Pu
 188 et al., 2008; Duarte et al., 2018), and the total area of vegetation and non-vegetation for each
 189 year then determined by generating binary dataset for each year using ArcGIS environment.
 190 This section elaborates on the image-processing procedures used, including; fishnet creation,
 191 binary outputs, standardised anomaly, image difference. A structure chart of the methodology
 192 is presented in Figure 2.

193 *2.3.1. Pre-processing*

194 All 18 of the MODIS images are clipped to the extent of the basin outside of the lake prior
 195 to further processing. NDVI values > 0.2 are extracted to create binary outputs depicting
 196 vegetation and non-vegetation pixels. The same threshold value was used in Pu et al. (2008);
 197 Duarte et al. (2018) as a means of representing shrubs and meadows. The NDVI statistics of
 198 the images are further normalised to provide the exact NDVI values of all vegetation pixels
 199 determined from the binary output, as well as categorising all non-vegetation pixels ≥ 0.2 as a
 200 ‘0’ value.

201 A fishnet of points is then created for each pixel in the study area (over 2.6 million; 250 m x
202 250 m pixels). Within the fishnet dataset, the NDVI values for each pixel is extracted from all
203 images and added to their respective points. The pixel mean NDVI value (i.e., the mean was
204 used for original products for 7 or 10 days interval) for each year (2003, 2006, 2009, 2012, 2015,
205 and 2018) is calculated from the three NDVI values of the original datasets of each year. The
206 overall mean NDVI value of those years is then obtained from these outputs, i.e., the pixel mean
207 for all years is then calculated using the pixel mean values for each year, which in turn allows
208 the overall mean NDVI values to be extracted. Using the pixel mean value of each year, new
209 rasters are derived depicting vegetation and non-vegetation areas using the NDVI threshold of
210 > 0.2 .

211 *2.3.2. Anomaly calculations*

212 Annual anomalies are computed with respect to the mean change of each pixel. These values
213 are utilised to demonstrate the extent of changes in NDVI of the basin. Two approaches are
214 used to present this; (i) the inter-annual, which highlight short-term 3-year annual trends for
215 NDVI differences, i.e., 2003-2006, 2006-2009 etc., and (ii), the intra-annual maps that present
216 short and long-term trends demonstrating the degree of NDVI difference for each of the years
217 in relation to 2003 (which is set as the base year).

218 *2.3.3. Hotspot significance maps*

219 The significance of the mean annual changes of NDVI P -value for each pixel is the deter-
220 mining factor in the identification of a “hotspot” signifying reduction in vegetation coverage.
221 Vegetation decreases are indicated for Z -values < 0 , with the significance of the trends of veg-
222 etation changes indicated by P -values < 0.05 . P and Z values are calculated for all pixels in
223 the study area where any pixel that fulfils both criteria is interpreted as decreasing at 95%
224 confidence.

225 *2.3.4. Principal Component Analysis (PCA)*

226 Principal component analysis (PCA) is a technique for reducing the dimensionality of datasets,
227 increasing interpretability but at the same time minimizing information loss (Awange et al.,
228 2020). It is useful for identifying variance in hydrometeorological parameters Dyer (1975);
229 Awange et al. (2011, 2014, 2016, 2019a). In this study, it is employed to analyse gridded

230 CHIRPS rainfall variation throughout the catchment area for the 1984-2018 period in order to
 231 infer the impacts of climate change on vegetation. As NDVI values are a measure of vegetation
 232 greenness, higher NDVI values are dependent on higher water content (Foody, 2003), eutrophication
 233 (especially around the lake perimeter, e.g., Coladello et al., (2020)), the vegetation
 234 type (e.g., (Omute et al., 2012)), density and height which translated in high NDVI values.
 235 Also, in such spatial resolution (250 m) mixed pixels significantly affect the NDVI values on
 236 the one hand while on the other hand, the meteorological conditions in the previous days of the
 237 selected MODIS images can affect the NDVI values. The patterns determined from the PCA
 238 analysis can determine whether the reduction in NDVI values from the ‘hotspots’ derived from
 239 the MODIS imagery processing are climate-driven through the analysis of rainfall variability
 240 (e.g., Awange et al., 2016, 2019a) and comparing with changes in NDVI (e.g., Omute et al.,
 241 2012).

