

24 baseline thresholds that are often applied arbitrarily, or in apparently subjective ways in the
25 literature. We use simulated SAR data to assess (1) the influence of residual noise and SBAS
26 network configuration on InSAR-derived deformation rates, and (2) how the number of
27 interferograms and data gaps in the time series may further impact the estimated rates. This
28 leads us to an approach for defining a SBAS network based on geodetic reliability theory
29 represented by the redundancy number (r -number). Simulated InSAR datasets are generated
30 with three subsidence signals of linear rates plus sinusoidal annual amplitudes of -2 mm/yr
31 plus 2 mm, -20 mm/yr plus 5 mm and -100 mm/yr plus 10 mm, contaminated by Gaussian
32 residual noise bounded within $[-2; +2]$ mm, $[-5; +5]$ mm and $[-10; +10]$ mm,
33 corresponding to standard deviations of approximately 0.5 mm, 1.5 mm and 3.0 mm,
34 respectively. The influence of data gaps is investigated through simulations with percentages
35 of missing data ranging from 5% to 50% that are selected (1) randomly across the 4-year
36 time series, and (2) for three-month windows to represent the northern winter season where
37 snow cover may cause decorrelation. These simulations show that small deformation rates are
38 most adversely affected by residual noise. In some extreme cases, the recovered trends can be
39 contrary to the signal (i.e., indicating uplift when there is simulated subsidence). We
40 demonstrate through simulations that the r -number can be used to pre-determine the
41 reliability of SBAS network design, indicating the r -values between ~ 0.8 and ~ 0.9 are
42 optimal. r -numbers less than ~ 0.3 can deliver erroneous rates in the presence of noise
43 commensurate with the magnitude of deformation. Finally, the influence of data gaps is not as
44 significant compared to other factors such as a change in the number of interferograms used,
45 although the blocks of “winter” gaps in the SBAS network show a larger effect on the rates
46 than gaps at random intervals across the simulated time series.

47

48 **1. Introduction and motivation**

49 Interferometric synthetic aperture radar (InSAR) has been demonstrated to be a powerful tool
50 for measuring the Earth's land-surface deformation owing to its high spatial and temporal
51 resolution, wide spatial coverage, and ability to acquire data remotely (e.g., [Hooper, 2008](#)).
52 However, InSAR measurements are contaminated by various error and noise sources, such as
53 those caused by digital elevation models (DEMs), atmospheric signal path delay, orbital
54 errors (ramps), temporal decorrelation, and other noise sources (e.g., [Lee et al., 2012](#); [Murray
55 et al., 2019](#)). Multi-temporal InSAR (MT-InSAR) methods were proposed to reduce these
56 error and noise sources (e.g., [Hooper, 2008](#)). These methods work by analyzing a network of
57 multiple acquisitions to derive the deformation time series and thus deformation rate (e.g.,
58 [Shanker et al., 2011](#)).

59 MT-InSAR methods can be classified into two principal categories, comprising the
60 persistent scatterer (PS) method (e.g., [Ferretti et al., 2001](#); [Hooper et al., 2007](#); [Hooper et al.,
61 2004](#)) and the small baseline subset (SBAS) method (e.g., [Berardino et al., 2002](#); [Cavalié et
62 al., 2007](#); [Hetland et al., 2012](#); [López-Quiroz et al., 2009](#); [Lundgren et al., 2001](#); [Schmidt &
63 Bürgmann, 2003](#); [Usai, 2003](#)). SBAS is among the most commonly used methods that makes
64 use of a network of interferograms from which temporal and perpendicular baselines are
65 limited in time and length to reduce the effects of geometric decorrelation (e.g., [Crosetto et
66 al., 2016](#); [Shanker et al., 2011](#); [Zebker & Villasenor, 1992](#)). This also incorporates an
67 approach to connect multiple SBASs that results in an increase in temporal and spatial
68 sampling ([Berardino et al., 2002](#)). The SBAS method has been used to measure land
69 deformation of various magnitudes, ranging from mm/yr (e.g., [Elliott et al., 2010](#); [Furuya et
70 al., 2007](#); [Jiang et al., 2011](#); [Schmidt & Bürgmann, 2003](#)) to cm/yr (e.g., [Amelung et al.,
71 1999](#); [Cavalié et al., 2013](#); [Chaussard et al., 2014](#); [Lee et al., 2012](#)) or even tens of cm/yr (e.g.,
72 [Chaussard et al., 2014](#); [López-Quiroz et al., 2009](#); [Motagh et al., 2007](#); [Short et al., 2011](#)).

73 InSAR data are degraded by various error and noise sources. The error caused by
74 DEM uncertainty can be reduced by a number of methods correcting for interferograms (e.g.,
75 [Berardino et al., 2002](#); [Bombrun et al., 2009](#)) or deformation time series (e.g., [Fattahi &](#)
76 [Amelung, 2013](#); [Pepe et al., 2011](#)). In order to reduce the effect of satellite orbital errors
77 (ramps), polynomial models based on network-sense ([Biggs et al., 2007](#); [Cavalié et al., 2008](#);
78 [Jolivet et al., 2012](#); [Lin et al., 2010](#)) or GPS data (e.g., [Neely et al., 2020](#); [Tong et al., 2013](#))
79 can be used. A number of methods can be applied to correct atmosphere phase errors utilizing
80 the stacking method (e.g., [Biggs et al., 2007](#); [Tymofyeyeva & Fialko, 2015](#)), using local data
81 assimilation, e.g., local atmospheric data (e.g., [Delacourt et al., 1998](#)) or zenith total delay
82 (ZTD) computed from GPS data (e.g., [Williams et al., 1998](#); [Yu, Li, & Penna, 2018](#); [Yu et al.,](#)
83 [2017](#)), utilizing global or regional atmospheric models (e.g., [Doin et al., 2009](#); [Jolivet et al.,](#)
84 [2011](#)), or integrating a global atmospheric model and GPS data to an atmospheric correction
85 model (e.g., [Yu, Li, Penna, et al., 2018](#)). Although these methods can be used to cope with
86 different errors and noise in InSAR measurements, they cannot be conducted perfectly, which
87 leads to remaining or residual errors and noise. Additionally, because of scheduling or other
88 technical issues, SAR images are not always regularly captured, or in other cases, blocks of
89 images acquired during extended periods (e.g., winter snowfall) may be omitted from
90 processing due to very low coherence, both of which may have a detrimental influence on the
91 estimated time series (e.g., [Kim et al., 2015](#); [Kohlhase et al., 2003](#)).

92 In InSAR SBAS data processing, pairs of scenes are chosen to form interferograms
93 from which an interferogram network is built in such a way to reduce decorrelation noise
94 through minimizing their time spans, and differences in look angle and squint angle ([Hooper](#)
95 [et al., 2012](#)). Coherent pixels to which a specific SBAS approach are applied can subsequently
96 be selected based on specific criteria, e.g., amplitude dispersion, spatial coherence, spectral
97 coherence or their combination ([Crosetto et al., 2016](#)). Different proposed SBAS approaches

98 are therefore based on thresholds that are, to a lesser or greater extent, different depending on
99 various factors, e.g., applications, data availability or the critical baseline, which in turn
100 depends on the wavelength of the radar sensor, spatial resolution and incidence angle ([Gatelli
101 et al., 1994](#); [Zebker & Villasenor, 1992](#)).

102 The temporal baseline threshold has been chosen varying from months to years (e.g.,
103 [Lanari et al., 2007](#); [López-Quiroz et al., 2009](#)), while the perpendicular baseline threshold has
104 been chosen ranging between hundreds of meters and over one thousand meters (e.g.,
105 [Berardino et al., 2002](#); [Chaussard et al., 2014](#)). The SBAS network thresholds are used with
106 the aim of maximizing the number of InSAR interferograms while minimizing their temporal
107 and spatial decorrelation, as well as reducing the computation time and data burden. Baseline
108 thresholds and pixel selection criteria used in several main SBAS approaches are listed in
109 Table 1. The question then arises as to whether there is some more objective means by which
110 to select these thresholds, which we consider herein. In this study, we deal with thresholds
111 used to select InSAR image pairs with an assumption that all pixels are of relatively high
112 coherence so as to be considered for SBAS processing.

113 We also consider the configuration of the SBAS network during our simulations. The
114 so-called network “optimization” problem has been applied to geodetic (surveying) networks,
115 which is traditionally divided among zero-, first-, second- and third-order problems (e.g.,
116 [Grafarend & Sansò, 1985](#)). The zero-order design (ZOD) is adopted for designing a reference
117 system, thus is also called “datum problem” ([Teunissen, 1985](#)). In the first-order design
118 (FOD), a network configuration is adopted by choosing the “optimal” locations of points in a
119 geodetic network that result in small changes in the positions of the preliminary chosen
120 network points ([Berné & Baselga, 2004](#); [Koch, 1985](#)). The objective of second-order design
121 (SOD) is to select “optimal” weights for the sometimes-different observations in which three
122 approaches can be utilized, including (i) direct approximation of the criterion matrix, (ii)

123 iterative approximation of the criterion matrix, and *(iii)* direct approximation of the inverse
124 criterion matrix (Schmitt, 1985a). By applying SOD, one seeks a network with high precision
125 (Amiri-Simkooei, 2004). In the third-order design, an existing network is improved, extended
126 or densified by introducing new points and/or additional measurements (Schmitt, 1985b).
127 This is also called the densification problem and can be understood to be a mixture of FOD
128 and SOD. A combined design, introduced by Vaníček and Krakiwsky (1986), refers to the
129 case where FOD and SOD problems are solved simultaneously.

130 In the experiments presented here, we use a time series of simulated InSAR data for
131 which we have control on the amount of error and residual noise introduced. We then
132 investigate the following parameters to determine what effect they have on InSAR-derived
133 rates of [simulated] land deformation. Our overarching aim is to find an “optimal” network of
134 interferograms that results in reduced data processing time. We assess 1) the influence of
135 residual errors and noise on SBAS-derived rates and the root mean square (RMS) of the
136 difference between simulated and SBAS-derived deformation time series for different
137 scenarios of the signal to noise ratio (SNR), 2) the effect of data gaps (i.e., missing scene
138 acquisitions) for both random and the three-month “winter” cases, and 3) the use of
139 redundancy numbers from geodetic network theory to design an “optimal” SBAS network.