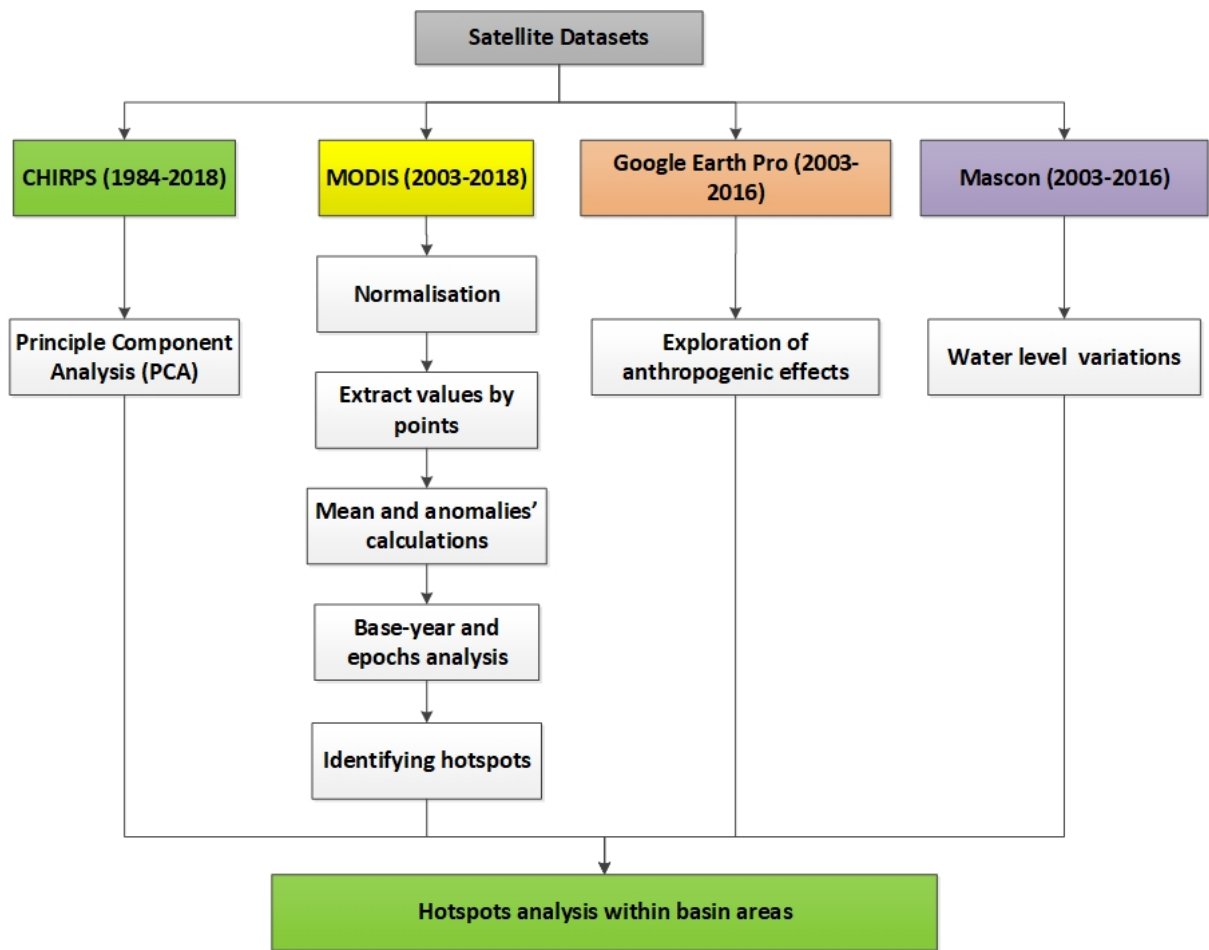


Figure 2: Workflow for the investigation.

242 **3. Results and discussion**

243 *3.1. Vegetation analysis within LVB*

244 The lower mean NDVI values for December 2006 and 2018 (Figure 3a) can be largely at-
 245 tributed to the lower than average rainfalls in Eastern Africa, which were caused by the oc-
 246 currence of La Niña events in 2006 and 2016-2017 respectively (Hoell et al., 2017; Setimela
 247 et al., 2018). Conversely, the higher mean NDVI value for December 2015 (Figure 3a) can be
 248 attributed to higher than average rainfalls from the occurrence of an El Niño event in 2015-2016
 249 (Setimela et al., 2018).

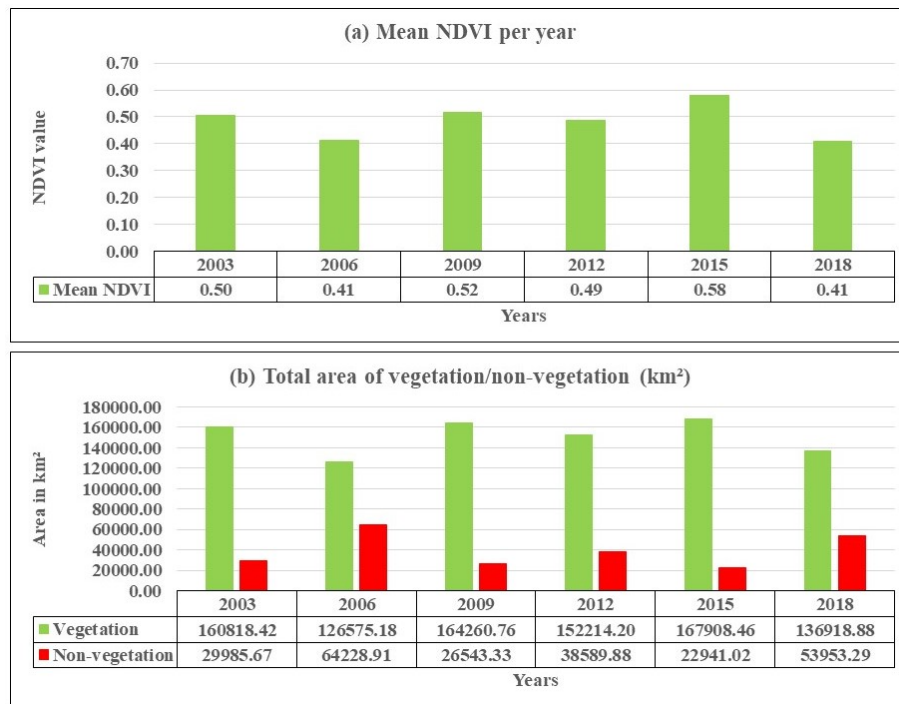


Figure 3: (a) Mean NDVI, and (b), Area of vegetation/non-vegetation using NDVI > 0.2 threshold.

250 *3.1.1. Epoch and base-year trend change maps*

251 As stated earlier (see Section 2.3.2), epoch and long-term trends of NDVI changes indicate
252 the potential whereabouts of significant hotspots in the basin. The primary purpose of these
253 maps is to identify areas in the LVB have undergone drastic transformations in overall vegetation
254 greenness within its respective recorded period.

255 As the epoch maps highlight the short-term 3-year annual trends, it therefore can highlight
256 the extent in the variation of NDVI values caused by extreme weather events. The 2003-2006
257 and 2015-2018 maps (Figure 4a) display the magnitude of environmental impact caused by
258 lower than average rainfall from their respective La Niña events that were highlighted in Hoell
259 et al. (2017); Setimela et al. (2018). Figure 4a demonstrates the capacity in which rainfall
260 variations can have on the environment in the short-term. However, vegetation that has yet
261 to be subjected to significant human interference has the capacity to replenish when annual
262 rainfall levels increase. This is displayed in the 2006-2009 (Figure 4a), which occurred after the
263 La Niña event of 2006 as well as the 2012-2015 period (Figure 4a, which occurred during the
264 El Niño event of 2015-2016. The decreases of NDVI values are more widespread and of greater
265 magnitude. In 2003-2006 the impact is generally more profound west of the lake. Significant
266 clusters of reduced NDVI values have been identified in the south-west of the catchment area
267 (Burundi), central to the western side of the lake (Tanzania), north-west of the catchment area
268 (Uganda) and Central Rwanda. East of LVB, there is a large general decrease of NDVI spread
269 across Kenya, which sprawls across into the Tanzanian border and continues nearby to the
270 southern border of the lake.

271 In the 2015-2018 period, the impact is even more extreme. There are additional areas that
272 have undergone significant NDVI decreases; namely along the northern boundary of the lake as
273 well as more extensively east of the lake across Kenya and Tanzania. However, it must be stated
274 that the extremity of the reduction is harnessed from the contrasting extreme weather events
275 that occurred in 2015 and 2018. As stated in Section 3.1, higher than average rainfall occurred
276 in the Lake Victoria region due to an El Niño event in 2015-2016, whereas lower than average
277 rainfall occurred due to a La Niña event in 2016-2017 (Setimela et al., 2018). These extreme
278 weather events would be the primary contributor as to why NDVI is higher than normal for
279 2015, and lower than normal for 2018. This means that when computing the difference in
280 the pixel anomalies inter-annually for 2015-2018, the result would indicate widespread NDVI

281 decrease.

282 The base-year map (Figure 4b) highlights the long-term trends for NDVI changes for all
283 years from 2003. In a similar vein to the outputs for the epoch maps (Figure 4a), the western
284 side of the basin is substantially more impacted. Within the 15-year time-span; south-western
285 Uganda, north-western Tanzania, central and eastern Rwanda as well as northern Burundi
286 have all been inflicted in this negative trend. East of the lake has also undergone some impact,
287 particularly in an area in south-western Kenya. However, the long-term trend indicates that
288 most of the primary hotspots occur west of the lake, see more details in Section 3.1.2.