140

141

142

143

144

145

146

147

148

Table 1. Summary of the main SBAS approaches

Reference	Interferogram selection thresholds	Pixel selection criterion
Berardino et al. (2002)	Perpendicular baseline (130 m)	Coherence
Mora et al. (2003)	Perpendicular baseline (24 m)	Coherence
Schmidt and Bürgmann (2003)	Perpendicular baseline (200 m)	Coherence
Lanari et al. (2004)	Perpendicular baseline (130 m)	Coherence
Hooper (2008)	Perpendicular baseline Temporal baseline Doppler baseline	Amplitude and phase stability
López-Quiroz et al. (2009)	Perpendicular baseline (500 m) Temporal baseline (9 months)	Coherence
Goel and Adam (2014)	Perpendicular baseline (150 m) Temporal baseline (150 days)	Statistical homogeneity test

149

150 2. InSAR SBAS algorithm used for this experiment

151 In summary, SBAS starts by forming an interferogram network using temporal and
152 perpendicular baseline thresholds, followed by selecting coherent pixels in which noise is
153 assumed to be negligible. Phase unwrapping is another step implemented in SBAS that can be
154 carried out either before or after pixel selection, depending on the implementation strategy
155 ([Gong et al., 2016](#)). The inversion step is subsequently implemented to convert small baseline
156 interferograms phase differences to a time series of displacements at the acquisition times.
157 With m interferograms generated from $(n + 1)$ InSAR images, the inversion equation can be
158 written as ([Berardino et al., 2002](#)):

$$\mathbf{A}\boldsymbol{\phi} = \boldsymbol{\delta\phi} \quad (1)$$

159 where \mathbf{A} is the design matrix of size $m \times n$, $\boldsymbol{\phi}$ is the vector of n (unknown) time series phase
 160 displacements of InSAR images at a pixel, $\boldsymbol{\delta\phi}$ is the vector of m (known) phase differences
 161 between each small baseline interferogram. In the SBAS approach applied in these
 162 simulations, the interferogram phase measurements can be expressed as (Agram et al., 2012;
 163 Gong et al., 2016):

$$\delta\phi_{ij} = \phi_j - \phi_i = \sum_{n=i}^{j-1} \delta\phi_n \quad (2)$$

164 where $\delta\phi_{ij}$ is the interferogram phase connecting i^{th} and j^{th} images, ϕ_i and ϕ_j are the phase
 165 values at i^{th} and j^{th} acquisitions, respectively, $\delta\phi_n$ is the pixel phase increment between n^{th}
 166 and $(n + 1)^{th}$ images. Equation (2) is utilized with an assumption of linear deformation
 167 between acquisitions that are adjacent in time (Berardino et al., 2002).

168 In SBAS data processing, a network is formed by choosing interferometric pairs with
 169 short temporal and perpendicular baselines limited by user-prescribed thresholds, and this
 170 controls the structure of the design matrix \mathbf{A} in Equation (1). With the above assumption of
 171 $(n + 1)$ InSAR images, the possible number of interferometric pairs (m) satisfies (Berardino
 172 et al., 2002):

$$\frac{n + 1}{2} \leq m \leq \frac{n(n + 1)}{2} \quad (3)$$

173 For each pixel selected, Equation (1) is applied to convert the phase difference from
 174 interferograms in the chosen network to the phase time series of displacements according to
 175 InSAR acquired times by applying least-squares (LS) (Schmidt & Bürgmann, 2003), singular
 176 value decomposition (SVD) (Berardino et al., 2002), or minimization of the L1-norm

177 (Lauknes et al., 2011). In most SBAS approaches, the design matrix \mathbf{A} is fixed to be used in
 178 the inversion step for all selected pixels. This is an advantage in terms of convenience and
 179 reduced processing time, but may suffer from decorrelation, particularly in vegetated or snow-
 180 covered areas where many pixels may decorrelate, so that there are large gaps in the spatial
 181 distribution of its products, e.g., a velocity map (Sowter et al., 2013). Methods using a flexible
 182 design matrix \mathbf{A} , e.g., the intermittent SBAS method (Sowter et al., 2013), have been
 183 proposed as a solution. In this simulation, however, we use a fixed-size \mathbf{A} matrix.

184

185 3. Network design used in geodesy

186 Geodetic surveying network “optimization” aims at finding a geometric configuration and a
 187 set of observations of sufficient precision to satisfy the desired positional quality criteria with
 188 lower financial and logistical costs (e.g., Kuang, 1993). The quality of a geodetic network is
 189 defined by the criteria of precision, reliability and economy (i.e., cost) of the network
 190 (Schmitt, 1985a). In geodetic network design, one seeks to minimize the objective function of
 191 economy and/or maximize that of precision or reliability of the network (e.g., Amiri-
 192 Simkooei, 2004).

193 The observational precision and network geometry are two crucial factors that
 194 influence the precision of a geodetic network. The variance-covariance (VCV) matrix is
 195 normally adopted to represent the network’s precision. With the assumption of a minimum
 196 constraint, the VCV matrix is expressed as (e.g., Kuang, 1996).

$$\mathbf{C}_x = \sigma_0^2 \left[(\mathbf{A}^T \mathbf{P} \mathbf{A} + \mathbf{D} \mathbf{D}^T)^{-1} - \mathbf{H} (\mathbf{H}^T \mathbf{D} \mathbf{D}^T \mathbf{H})^{-1} \mathbf{H}^T \right] \quad (4)$$

197 where σ_0^2 is the a priori variance factor, \mathbf{A} and \mathbf{P} are the design and weight matrices of
 198 observations, \mathbf{D} and \mathbf{H} are the minimum and inner constraint datum information matrices,
 199 respectively.

200 The reliability of geodetic networks, as defined classically by Baarda (1968), is the
 201 ability of a network to detect and resist against gross errors in observations. It is further
 202 divided into internal and external reliability as follows.

203 1) The internal reliability is defined as the ability of a network to detect gross errors, referring
 204 to the lower bounds of detectable gross errors (aka. the minimum detectable bias, MDB) that
 205 is expressed as (e.g., Baarda, 1968):

$$\nabla_0 l_i = \frac{\delta_0 \sigma_{l_i}}{\sqrt{r_i}} \quad (5)$$

206 where δ_0 is the lower bound for the non-centrality parameter, σ_{l_i} and r_i are the standard
 207 deviation and the redundancy or r -number of the i^{th} observation, respectively. The r -numbers
 208 of the observations are the diagonal elements of the matrix \mathbf{R} that are expressed as (e.g.,
 209 Amiri-Simkooei et al., 2012):

$$\mathbf{R} = \mathbf{I} - \mathbf{A}(\mathbf{A}^T \mathbf{P} \mathbf{A})^{-1} \mathbf{A}^T \mathbf{P} \quad (6)$$

210 where \mathbf{I} is the identity matrix.

211 2) The external reliability refers to the maximum effect of an undetectable gross error ($\nabla_0 l_i$)
 212 on the estimates of unknown parameters as:

$$\nabla_{0,i} \hat{\mathbf{x}} = (\mathbf{A}^T \mathbf{P} \mathbf{A})^{-1} \mathbf{A}^T \mathbf{P} \nabla_{0,i} l \quad (7)$$

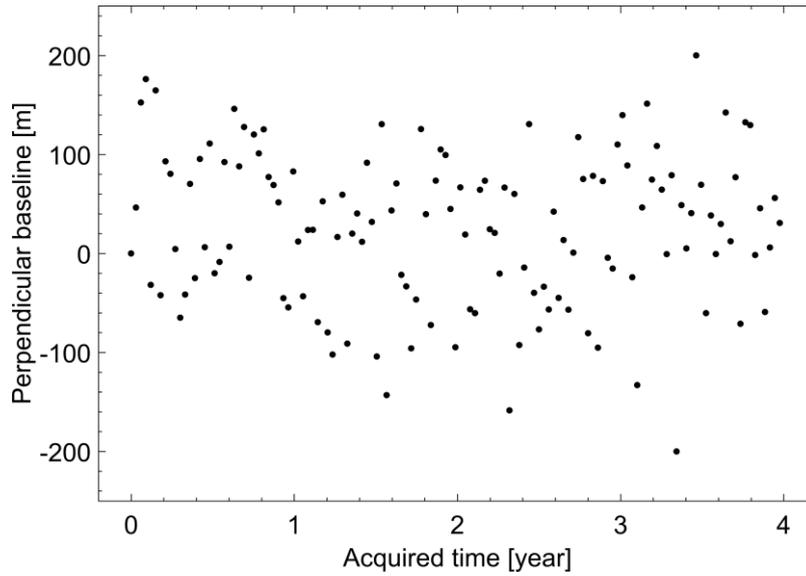
213 The internal reliability criterion is generally used as the measure for an “optimal”
 214 design of geodetic networks aiming at high reliability (Amiri-Simkooei, 2001), as shown in
 215 Equation (6). In this paper, we examine the redundancy number (r -number) as a diagnostic
 216 metric to determine the likely effectiveness of the SBAS network design and verify this with

217 simulation experiments. Specifically, for a given SBAS network with a corresponding design
218 matrix \mathbf{A} as shown in Equation (1), the r -number is computed using Equation (6) with the
219 weights \mathbf{P} of interferograms computed as the inverse of normalized (perpendicular and
220 temporal) baseline lengths, which will be described in Section 4.

221

222 **4. Generation of simulated data**

223 A time series of independent pixels that are reasonably representative of the range of Earth
224 deformations detected by InSAR are simulated, these being: mm/yr (e.g., [Elliott et al., 2010](#);
225 [Furuya et al., 2007](#); [Jiang et al., 2011](#); [Schmidt & Bürgmann, 2003](#)), cm/yr (e.g., [Amelung et](#)
226 [al., 1999](#); [Cavalié et al., 2013](#); [Chaussard et al., 2014](#); [Lee et al., 2012](#)), and tens of cm/yr
227 (e.g., [Chaussard et al., 2014](#); [López-Quiroz et al., 2009](#); [Motagh et al., 2007](#); [Short et al.,](#)
228 [2011](#)). Our simulated data cover a four-year time span with 11-day sampling interval that
229 corresponds to 133 equally time-spaced InSAR images. The baseline history of these 133
230 images, which is defined as the perpendicular baselines between images and the reference one
231 (i.e., the first scene), is assumed to be within $[-200, +200]$ m, which is approximately the
232 order of modern SAR missions such as C-band Sentinel-1 ([Yague-Martinez et al., 2016](#)) or
233 TerraSAR-X (TSX) (e.g., [Chen et al., 2016](#); [Lubitz et al., 2013](#)). The simulated baseline
234 history of 133 images is generated randomly with ranges between -200 m and $+200$ m with
235 that of the first scene being fixed to be zero (so leaving 132), and are shown as a scatter plot
236 in Figure 1.



237 Figure 1. Scatter plot of simulated perpendicular baseline history. Each black dot represents a
 238 SAR scene acquisition.

239

240 We take an interest in a land subsidence signal with both a linear trend and a
 241 superposed annual sinusoidal oscillation; all pixels are simulated to experience surface
 242 deformation in the SAR line of sight (LoS) with linear plus annual periodic terms, which are
 243 expressed as:

$$d_{i,j} = v_i t_j + a_i \sin(2\pi t_j) \quad (8)$$

244 where $d_{i,j}$ is the deformation of the i^{th} pixel at the j^{th} image with corresponding acquired
 245 time t_j , v_i and a_i are the linear rate (velocity) and annual amplitude of the same pixel,
 246 respectively. We select this model form because time series analysis of other environmental
 247 phenomena do likewise (e.g., [Davis et al., 2012](#); [Didova et al., 2016](#)).

248 The linear rates are chosen as -2 mm/year, -20 mm/year and -100 mm/year over
 249 the four-year period, which are representative of Earth deformation rates measured by InSAR
 250 (e.g., [Cavalié et al., 2013](#); [Chaussard et al., 2014](#); [Elliott et al., 2010](#)). The sinusoidal annual
 251 amplitude of Earth surface deformation has been drawn from the literature, which can range

252 from the order of millimeters to centimeters (e.g., Baldi et al., 2009; Bock et al., 2012; Davis
 253 et al., 2012; Dzurisin et al., 2009; Murray & Lohman, 2018; Osmanoglu et al., 2011). For
 254 example, Osmanoglu et al. (2011) report annual amplitudes of GNSS stations ranging from
 255 several millimeters up to ~2.6 centimeters. Murray and Lohman (2018) found seasonal
 256 amplitudes up to ~5 centimeters in California detected by InSAR and peak-to-peak
 257 amplitudes of ~6 centimeters from GNSS in the Amazon Basin (cf.
 258 <http://geodesy.unr.edu/NGLStationPages/stations/NAUS.sta>).

259 While there can be large annual signals in various parts of the world, we simulate
 260 more conservative cases of simulated deformation signal with pairs of signal parameters of
 261 linear rate plus annual amplitudes that are shown in Table 2. We then apply a Monte Carlo
 262 simulation (e.g., Kroese et al., 2014) with 1,000 pixels for each scenario. The number of
 263 tested pixels is chosen to avoid prohibitive computation times for the simulation experiments.
 264 The deformation time series of the 1,000 pixels are then computed for the 133 equally spaced
 265 11-day acquisition times using Equation (8), and are termed herein the “simulated
 266 deformation time series”. These are considered to be the “true” or noise-free signal, and will
 267 be used to validate the SBAS InSAR data processing results later in this paper.

268

269 Table 2. The three cases of simulated signals showing linear rates and annual amplitudes used

270

for experiments in Sections 5-7

Signal case	Linear rate	Annual amplitude
	[mm/yr]	[mm]
1	-2	2
2	-20	5
3	-100	10

271 With 133 InSAR images, the maximum possible number of interferograms is 8,778
 272 (Equation (3)). These 8,778 noise-free interferograms are then computed based on this
 273 simulated deformation SAR time series: the phase difference of an interferogram connecting
 274 i^{th} and j^{th} images is computed by subtracting the simulated time series value at i^{th} time from
 275 that at j^{th} time.

276 The simulated residual errors and noise, herein called the “residual interferogram
 277 noise”, are then added to the noise-free interferograms. Three sets of assumed 8,778 Gaussian
 278 noise values with zero mean are generated for each of the 1,000 pixels and bounded within
 279 $[-2; +2]$ mm, $[-5; +5]$ and $[-10; +10]$ mm, which correspond to standard deviations of
 280 approximately ± 0.5 mm, ± 1.5 mm and ± 3.0 mm, respectively (Table 3). Specifically, for
 281 each pixel, we first generate 8,778 random samples of a Gaussian distribution with a zero
 282 mean and a standard deviation of one. These are subsequently rescaled so that their ranges lie
 283 exactly within the bounds set in Table 3. We acknowledge that the residual errors and noise in
 284 real SAR data may not be Gaussian with zero mean because they originate from a variety of
 285 sources (e.g., DEM error, orbital ramp, atmospheric delay, etc). However, we would only ever
 286 be able to postulate the actual statistical distribution of real InSAR data errors, so instead
 287 make the simple assumption of Gaussian zero mean for our simulations.

288 Table 3. Simulated noise with various ranges and standard deviations

Noise case	Range	Standard deviation
	[mm]	[mm]
A	$[-2; +2]$	± 0.5
B	$[-5; +5]$	± 1.5
C	$[-10; +10]$	± 3.0

289

290 The simulated residual interferogram noise is generated in such a way that longer
 291 baseline lengths are assigned with noise of higher magnitude. Additionally, they have
 292 different ranges with the temporal baselines being from ~ 0.03 year to ~ 3.97 years, whilst the
 293 perpendicular baselines being between -376 meters and 400 meters. Therefore, they are first
 294 “normalized” by dividing all elements by the maximum value:

$$\begin{aligned} norm_btemp_i &= \frac{btemp_i}{\max(btemp)} \\ norm_bperp_i &= \frac{\text{abs}(bperp_i)}{\max[\text{abs}(bperp)]} \end{aligned} \quad (9)$$

295 where $norm_btemp_i$ and $norm_bperp_i$ are the “normalized” temporal and perpendicular
 296 baselines of the i^{th} interferogram, respectively which correspond to their values before
 297 “normalization” $btemp$ and $bperp$, $\text{abs}(\cdot)$ and $\max(\cdot)$ indicate the absolute and maximum
 298 values, respectively.

299 By this “normalization”, the normalized temporal and perpendicular baselines will
 300 have ranges between ~ 0 and 1 . The normalized baseline lengths of all interferograms are then
 301 computed with the i^{th} interferogram being:

$$norm_bsln_i = \sqrt{norm_btemp_i^2 + norm_bperp_i^2} \quad (10)$$

302 The normalized baseline lengths computed from Equation (10) are then used to assign
 303 the residual interferogram noise. Specifically, for each pixel with corresponding noise set of
 304 $8,778$ samples, the noise is assigned to interferograms by a way that an interferogram with a
 305 longer normalized baseline length will be assigned with noise of larger magnitude. We
 306 acknowledge that the influences of temporal and perpendicular baselines on interferometric
 307 noise are different. While the influence of perpendicular baselines can be quantified via their
 308 relationship with DEM error (e.g., [Lee et al., 2012](#)), the influence of temporal baselines is
 309 more sophisticated, which is dependent on the change of atmosphere and target environment

310 over time (Zebker et al., 1997; Zebker & Villasenor, 1992). Here, for the sake of simplicity,
311 we assume the two types of baseline are equal in terms of their weights in calculating
312 normalized baselines using Equation (10).

313

314 **5. Disruptive influences of residual noise and network configuration**

315 In order to assess the influence of residual noise and small baseline network configuration on
316 SBAS-derived land deformation rates, various interferogram networks were formed through
317 the use of different thresholds for the temporal baselines. Here, for the sake of simplicity
318 initially, we restrict the perpendicular baseline length to 200 m and only vary the temporal
319 baseline. Table 4 shows the temporal baseline thresholds that are applied with the resulting
320 number of interferograms.

321 We apply the SBAS approach to subsets of our simulated noisy interferograms (Table
322 4) using the GIANt software package (Agram et al., 2013; Agram et al., 2012). GIANt
323 incorporates most of the SBAS-based data processing approaches mentioned in the
324 Introduction, including the “traditional” SBAS (e.g., Berardino et al., 2002; Cavalié et al.,
325 2007; Schmidt & Bürgmann, 2003; Usai, 2003), the new SBAS (NSBAS) (Doin et al., 2011;
326 López-Quiroz et al., 2009), and the Multiscale InSAR Time-Series (MInTS) (Hetland et al.,
327 2012); cf. Table 1. Time series of deformation relative to the first-acquired SAR image time
328 for each of the 1,000 test pixels are generated assuming that there is no deformation in the
329 first acquisition. Both unweighted linear regression and unweighted LS are then applied to
330 those SBAS time series in order to compute SBAS-derived linear rates and annual sinusoids,
331 which are then compared with our simulated parameters listed in Table 2. The RMS of the
332 difference between simulated deformation time series (the “true” signal) and SBAS-derived
333 deformation time series is also computed in order to test dependence on the number of
334 interferograms chosen.

335 Table 4. List of networks tested in this study based on various temporal baseline thresholds.

336 The perpendicular baseline threshold is set fixed at 200 m (Figure 1).

Temporal baseline threshold [days]	Number of interferograms
22	263
33	376
44	498
55	621
66	745
77	863
88	986

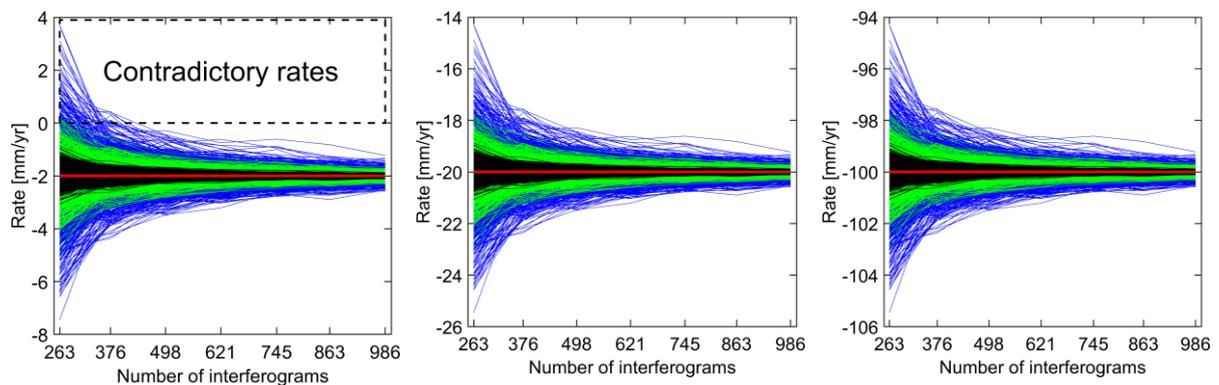
337

338 *5.1. Influences on simulated linear signals*

339 We first examine a signal where Equation (8) is adopted solely with the linear rate
 340 components of -2 mm/yr, -20 mm/yr and -100 mm/yr (Table 2). Figure 2 shows results
 341 from different combinations of simulated deformation rates and residual interferogram error
 342 and noise. Here, the assumed simulated linear rates are considered as the “true” rates to which
 343 the SBAS-derived rates are compared and the differences between them are herein termed the
 344 “errors in rate determination”. The SBAS rates are derived by fitting a linear regression to the
 345 corresponding deformation time series, then the errors in rate determination are calculated.
 346 The errors are shown in Figure 2, and are the same in both magnitude and sign among all
 347 three simulated linear rate cases from Table 2. Generally, the larger simulated residual
 348 interferogram noise (i.e., $[-10; +10]$ mm vs. $[-5; +5]$ mm vs. $[-2; +2]$ mm) leads to larger
 349 errors in the rate determination (cf. blue, green and black plotlines in Figure 2), whereas an

350 increase in the number of chosen interferograms (by choosing a larger temporal baseline
 351 threshold) can reduce this error.

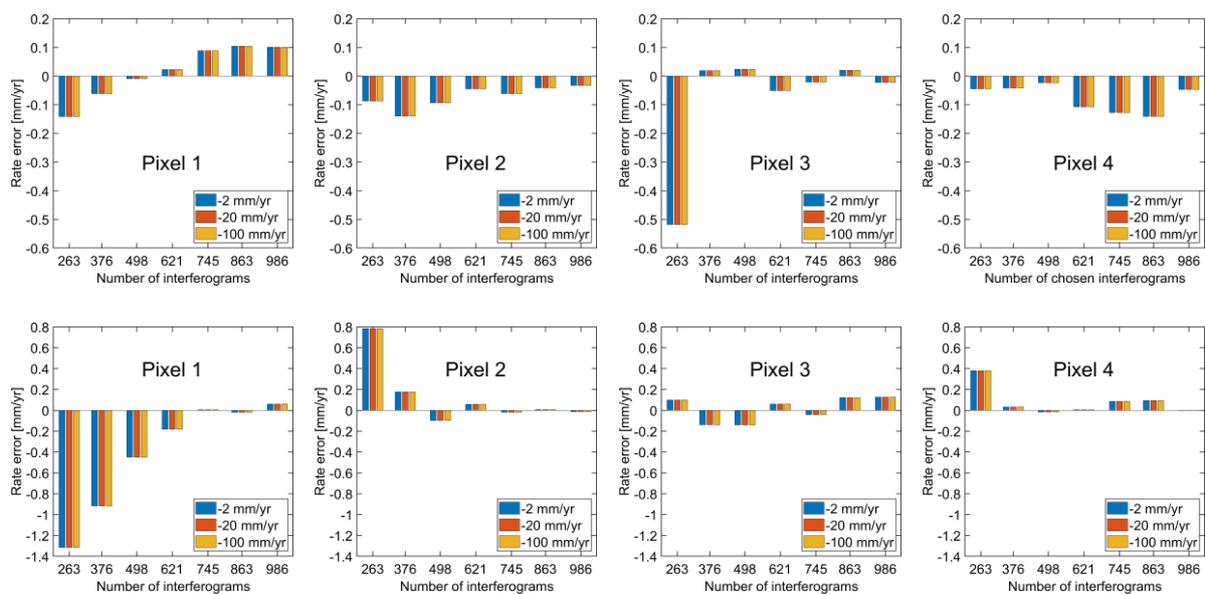
352 Additionally, while their trends are in an agreement for the cases of larger signal rates
 353 (i.e., -20 mm/yr and -100 mm/yr, Figure 2, middle and right), contradictory trends exist in
 354 the cases of small deformation (i.e., -2 mm/year, Figure 2, left), particularly when networks
 355 of fewer interferograms are used together with higher residual noise of $[-5; +5]$ mm and
 356 $[-10; +10]$ mm. Importantly, the SBAS-derived deformation trends are affected by not only
 357 the magnitude of noise, but also its relation to the signal size (see Figure 2, left), thus low
 358 SNR is more likely to result in incorrect or even contradictory trend estimates. In essence,
 359 small deformation rates in the presence of proportionally large noise may lead to spurious
 360 results, which become exacerbated in the presence of significant data gaps.