289 3.1.2. Identification of significant hotspots

290 The next step is to determine the whereabouts of statistically significant hotspots that in-
291 dicate a long-term trend in vegetation decline. The criteria for identifying pixels that would
292 formulate the hotspots are $Z \leq -2$ and $P < 0.05$, indicating significant decrease at 95% confi-
293 dence. All pixels meeting this criteria are extracted from their respective datasets. All pixels
294 that represent both criteria could potentially be located within clusters of similar pixels. There
295 are such pixels located widespread in the study area, however, the purpose is to identify more
296 aggregations, i.e., areas, in which the separation of neighbouring pixels is no greater than 30
297 km. Due to the large size of the study area, it is understood that using a larger distance
298 threshold could provide a more sustainable output in terms of the number of clusters that will
299 result.

300 That resultant output when utilising the 30-km separation threshold are 8 significant hotspots
301 (Figure 4c); 5 of which were in Uganda, and 1 each for Kenya, Tanzania and Rwanda respec-
302 tively. With Uganda containing the most hotspots, those findings correlate with the long-term
303 trends analysed in Section 3.1.1, in which vegetation located north-west of Lake Victoria was
304 found to have undergone the largest areal decrease in NDVI from 2003-2018. Evidently, parts of
305 the study area that were subjected to short-term vegetation changes were not deemed significant
306 enough in the P and Z outputs in Figure (4c) to be deemed a hotspot.

307 The results show that there is some correlation between the NDVI hotspots deduced in
308 this investigation and the hotspots discovered in Awange et al. (2019a) that demonstrated
309 a significant decrease in the surface area of Lake Victoria itself. That study concluded that
310 Winam gulf (Kisii, Kenya), Emin Pasha gulf (Katoro, Tanzania), Mwanza gulf and Birinzi

311 (Kampala & Masaka, Uganda) were the hotspots. Each of these places has expanded their
 312 urban environments to various extents since 2003, whilst also showing signs of deforestation and
 313 clearing to enable the provision of agriculture and other rural industry. The assumption can be
 314 made that this outwards urban expansion has not only reduced all vegetation characteristics
 315 but has also necessitated the extraction of nearby freshwater that Lake Victoria provides for
 316 these areas.