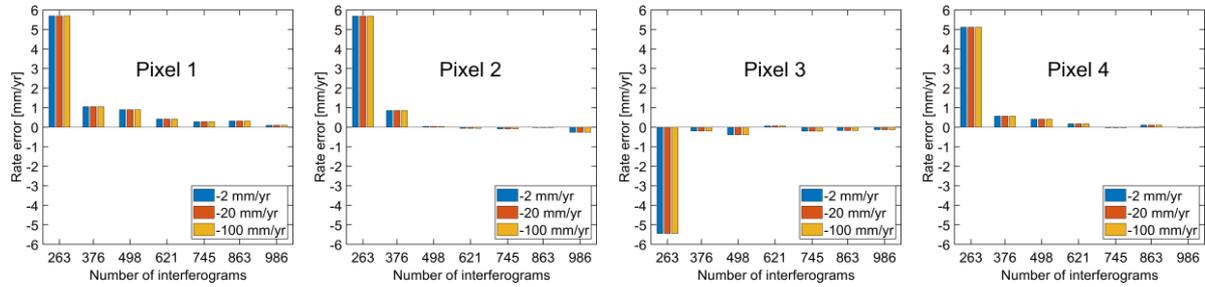


361 Figure 2. Comparison of rates computed by unweighted linear fit from combinations of
 362 different deformation signals. From left to right are simulated linear rate cases 1 to 3 (Table 2)
 363 contaminated by simulated residual interferogram noise. Black, green and blue polylines are
 364 SBAS derived rates computed from simulated data with simulated noise cases A to C,
 365 respectively (Table 3). Red horizontal lines represent the simulated rates. The black dashed
 366 box in the left panel is used to contrast between positive and negative rates that indicates
 367 contradictory trends.

368

369 The “errors in rate determination” are next compared for the networks listed in Table 4
 370 and shown in Figure 3 for four example pixels. Within a specific network and pixel, the
 371 retrieved rate errors are identical when the same residual noise is applied regardless of the
 372 signal rates. In other words, if a specific network chosen from Table 4 with corresponding
 373 interferogram noise set is applied, then its error in rate determination will not depend on the
 374 magnitude of simulated rate (cf. blue, orange and yellow bars in Figure 3). This is attributable
 375 to SBAS using the LS principle (Schmidt & Bürgmann, 2003) or the SVD method (Berardino
 376 et al., 2002). The results computed from applying the LS principle depend on redundant
 377 interferograms, together with residual interferogram error and noise that in turn depends on
 378 the configuration of the network (Berardino et al., 2002). The SBAS network configuration is
 379 specified by the design matrix \mathbf{A} as per Equation (1). Both the LS principle and SVD method
 380 result in the same InSAR-derived rates, except that the latter can cope with disconnected
 381 subsets of interferogram networks, whereas the former cannot (Berardino et al., 2002; Gong et
 382 al., 2016). Consequently, the same error in rate determination will result if the same residual
 383 noise is applied to a network regardless of the deformation rate.
 384





385 Figure 3. Comparison of rate errors computed from different networks for four example
 386 pixels. The top, center and bottom rows correspond to simulated noise cases A, B and C
 387 (Table 3). Note the different scale on the y-axis for each noise case.

388

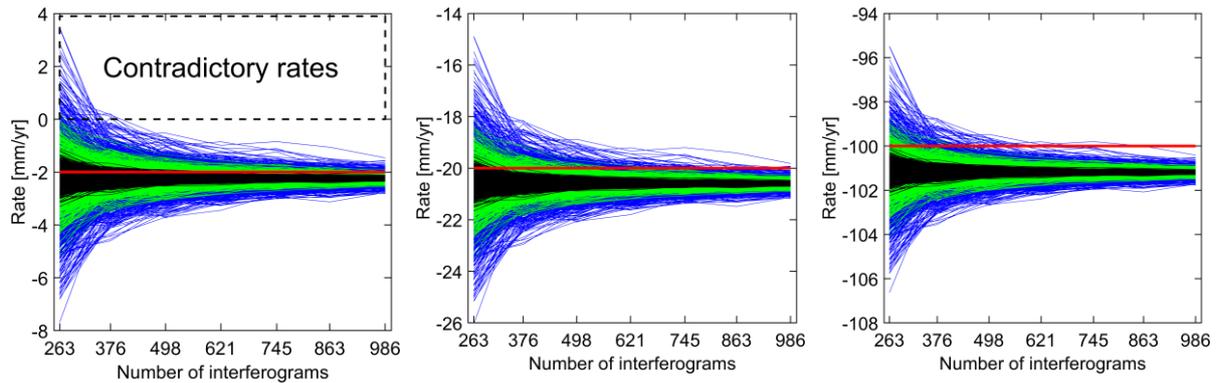
389 5.2. Influences on non-linear signals

390 We next examine the signal combining both a linear trend and sinusoidal annual terms. As
 391 mentioned in Section 4, we apply pairs of signal parameters of linear rate plus annual
 392 amplitude, which are -2 mm/yr plus 2 mm, -20 mm/yr plus 5 mm, and -100 mm/yr plus
 393 10 mm (Equation (8)) as listed in Table 2. Via this simulation, we will test the influence of
 394 non-linearity of signal on unweighted linear fit rates, which are derived by fitting a linear
 395 regression to the SBAS-derived deformation time series.

396 Like the previous test of a linear signal only, the simulated deformation time series is
 397 generated by first applying Equation (8) for all 1,000 pixels prior to forming 8,778 noise-free
 398 interferograms and applying simulated residual interferogram noise. The networks shown in
 399 Table 4 are then applied in sequence to select corresponding stacks of interferograms, which
 400 are then utilized with the SBAS method. Both the unweighted linear fit and unweighted LS
 401 methods are subsequently adopted to derive linear rates and annual amplitudes. Additionally,
 402 the RMSs between simulated and SBAS-derived time series are calculated.

403 Figure 4 shows unweighted linear-fit rates computed using the linear rates from Table
 404 2 and the simulated noise in Table 3. These results in Figure 4 reflect the influence of signal

405 non-linearity on linear-fit rates through biases in rate errors, particularly the case of large
 406 annual amplitudes, i.e., strongly non-linear, (cf. Figure 4 between red lines and coloured
 407 polylines). This is due to the inappropriate functional model used here to derive the linear
 408 rates, i.e., linear regression, which is applied to linear plus annual simulated signal.
 409



410 Figure 4. Comparison of unweighted linear-fit rates from linear plus annual signals. From left
 411 to right are simulated signal cases 1 to 3 (Table 2) contaminated by various simulated residual
 412 interferogram noise. Black, green and blue polylines indicate the results computed from
 413 simulated data with noise cases A, B and C (Table 3). Red horizontal lines represent the
 414 simulated rates. The black dashed box in the left panel used to contrast between positive and
 415 negative rates that indicates contradictory trends in some cases.

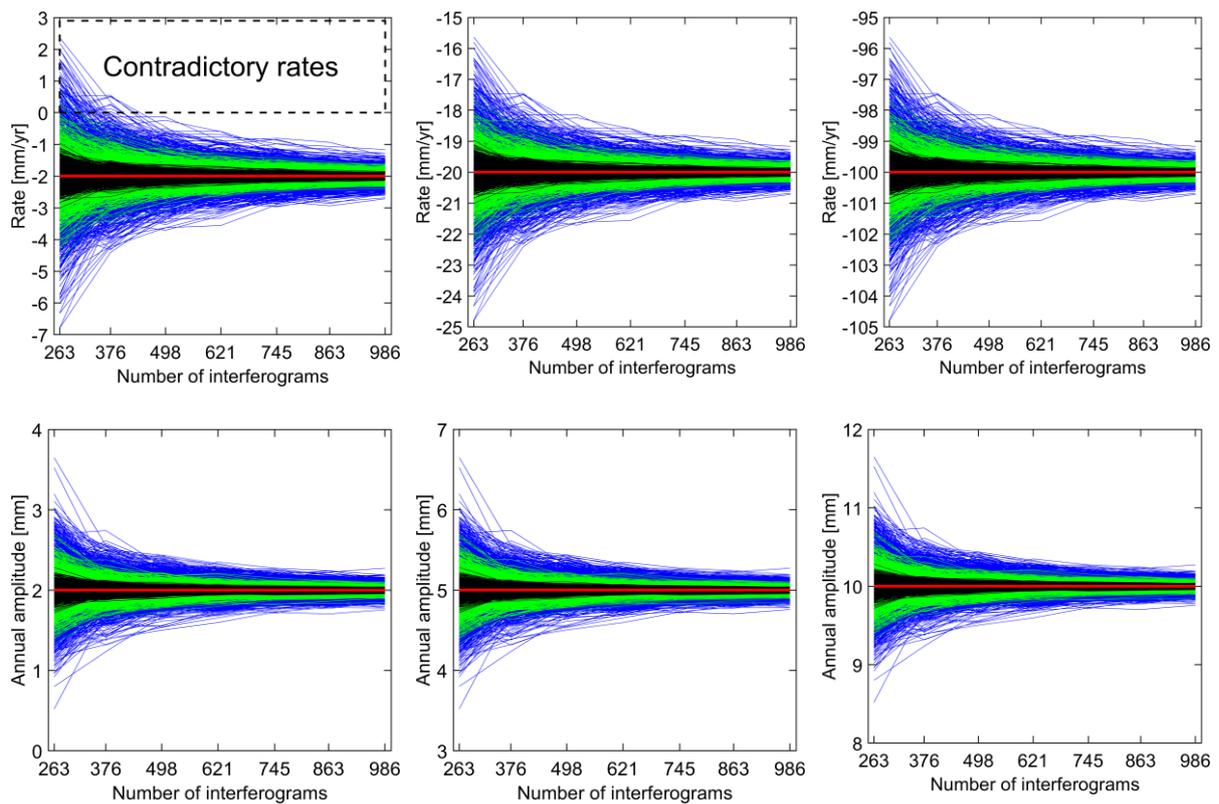
416

417 The simulated signal function is known (Equation (8)), so we adopt this for estimating
 418 both rates and annual amplitudes utilizing unweighted LS (Figure 5). The results indicate
 419 similar behavior as that in the case of solely linear signals (cf. Figure 5 (top) with Figure 2)
 420 and those with biases removed (cf. Figure 5 (top) and Figure 4). Again, this is attributable to
 421 the SBAS method in which the results computed depend on the configuration of the network
 422 and residual interferogram noise but not the deformation rate. Also, it is due to the more
 423 appropriate functional model used to obtain the linear rates where the influence of the signal

424 non-linearity cancel out. It is therefore an important warning that a suitable function should be
 425 utilized to calculate linear rates in case the Earth's surface experiences non-linear
 426 deformation, particularly in strongly non-linear cases.

427 In the case of applying LS estimation with an appropriate function, not only the linear
 428 rate, but also its accompanying parameters, e.g., the annual amplitude in this study, will be
 429 obtained. This is shown in Figure 5 (bottom), where the computed annual amplitudes indicate
 430 that more interferograms in the SBAS network result in more accurate LS estimation of the
 431 annual amplitude. In addition, the errors in those computed parameters are dependent on the
 432 SBAS network configuration and residual interferogram noise, but not the signal magnitude.

433



434 Figure 5. Comparison of unweighted LS rates (top panel) and annual amplitudes (bottom
 435 panel) computed from linear plus annual signals. From left to right correspond to simulated
 436 signal cases 1 to 3 (Table 2) contaminated by various simulated residual interferogram noise.

437 Black, green and blue polylines indicate the results computed from simulated data with noise
438 cases A, B and C (Table 3). Red horizontal lines represent the simulated rates or annual
439 amplitudes. The black dashed box in the top-left panel used to contrast between positive and
440 negative rates that indicates contradictory trends in some cases.

441

442 **6. Influence of data gaps on SBAS-derived rates**

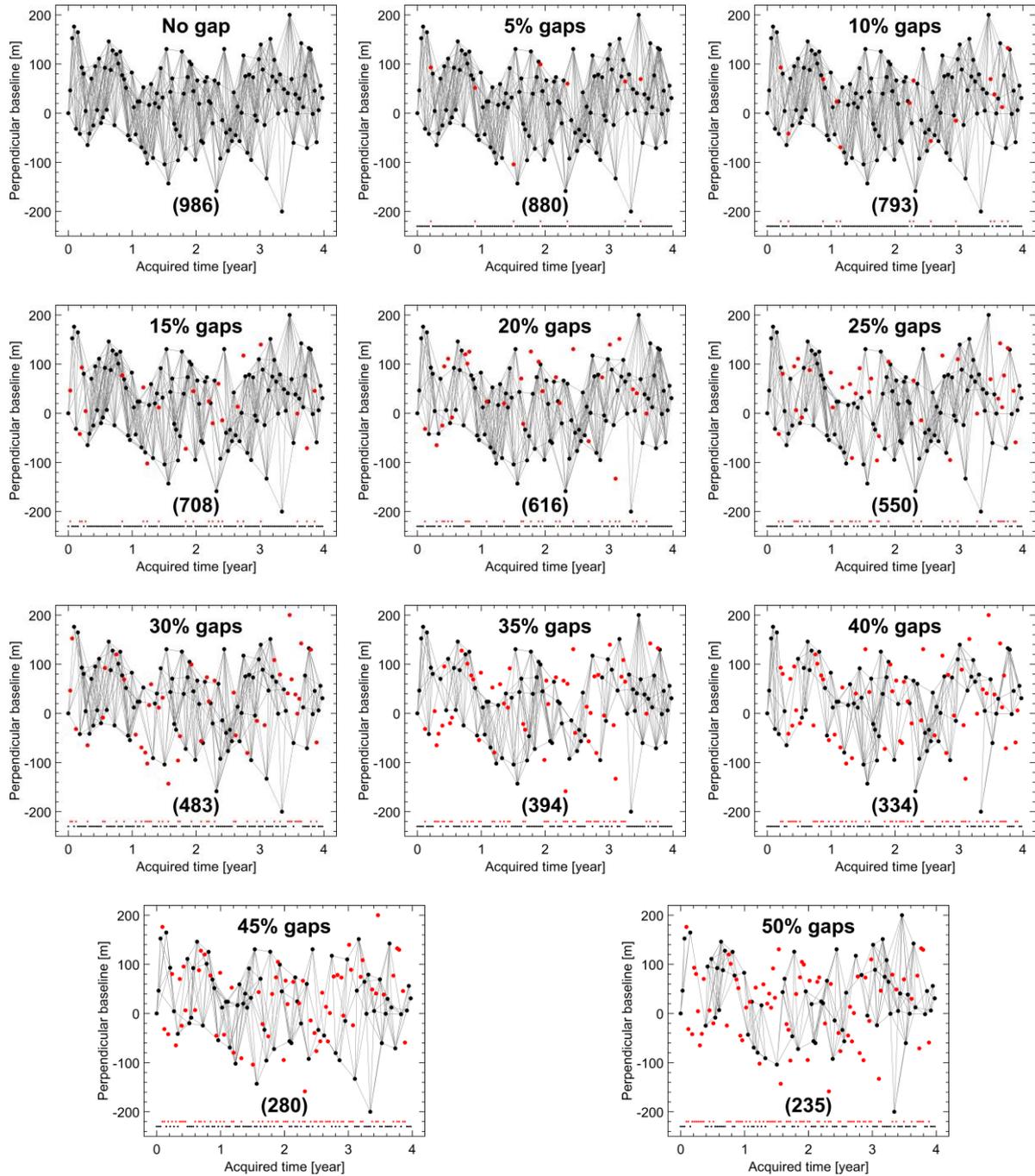
443 In this Section, we study the influence of SAR data gaps on SBAS-retrieved rates. This is
444 motivated by the likelihood of irregular temporal sampling of SAR data due to scheduling or
445 other technical issues, such as decorrelation during winter snow cover. We now conduct
446 simulations with a network of 986 interferograms formed by applying a temporal baseline
447 threshold of 88 days (~ 3 months, Table 4), with two scenarios of data gaps. In the first
448 scenario, missing images are due to technical and/or scheduling issues, which are considered
449 random, and, in the second scenario, missing images are chosen in the northern winter season
450 which are assumed to have low coherence due to extreme weather.

451

452 *6.1. Random data gaps*

453 In this Sub-section, we assume there are, in turn, 5%, 10%, ..., 50% of acquisitions missing
454 from our simulated time series. First, missing images are randomly chosen. Interferograms
455 having connections with those missing images are subsequently identified and eliminated
456 from the original list of 986 interferograms. Figure 6 compares the network without gaps and
457 those corresponding to various amount of gaps in percentage from 5% to 50% with an
458 increment of 5%.

459



460 Figure 6. Comparison of the interferogram network gaps in percentage. Gray lines indicate
 461 InSAR interferograms connecting images denoted by black dots. Red dots indicate missing
 462 images (i.e., gaps). The number under each network refers to the number of interferograms.

463

464

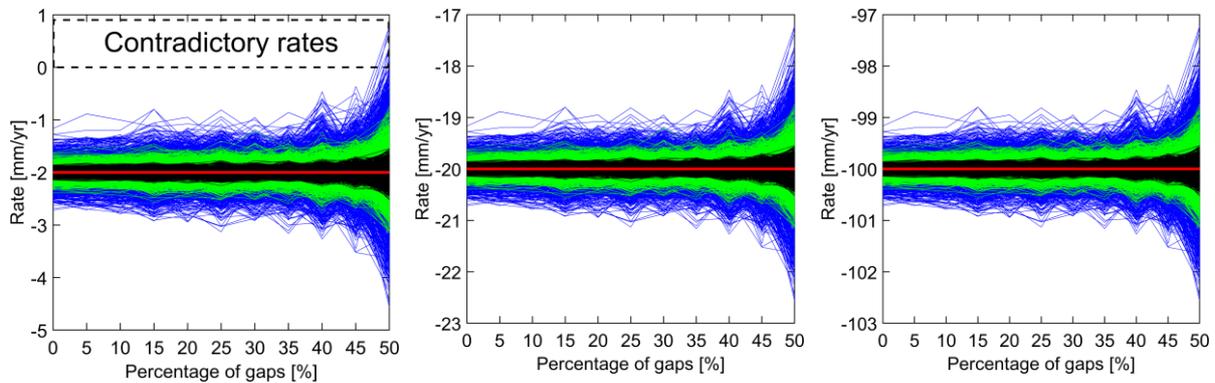
465

Here, we use the same linear plus annual signals as those used in Section 5.2 according to simulated signal cases shown in Table 2. For each network shown in Figure 6,

466 the SBAS approach in GIAN-T is applied to all 1,000 pixels in which the deformation time
 467 series at each pixel is derived. The unweighted LS is then applied to calculate the deformation
 468 rates and the RMSs of the difference between simulated and SBAS-derived time series are
 469 then calculated.

470 Figure 7 compares SBAS-derived unweighted LS rates between the SBAS network
 471 with no gaps and those of different percentages of data gaps. Figure 8 shows the
 472 corresponding RMSs of the difference between simulated and SBAS-derived deformation
 473 time series. These RMSs are the same for all three cases of linear plus annual signal (Table 2).
 474 Figure 7 and Figure 8 confirm that data gaps have an effect on the retrieved rates and RMSs
 475 with a noticeably larger influence in cases of higher gap percentages, particularly the 50%
 476 case. Contradictory trends are obtained for some pixels the case of large residual
 477 interferogram noise and low magnitude rates (Figure 7, left). This is likely caused by a weak
 478 SBAS network configuration (see Figure 6 with the 50% gaps case).

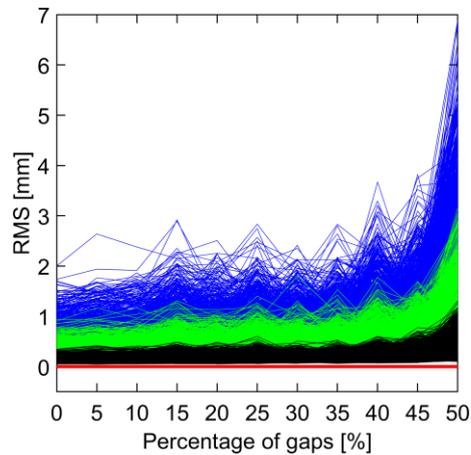
479



480 Figure 7. Comparison of unweighted LS rates computed from linear plus annual signals
 481 between the interferogram network of no gaps and those with randomly chosen gaps of
 482 various percentages. From left to right correspond to simulated signal cases 1 to 3 (Table 2)
 483 contaminated by various simulated residual interferogram noise. Black, green and blue
 484 polylines indicate the results computed from simulated data with noise cases A, B and C

485 (Table 3). Red horizontal lines represent the simulated rates. The black dashed box in the left
 486 panel used to contrast between positive and negative rates that indicates contradictory trends
 487 in some cases.

488



489 Figure 8. Comparison of the RMSs of the difference between simulated and SBAS-derived
 490 deformation time series of all pixels between the SBAS interferogram network of no gap and
 491 those with random gaps. Black, green and blue polylines indicate the results computed from
 492 simulated data with noise cases A, B and C (Table 3).