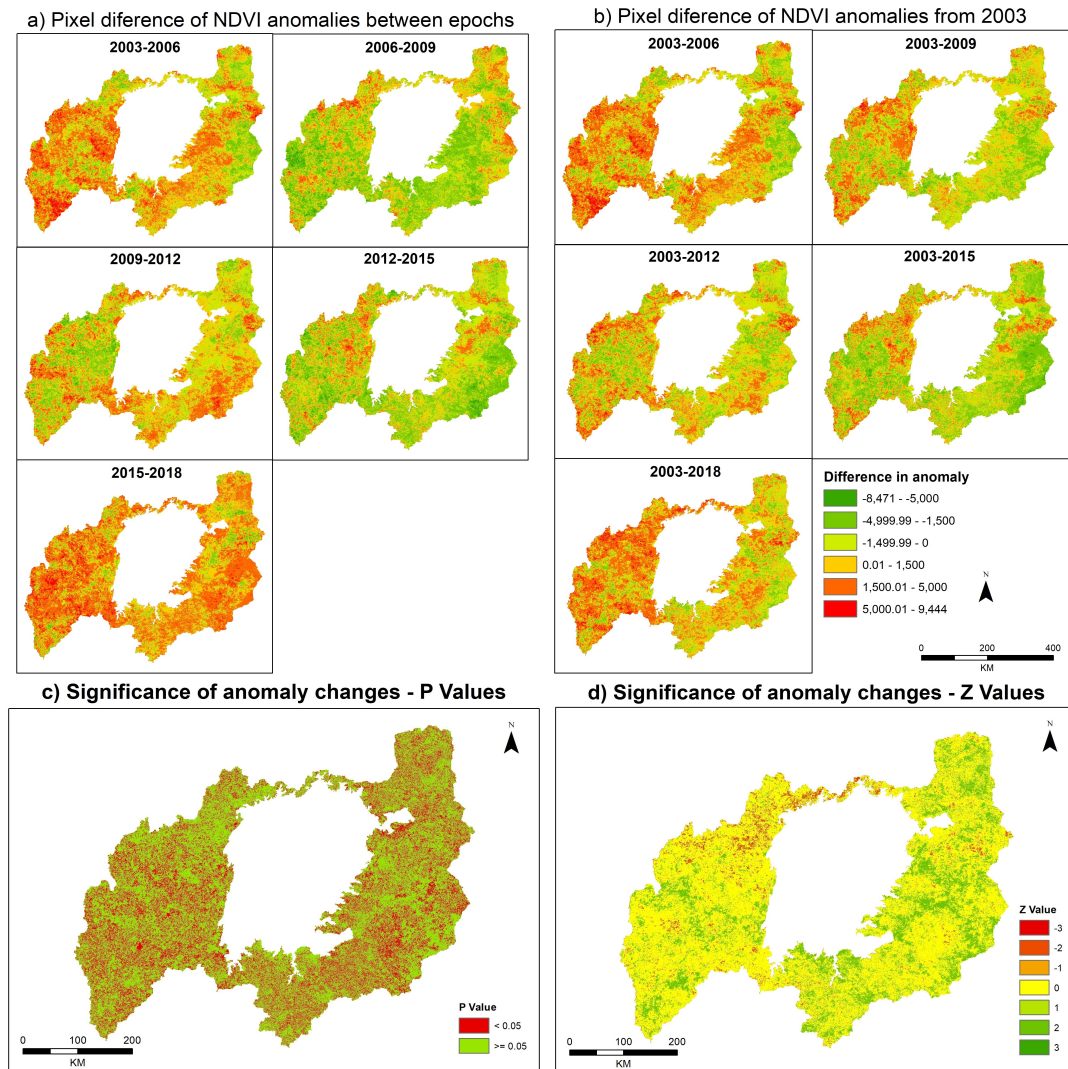


Figure 4: (a) Short-term three-year epoch differences in NDVI, (b) long-term NDVI trends from 2003, (c), significance of anomaly changes using maps depicting P values, and (d), significance of anomaly changes using maps depicting Z values.

317 *3.2. Rainfall variability within LVB*

318 For the 1984-2018 period, a PCA analysis is performed on the rainfall parameters within
319 LVB. Figure 5 shows the principal components (PC; time series) and the empirical orthogonal
320 functions (EOF; spatial maps) both of which have to be interpreted together to understand
321 rainfall variability within the LVB. Both PCs and EOFs account for the total variation in the
322 rainfall (Dyer , 1975). The results of the PCA, where PC1 (accounting for 75.2% of the overall
323 variance in rainfall) depicts the dominant seasonal rainfall superpositioned with the annual
324 signal, PC2 (19.1%) shows annual rainfall variation, while PC3 (5.6%) shows extreme rainfall
325 events (i.e., those associated with El Niño and La Niña)), are consistent, e.g., with the results
326 of Awange et al. (2013) who found four modes over the basin for the period 2003-2013. For
327 instance, Awange et al. (2013)'s PC1 (representing 63% of total variance of the rainfall) showed
328 a superposition of the annual and seasonal variabilities while PC2 (13%) related to the annual
329 variation and PC3 showed a summation of interannual changes and a linear trend over the
330 basin. The first three EOF modes in the present study are identical to those of Awange et al.
331 (2013). Other similar findings are presented, e.g., in (e.g., Khaki & Awange , 2019; Awange et
332 al., 2019a).

333 Identifying if there are climatic drivers in the formation of NDVI hotspots is accomplished
334 through analysing the variation trends derived from the PCA output. In general, overall rainfall
335 decreased throughout the Lake Victoria basin, in some places by as high as 250 millimetres,
336 such as along the western boundary of the lake and along the south-western boundary of
337 the study area within Burundi. Seasonality caused its most profound rainfall decrease along
338 the southern extremities of the study area. The decrease became gradually less profound
339 progressively north from the southern extent. The north-eastern corner overlapping Kenya and
340 Uganda recorded increased rainfall due to seasonal variations. Extreme weather events result in
341 horizontal contrast with increased rainfall recorded over most of the western half of the study
342 area and decreased rainfall over most of the eastern half.

343 For the NDVI hotspot identified in Kisii, Kenya the only rainfall variable that could have had
344 any effect in the long-term reduction in NDVI is the La Niña extreme weather events. However,
345 as that only accounts for 5.6% of rainfall variation, it is safe to assume that the hotspot is
346 the result of a non-climatic driver. For all the Ugandan hotspots, PCA does not provide any
347 meaningful indicator for climate contributing to its formation. Conversely, the total rainfall and

348 seasonality PCs indicate some possibility that they contribute to the formation of the Kigali
349 (Rwanda) and Katoro (Tanzania) hotspots. These PCs account for approximately 94.3% of
350 rainfall variance, which is reason enough to suggest that climatic impact could be meaningful.