493

494 The influence of random data gaps on the errors in rate determination and the RMSs of
 495 the difference between simulated and SBAS-derived deformation time series is caused by a
 496 reduction in the number of interferograms when the percentage of gaps increases. However, a
 497 reduction in interferograms in the SBAS network can be caused by random data gaps (Figures
 498 7 and 8) or by changing the temporal baseline thresholds (as shown in Section 5). We
 499 compare errors resulting from fewer interferograms in a SBAS network due to (1) random
 500 gaps and (2) temporal baseline thresholds in Figure 9 (cf. blue and green polylines). This
 501 demonstrates the role of the network configuration, where a network may have the same

502 number of interferograms, but will have higher errors depending on which interferograms are
503 selected.

504 The random gap scenario results in more redundant interferograms, making the

505 network more robust, especially in the case of noisier time series (Figure 9, right plots).

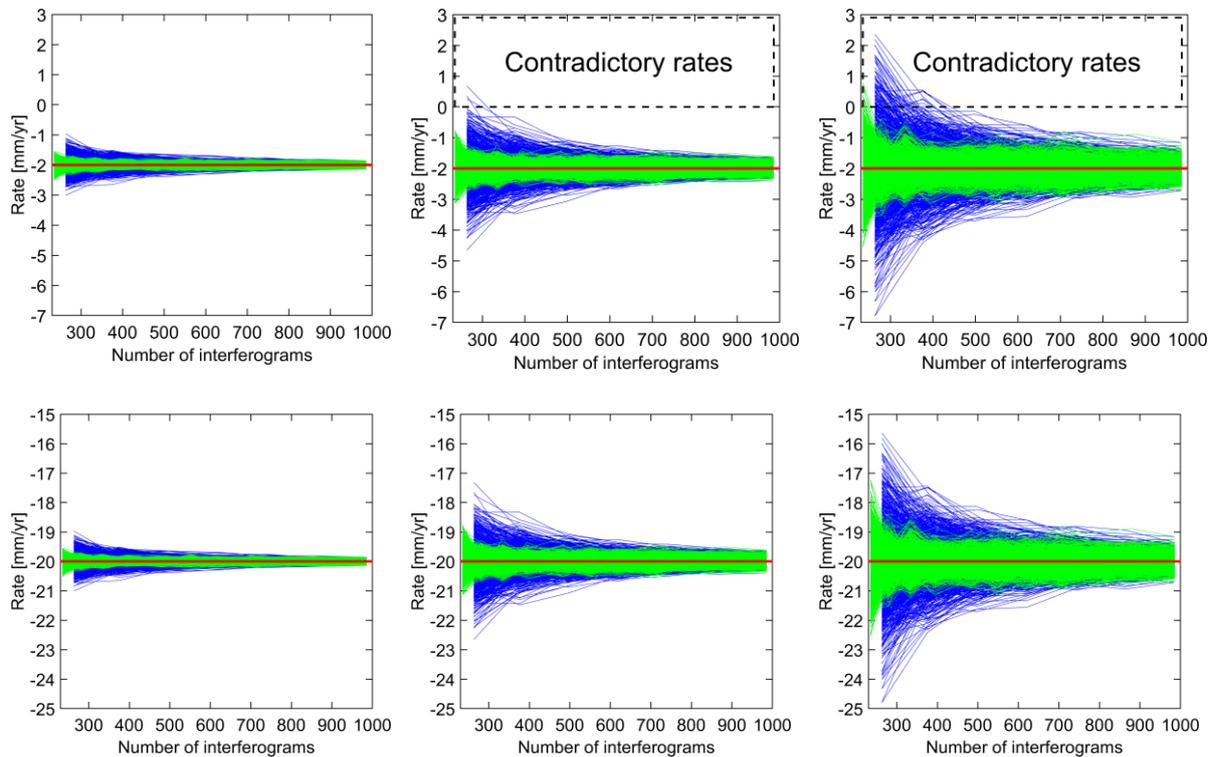
506 Therefore, in this case of randomly selected data gaps, mixed interferograms covering both

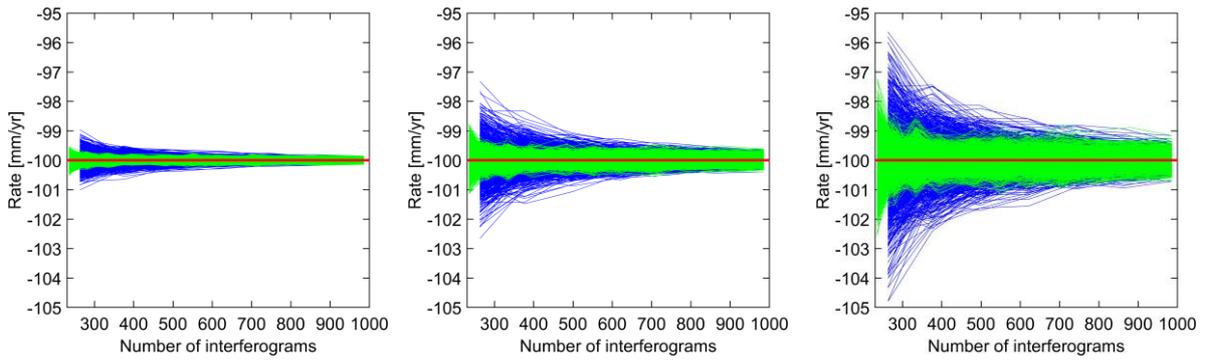
507 short and long time spans makes the network more robust in recovering the deformation

508 signal compared to the case of no gaps in which only short-time interferograms are chosen,

509 which are limited by the threshold.

510





511 Figure 9. The influence of the change in number of interferograms chosen by various
 512 temporal baseline thresholds (blue) and due to random data gaps (green) on SBAS-derived
 513 unweighted LS rates. From top to bottom: simulated signal cases 1, 2 and 3 (Table 2). From
 514 left to right: residual interferogram noise cases A to C (Table 3). Black dashed boxes in the
 515 top panel used to contrast between positive and negative rates that indicates contradictory
 516 trends in some cases.

517

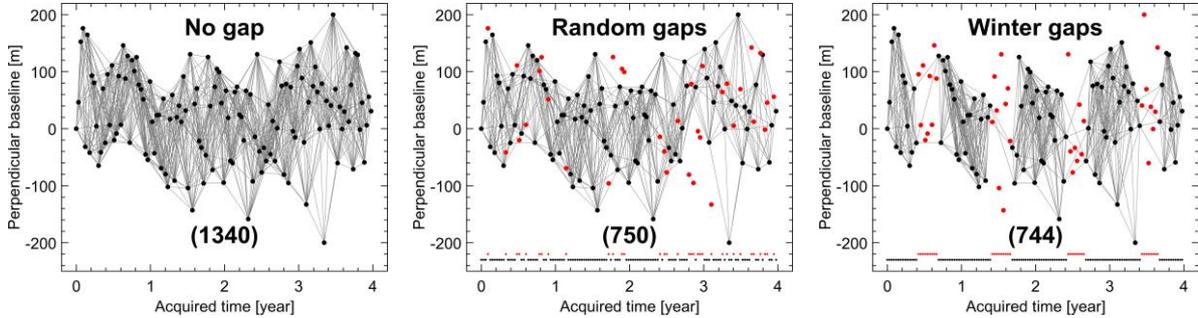
518 6.2. “Winter” data gaps

519 The previous test on data gaps in Section 6.1 is based on the fact that SAR data is missing
 520 sometime due to technical and/or scheduling issues, which we consider random. There is an
 521 alternative situation where there may be “user-defined” data gaps in which data missing is due
 522 to, e.g., very low coherence caused, for instance, by snow cover. We term this situation
 523 “winter data gaps” where all images acquired in the winter season (we use December to
 524 February for the Northern Hemisphere) are removed (Figure 10).

525 The results of this simulation experiment are shown in Figure 11. We compute
 526 unweighted LS rates and RMSs of the difference between simulated and SBAS-derived
 527 deformation time series for networks with no gaps, random data gaps and “winter data gaps”,
 528 with the latter two having the same number of images. To avoid a disconnection in the SBAS
 529 network, we apply a network of 1,340 interferograms formed by applying a temporal baseline

530 threshold of 121 days (~ 4 months), instead of ~ 3 months as in Section 6.1, and a
 531 perpendicular baseline threshold of 200 meters.

532

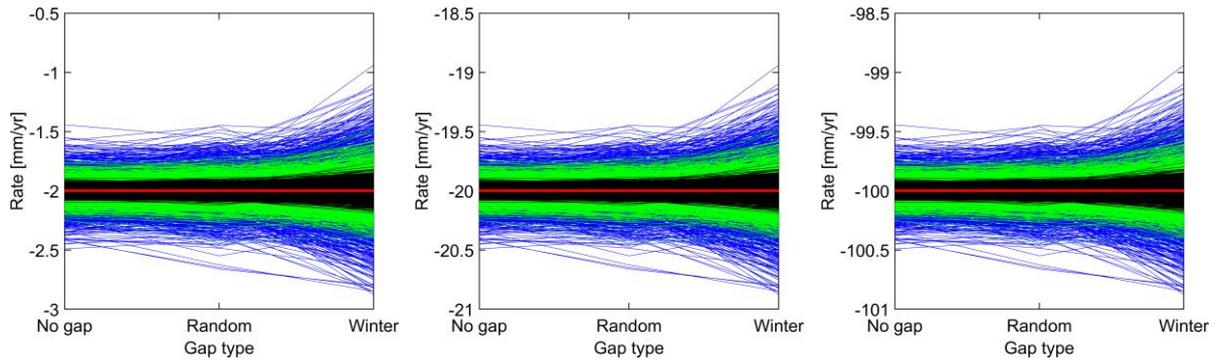


533 Figure 10. Interferogram networks without (left) and with (middle, right) missing images. The
 534 number of missing images is 34 out of 133 corresponding to about 25%, which are selected
 535 randomly (middle) and in the northern winter season (right). The networks are formed using a
 536 temporal baseline threshold of ~ 4 months and a perpendicular baseline threshold of 200
 537 meters. Gray lines indicate interferograms, with images denoted by black dots. Red dots
 538 indicate missing images (i.e., gaps). The number under each network refers to the number of
 539 interferograms.

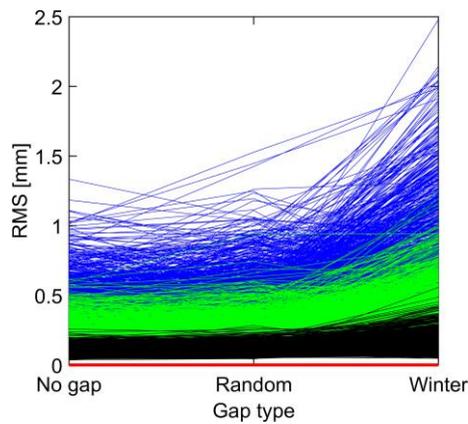
540 Figure 11 compares unweighted LS rates for each network with RMSs between
 541 simulated and SBAS-derived deformation time series shown in Figure 12. Figure 10 shows
 542 the number of missing images is the same between the two cases of data gaps, which is 34 out
 543 of 133, and, though the missing images are selected differently, the number of interferograms
 544 linking the remaining images are nearly the same; 750 for random gaps and 744 winter gaps.
 545 However, the influence of these two different data gap cases are distinct with the “winter”
 546 gaps having a larger influence, as confirmed by both retrieved rates in Figure 11 and RMSs in
 547 Figure 12.

548 This is caused by the strength of the network configuration, which is more robust with
 549 interferograms at regular intervals in the random gaps network but with “blocks” of gaps in

550 the “winter” case, leading to a less robust network (cf. Figure 10 (middle) and (right)). This
 551 alerts users that, in addition to the effect of fewer interferograms and gap percentages, the
 552 strength of network configuration is another factor influencing the SBAS results, in which one
 553 should try to design a SBAS network that does not contain long gaps in the time series.
 554



555 Figure 11. Comparison of unweighted LS rates computed from linear plus annual signals
 556 according to interferogram networks with no gaps, random gaps and “winter” gaps. The
 557 networks adopt a temporal baseline threshold of ~ 4 months and a perpendicular baseline
 558 threshold of 200 meters. From left to right are simulated signal cases 1 to 3 (Table 2). Black,
 559 green and blue polylines indicate the results computed from simulated data with noise cases
 560 A, B, and C (Table 3). Red horizontal lines represent the simulated rate.
 561



562 Figure 12. Comparison of the RMSs of the difference between simulated and SBAS-derived
 563 deformation time series for all pixels between the interferogram networks of no gaps and
 564 those with randomly chosen gaps and “winter” gaps corresponding to ~25% missing images.
 565 Black, green and blue polylines indicate the results computed from simulated data with noise
 566 cases A, B and C (Table 3).

567

568 **7. Optimal design of InSAR SBAS networks using redundancy numbers**

569 As has been demonstrated in Section 5, a spurious deformation trend (uplift instead of
 570 simulated subsidence) can be retrieved by applying SBAS, particularly in the case of small
 571 deformation in relation to large residual error and noise (i.e., a low SNR). By using more
 572 interferograms, the rate error can be decreased as the redundancy in the network is increased.
 573 However, an increased number of interferograms will also result in a higher computational
 574 burden. In this Section, “optimal” network design from geodesy is adopted for InSAR based
 575 on redundancy or r -numbers (Section 3). The motivation here is to investigate the relation
 576 between RMSs of the difference between simulated and SBAS-derived deformation time
 577 series, number of selected interferograms and the redundancy number.

578 Here, we test interferogram networks determined by combinations of temporal
 579 baseline thresholds, from one month to four years long, with a one-month increment, and
 580 perpendicular baseline thresholds of 100 meters, 200 meters and 300 meters. As a result, 144
 581 networks are formed with the minimum and maximum number of interferograms being 251
 582 and 8,778, respectively. Equation (6) is then applied to each of these networks to compute the
 583 r -numbers.

584 The reliability matrix \mathbf{R} computed from Equation (6) contains the r -numbers located
 585 on its diagonal (r_i). The objective of this optimization is to maximize these r -numbers by
 586 using their minimum value to represent the reliability of a network so that the r -numbers of

587 all measurements in that network are larger or equal to this minimum value. The r -number of
 588 a network is thus defined as:

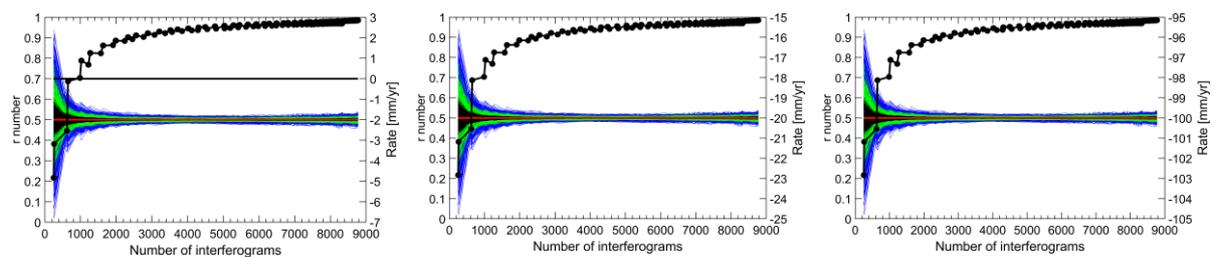
$$r = \min(r_i) \quad (11)$$

589

590 The SBAS method was then applied to derive deformation time series for all 1,000
 591 pixels, again using GIANt. We examine the same linear plus annual signals as those tested in
 592 Sections 5.1 and 6 (Table 2). The unweighted LS method is then utilized to derive SBAS-
 593 retrieved rates and the RMSs of the difference between simulated and SBAS-derived
 594 deformation time series are calculated.

595 The dependence of computed r -numbers and SBAS-retrieved annual rates on the
 596 number of selected interferograms are shown in Figure 13, where the change in SBAS-
 597 derived unweighted LS rates presents the same patterns among the three cases (Table 2) of
 598 simulated signals. Furthermore, the higher the r -number, the closer the agreement between
 599 simulated and SBAS-retrieved rates. The two rates are, in particular, nearly identical when the
 600 r -numbers are greater than ~ 0.9 .

601



602 Figure 13. The dependence of the r -numbers and SBAS-derived unweighted LS rates for
 603 1,000 pixels on the number of chosen interferograms with various linear plus annual signals.
 604 From left to right are simulated signal cases 1 to 3 (Table 2). Black, green and blue polylines

605 show the results for noise cases A, B and C (Table 3). Red horizontal lines represent the
606 simulated rates.

607

608 The dependence of the RMSs of the difference between simulated and SBAS-retrieved
609 deformation time series on the number of chosen interferograms are shown in Figure 14 (left)
610 for all three cases of simulated noise (Table 3). The r -number increases as the number of
611 interferograms increases, constrained by temporal baseline thresholds, and a reduction in the
612 RMSs. The RMSs decrease from a small r -number until ~ 0.8 , after which the change
613 becomes negligible.

614 We then apply 1/10 RMS as a trade-off value to identify the “optimal” r -number in
615 which a network with a minimal number of interferograms selected and with all RMSs
616 smaller than 1/10 noise range, which are 0.2 mm, 0.5 mm and 1.0 mm for the simulated
617 residual interferogram noise ranges shown in Table 3. Recall that the RMSs are dependent on
618 the SBAS network and residual interferogram noise but not signal magnitude (Figure 3).
619 Thus, Figure 14 indicates an “optimal” r -number being ~ 0.86 for a SBAS network of 1,911
620 interferograms, suggesting r -numbers between ~ 0.8 and ~ 0.9 to be a suitable range for the
621 “optimal” design of SBAS networks.

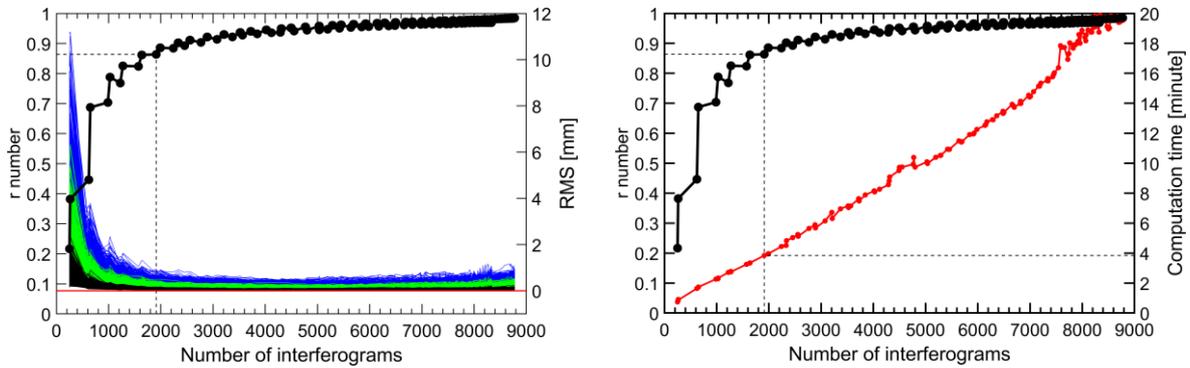
622 The efficiency of the obtained “optimal” r -numbers are confirmed by not only the
623 RMS trade-off, but also the computation time as shown in Figure 14 (right), where the
624 network of 1,911 interferograms (for the “optimal” r -number) runs for less than four minutes
625 compared to nearly 20 minutes for the largest network of 8,778 interferograms. This shows
626 efficiency in processing time for the SBAS inversion step only. In reality, SBAS data
627 processing with a full workflow, which comprises additional steps of interferogram formation
628 and error correction (e.g., DEM, orbital and atmospheric errors) the time difference between

629 processing all 8,778 interferograms and the optimal 1,911 interferograms can be substantial.

630 Additionally, this “optimal” design of SBAS networks keeps the number of interferograms to

631 a minimum, which limits the disk storage space required.

632



633 Figure 14. (left): Comparison of the change in the r -numbers and the RMSs of the difference

634 between simulated and SBAS-derived deformation time series. Black, green and blue

635 polylines indicate the results computed from simulated data with noise cases A, B and C

636 (Table 3). Dashed lines indicate the “optimal” r -numbers corresponding to the networks with

637 smallest amount of interferograms chosen with all RMSs being smaller than the chosen trade-

638 off values of 1/10 of the residual interferogram noise (Table 3). (right): Comparison of the

639 change in the r -numbers according to SBAS network interferogram numbers and computation

640 time.

641

642 Figure 15 shows an example of the SBAS-derived deformation time series of a pixel

643 compared with a simulated linear plus annual signal of -2 mm/yr plus 2 mm (i.e., the

644 simulated signal case 1 in Table 2) according to selected cases of computed r -number of

645 ~ 0.2 , ~ 0.8 and ~ 0.9 . The results of applying residual noise cases A, B, and C (Table 3) are

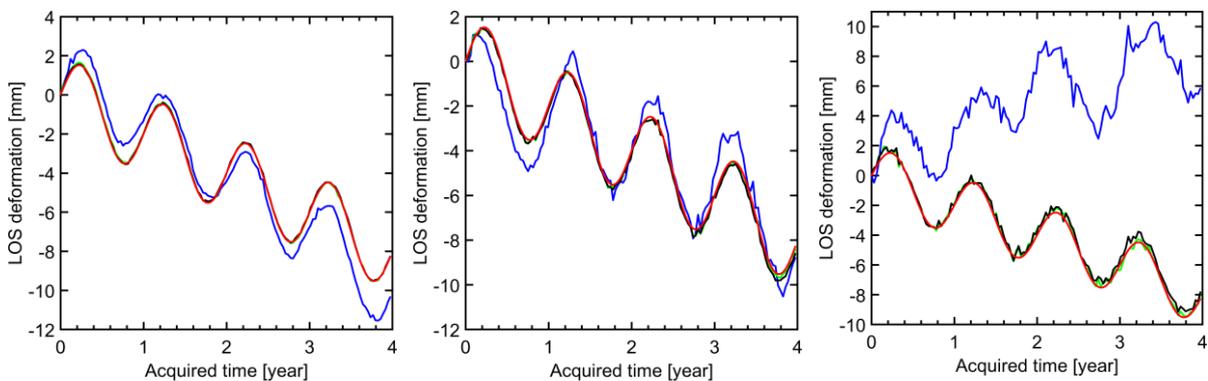
646 shown in Figure 15. In all cases, the results corresponding to the computed r -numbers of

647 ~ 0.8 and ~ 0.9 show close agreement with the simulated signal. In contrast, however, large

648 differences between simulated and SBAS-derived deformation time series are shown in the
 649 case when the r -number is ~ 0.2 (251 interferograms), particularly in the case of large
 650 residual interferogram noise (i.e., bounded within $[-10; +10]$ mm) where the difference in
 651 both its trend (i.e., deformation or uplift) and magnitude is shown (cf. Figure 15 (right))
 652 between the blue and red polylines).