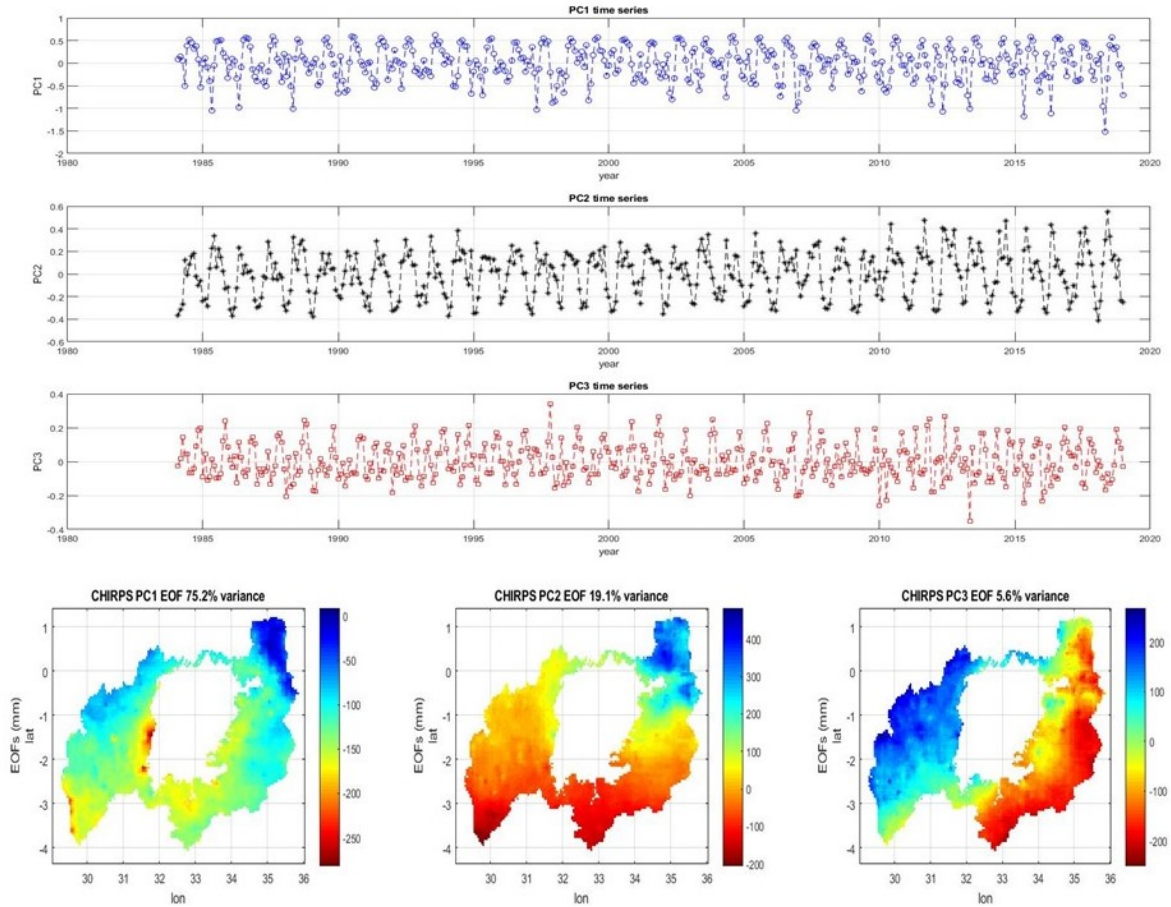


Figure 5: PCA analysis (a) Timeseries of PCA components, and (b), spatial pattern of components.

351 *3.3. Total Water Storage changes within the hotspots*

352 Annual total water storage (TWS; surface, groundwater, soil moisture and vegetation)
 353 changes have been obtained for all Gravity Recovery and Climate Experiment (GRACE) Mass
 354 Concentration (mascon) grids where hotspots lie (see Figure 6). The La Niña event in 2006
 355 presents a decrease in TWS by approximately 2 gigatons for all mascons. In contrast, the El
 356 Niño event in 2015-2016 presented an annual increase of approximately 3-4 gigatons for all
 357 mascons.

358 However, a short-coming in these assessments is their time-span. Analysis of NDVI changes
 359 occurred every three years from 2003-2018 while the mascon analysis occurs from January 2003
 360 – July 2016. The major detriment for that is that a La Niña event occurred in 2016-2017,
 361 resulting in lower rainfall for Eastern Africa. The effects that this had on TWS changes is not
 362 recorded. As December 2018 is the final period for NDVI analysis, the short-term and long-term
 363 effects of the La Niña event is displayed in Section 3.1.1. With the missing TWS change data,
 364 it makes it difficult to determine if rainfall variables are a major driver in the formation of the
 365 NDVI hotspots.

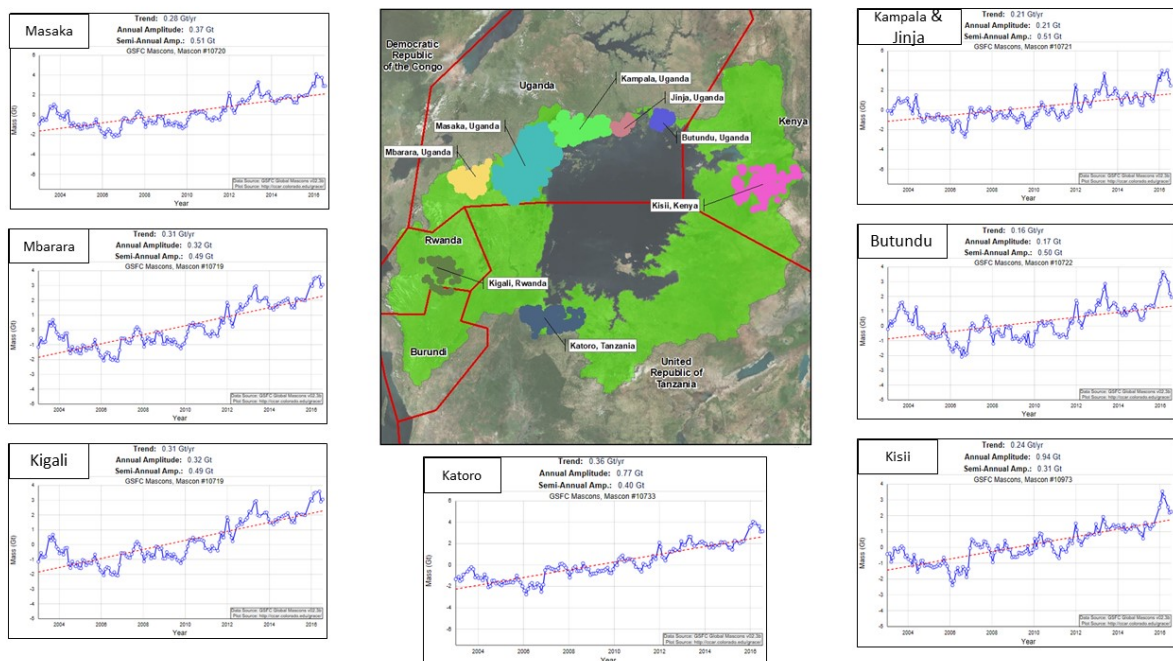


Figure 6: Gravity Recovery and Climate Experiment (GRACE) Mass Concentration mascon analysis of the total water storage changes within the NDVI hotspots.

366 3.4. Google Earth Pro imagery

367 The identified NDVI hotspots from MODIS data is assessed further using the satellite im-
 368 agery from GEP at different time-scales to provide insight into the extent of the anthropogenic
 369 impacts. From the analysis it is evident that for the Ugandan hotspots of Jinja, Kampala,
 370 Masaka and Mbarara, that outward urban expansion is the primary contributor for the long-
 371 term NDVI decline. The same conclusion can be drawn for the Rwandan and Kenyan hotspots
 372 also. The Katoro (Tanzania) hotspot can also be seen to have undergone urban expansion,
 373 but to a lesser extent. The Ugandan hotspot of Butundu is not necessarily derived from ur-
 374 banisation, but it was still caused by anthropogenic activities, as it appears that large-scale
 375 deforestation and clearing occurred in the region for widespread agricultural practices to begin.

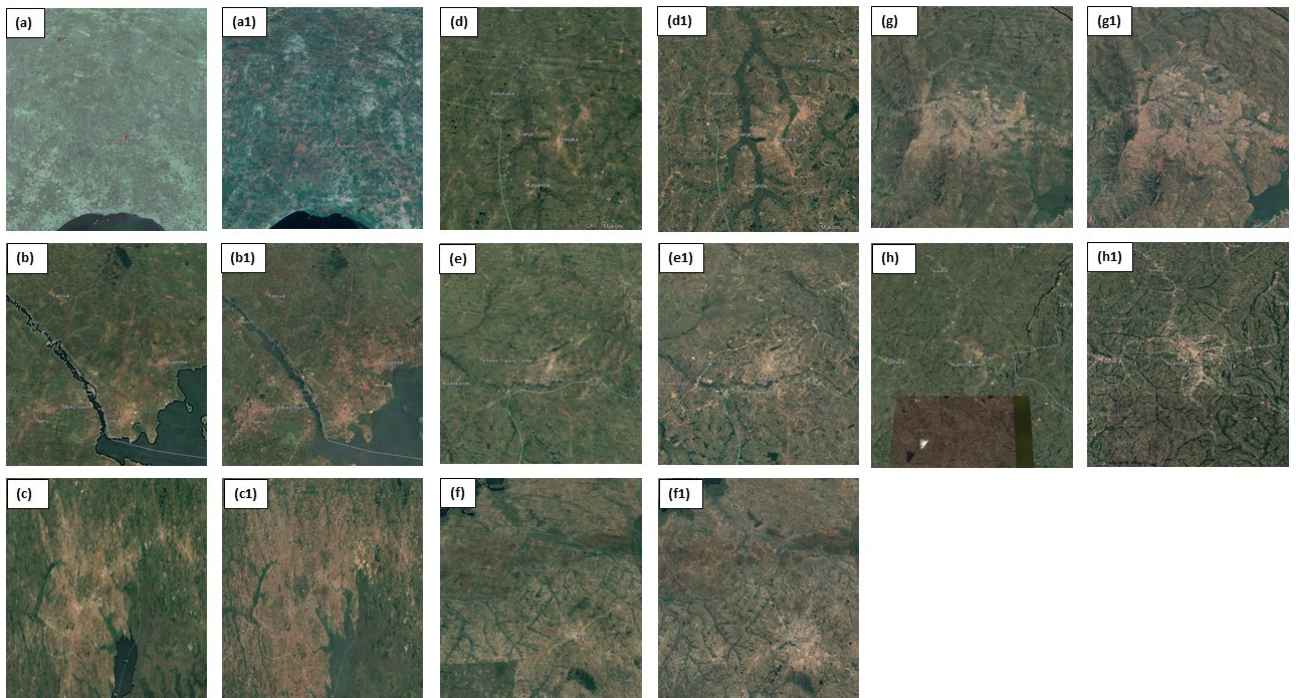


Figure 7: Google Earth Pro imagery with Landsat images as base map (a) Butundu, Uganda 2004 and (a1) Butundu, Uganda 2017; (b) Jinja, Uganda 2003 and (b1) Jinja, Uganda 2016; (c) Kampala, Uganda 2003 and (c1) Kampala, Uganda 2016; (d) Masaka, Uganda 2003 and (d1) Masaka, Uganda 2016; (e) Mbarara, Uganda 2003 and (e1) Mbarara, Uganda 2016; (f) Katoro, Tanzania 2003 and (f1) Katoro, Tanzania 2016; (g) Kigali, Rwanda 2003 and (g1) Kigali, Rwanda 2016; (h) Kisii, Kenya 2003 and (h1) Kisii, Kenya 2018.

376 **4. Conclusion**

377 Following a recent study that indicated reduction of Lake Victoria’s surface over the period
378 1984-2018, this contribution aimed at investigating changes in vegetation cover over the Lake
379 Vectorial Basin (LVB) over the period 2003-2018. To achieve this, the study employed MODIS
380 (Moderate Resolution Imaging Spectroradiometer), Google Earth Pro, CHIRPS (Climate Haz-
381 ards Group InfraRed Precipitation with station data) precipitation data, Google Earth Pro
382 imagery, Gravity Recovery and Climate Experiment (GRACE)-based Mascon’s water storage
383 products, and the statistical method of Principal Component Analysis (PCA). The assumption,
384 here, is that changes in vegetation within LVB is related to the lake’s physical dynamics and
385 as such, understanding vegetation changes within the basin and identifying the hotspots where
386 they occur could be essential to the overall management of the lake. The results show that
387 the vegetation within the LVB experienced temporal variations throughout the study period
388 (2003-2018). Specifically, the study found that:

- 389 1. Long-term vegetational changes within LVB over the period 2003-2018 were primarily
390 anthropologically driven, with urbanization expanding at the expense of vegetation as
391 seen from the Google Earth Pro imagery.
- 392 2. Eight “hotspots” (i.e., areas with significant vegetational changes) in total were identified
393 over LVB: 5 in Uganda, and one each in Kenya (Kisii), Kigali (Rwanda) and Tanzania.
394 Other than the Rwandan and Tanzanian hotspots where climate variability impacts were
395 visible, there is no meaningful evidence presented from the rainfall and Mascon’s TWS
396 analysis to suggest that anything other than human processes is causing long-term changes
397 in vegetation characteristics over the other hotspots.
- 398 3. Out of all the countries within the LVB, it can be said that Uganda has undergone the
399 most profound urbanisation processes since 2003, largely due to the expansion of its major
400 cities such as Kampala, Masaka and Jinja that were identified as hotspots. Small-scale
401 urban expansion also occurred in the Butundu, Mbarara and Katoro cities that do not
402 serve as major urban hubs, but instead service agricultural and industrial practices. The
403 expansion of these regional practices can be attributed to why they have been identified
404 as vegetation hotspots, as clearing of land is required to facilitate these practices.

405 Understanding the locations of vegetational changes is most profound, as well as the driving
406 forces associated with such changes, in that it provides critical information to major stakehold-

407 ers regarding future environmental management, policies and planning. Management of Lake
408 Victoria and its basin, therefore, would benefit from such analysis presented in this work.

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416 **References**

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