653 Figure 15 shows that if the r -number is too small (< 0.2), spurious or even
 654 contradictory rates can result, as was shown similarly in Section 5. Therefore, caution must be
 655 exercised when using InSAR to detect small rates of deformation in the presence of large
 656 noise (low SNR). This is where the redundancy number may be of assistance in gauging the
 657 reliability of the estimated rates. This also shows that, though the reliability of a network is
 658 relevant to its ability to detect and resist against gross errors, in this specific case of InSAR
 659 SBAS networks, a good agreement between the r -numbers and errors in rate determination is
 660 present that is useful for “optimal” design of InSAR SBAS networks.

661



662 Figure 15. An example of simulated and SBAS-derived deformation time series of simulated
 663 signal of linear rate plus annual amplitude of -2 mm/yr plus 2 mm (simulated signal case 1 in
 664 Table 2) contaminated by residual interferogram noise cases A (left), B (middle) and C (right)
 665 as listed in Table 3. The results are computed from applying various SBAS interferogram
 666 networks corresponding to computed r -numbers of ~ 0.2 (251 interferograms, blue polylines),

667 ~0.8 (1,571 interferograms, green polylines), and ~0.9 (2,330 interferograms, black
668 polylines), respectively. Red polylines indicate the simulated deformation time series. The
669 blue line in the right panel shows the extreme case where spurious uplift is indicated, whereas
670 subsidence is simulated.

671

672 **8. Conclusions**

673 This study has used simulated Gaussian noise with zero mean applied to interferograms
674 computed from simulated linear and annual sinusoidal trends to demonstrate the effects of
675 interferometric noise on InSAR SBAS derived deformation. This extends to how different
676 SBAS network configurations may influence the estimated deformation rates. Different
677 simulated rates are tested (Table 2), including the addition of annual periodic amplitudes so as
678 to represent a range of real SAR data. A Monte Carlo simulation with 1,000 pixels for each
679 scenario was adopted. Firstly, we investigated the linear deformation signal, finding that the
680 SBAS linear-fit deformation trends were sensitive to both the magnitude of interferometric
681 noise and signal size. The unweighted linear-fit rate error was the same in both magnitude and
682 size for all rates if the same residual noise is applied to a given network. The trend may
683 become contradictory for small magnitude deformation where, for example a -2 mm/yr rate
684 could be estimated from the SBAS least squares or SVD method as a spurious uplift. This
685 contradictory result was shown when small temporal thresholds of 33 days or less were used,
686 which resulted in a less robust SBAS network configuration with fewer interferograms.

687 When we tested the linear plus annual periodic signal with interferometric noise, the
688 linear-fit rates were biased in the linear rate (from the 1,000 pixels) compared to the simulated
689 rate. Alternatively, when we estimated the rates using a more suitable periodic functional
690 model, rather than just linear regression in the presence of non-linear terms, the estimated

691 rates were not biased. This demonstrates the potential for errors to be introduced by using
692 simple linear regression when non-linear deformation may also be occurring.

693 Because one of the strengths of the SBAS method is to provide redundant small
694 interferogram baselines (in space and time), we simulated the effect of missing SAR
695 acquisitions in the time series. We presumed that these gaps in the time series would be (1)
696 random that may be due to satellite mission scheduling issues, or (2) blocks of missing
697 interferograms over, for example, a northern winter with snow covered ground that causes
698 decorrelation. Our simulation results indicate that “winter” gaps causes a larger error in the
699 estimated rates and in the RMSs of the differences between simulated and SBAS-derived
700 deformation time series than for random gaps resulting from missed acquisitions. However,
701 the RMS for both random gaps and no gaps were mostly 1 mm, while the winter gaps RMS
702 was generally <2 mm, suggesting that random gaps have little influence. This is highlighted
703 when random gaps are compared to temporal threshold limits, showing that for the same
704 number of interferograms, limiting temporal thresholds can cause errors of up to 6 mm/yr
705 with noisy simulated data, compared to ~3 mm/yr for random gaps when using similar
706 interferogram numbers. This suggests that it is the configuration of the SBAS network that is
707 more important, to the point that caution should be exercised when reducing the temporal
708 baseline to increase the coherence of the interferograms, because the trade-off may be a
709 geometrically weak SBAS network that is vulnerable to incorrect rate estimation in the
710 presence of noisy data and non-linear deformation.

711 We ran an additional simulation investigating whether redundancy numbers from
712 geodetic theory could be adapted to design an optimal SBAS network. The simulation results
713 suggest that r -values between ~0.8 and ~0.9 indicated a robust SBAS network design, and
714 that including more interferograms beyond this provided little improvement in the accuracy of
715 the rate estimation.

716 We conclude finally that SBAS network design can be critical to correctly estimate
717 deformation rates, particularly in the case of low signal to noise ratios, and where the
718 deformation may be non-linear. Notably, we found an alarming artifact in a couple of
719 different simulation scenarios, where uplift was indicated by the SBAS rather than true
720 simulated subsidence. It therefore appears that the configuration (network design) is more
721 important than simply the number of interferograms used, which is important given any limits
722 on computing resources. For this reason, we recommend the use of redundancy numbers to
723 help optimize SBAS network design.

724

725 **Acknowledgements**

726 Luyen Bui is supported by the Australia Awards Scholarships (AAS) provided for
727 postgraduate study. Will Featherstone's InSAR projects are supported financially by
728 Australian Research Council linkage project LP140100155, Landgate (the Western Australian
729 geodetic agency), the Western Australian Department of Water, and Curtin University. We
730 would like to thank California Institute of Technology (Caltech) for providing the source code
731 of the Generic InSAR Analysis Toolbox (GIAN-T). Finally, we thank the editor and two
732 anonymous reviewers for their thorough and constructive handling of our manuscript,
733 particularly the advice to consider non-linear signals and "winter" gaps.

734 **References**

- 735 Agram, P. S., Jolivet, R., Riel, B., Lin, Y. N., Simons, M., Hetland, E., Doin, M. P., & Lasserre, C. (2013). New
736 radar interferometric time series analysis toolbox released. *Eos, Transactions, American Geophysical*
737 *Union*, 94(7), 69-70. <https://doi.org/10.1002/2013EO070001>
738
- 739 Agram, P. S., Jolivet, R., & Simons, M. (2012). GIANt - The Generic InSAR Analysis Toolbox - User Manual.
740 Retrieved 09 June 2018, from http://earthdef.caltech.edu/attachments/download/15/GIANt_doc.pdf
741
- 742 Amelung, F., Galloway, D. L., Bell, J. W., Zebker, H. A., & Lacznik, R. J. (1999). Sensing the ups and downs
743 of Las Vegas: InSAR reveals structural control of land subsidence and aquifer-system deformation.
744 *Geology*, 27(6), 483-486. [https://doi.org/10.1130/0091-7613\(1999\)027<0483:STUADO>2.3.CO;2](https://doi.org/10.1130/0091-7613(1999)027<0483:STUADO>2.3.CO;2)
745
- 746 Amiri-Simkooei, A. R. (2001). Comparison of reliability and geometrical strength criteria in geodetic networks.
747 *Journal of Geodesy*, 75(4), 231-233. <https://doi.org/10.1007/s001900100170>
748
- 749 Amiri-Simkooei, A. R. (2004). A new method for second order design of geodetic networks: Aiming at high
750 reliability. *Survey Review*, 37(293), 552-560. <https://doi.org/10.1179/sre.2004.37.293.552>
751
- 752 Amiri-Simkooei, A. R., Asgari, J., Zangeneh-Nejad, F., & Zaminpardaz, S. (2012). Basic concepts of
753 optimization and design of geodetic networks. *Journal of Surveying Engineering*, 138(4), 172-183.
754 [https://doi.org/10.1061/\(ASCE\)SU.1943-5428.0000081](https://doi.org/10.1061/(ASCE)SU.1943-5428.0000081)
755
- 756 Baarda, W. (1968). A testing procedure for use in geodetic networks. Publication on Geodesy, New Series,
757 Volume 2, Number 5. In *Netherlands Geodetic Commission, Delft*.
758
- 759 Baldi, P., Casula, G., Cenni, N., Loddo, F., & Pesci, A. (2009). GPS-based monitoring of land subsidence in the
760 Po Plain (Northern Italy). *Earth and Planetary Science Letters*, 288(1-2), 204-212.
761 <https://doi.org/10.1016/j.epsl.2009.09.023>
762
- 763 Berardino, P., Fornaro, G., Lanari, R., & Sansosti, E. (2002). A new algorithm for surface deformation
764 monitoring based on small baseline differential SAR interferograms. *IEEE Transactions on Geoscience*
765 *and Remote Sensing*, 40(11), 2375-2383. <https://doi.org/10.1109/TGRS.2002.803792>
766
- 767 Berné, J. L., & Baselga, S. (2004). First-order design of geodetic networks using the simulated annealing
768 method. *Journal of Geodesy*, 78(1-2), 47-54. <https://doi.org/10.1007/s00190-003-0365-y>
769
- 770 Biggs, J., Wright, T., Lu, Z., & Parsons, B. (2007). Multi-interferogram method for measuring interseismic
771 deformation: Denali Fault, Alaska. *Geophysical Journal International*, 170(3), 1165-1179.
772 <https://doi.org/10.1111/j.1365-246X.2007.03415.x>
773
- 774 Bock, Y., Wdowinski, S., Ferretti, A., Novali, F., & Fumagalli, A. (2012). Recent subsidence of the Venice
775 Lagoon from continuous GPS and interferometric synthetic aperture radar. *Geochemistry, Geophysics,*
776 *Geosystems*, 13(3), Q03023. <https://doi.org/10.1029/2011GC003976>
777
- 778 Bombrun, L., Gay, M., Trouvé, E., Vasile, G., & Mars, J. (2009). DEM error retrieval by analyzing time series
779 of differential interferograms. *IEEE Geoscience and Remote Sensing Letters*, 6(4), 830-834, Article
780 5226568. <https://doi.org/10.1109/LGRS.2009.2026434>
781
- 782 Cavalié, O., Doin, M. P., Lasserre, C., & Briole, P. (2007). Ground motion measurement in the Lake Mead area,
783 Nevada, by differential synthetic aperture radar interferometry time series analysis: Probing the
784 lithosphere rheological structure. *Journal of Geophysical Research: Solid Earth*, 112(3), 1-18, Article
785 B03403. <https://doi.org/10.1029/2006JB004344>
786
- 787 Cavalié, O., Lasserre, C., Doin, M.-P., Peltzer, G., Sun, J., Xu, X., & Shen, Z.-K. (2008). Measurement of
788 interseismic strain across the Haiyuan fault (Gansu, China), by InSAR. *Earth and Planetary Science*
789 *Letters*, 275(3-4), 246-257. <https://doi.org/10.1016/j.epsl.2008.07.057>
790

- 791 Cavalié, O., Pathier, E., Radiguet, M., Vergnolle, M., Cotte, N., Walpersdorf, A., Kostoglodov, V., & Cotton, F.
792 (2013). Slow slip event in the Mexican subduction zone: Evidence of shallower slip in the Guerrero
793 seismic gap for the 2006 event revealed by the joint inversion of InSAR and GPS data. *Earth and*
794 *Planetary Science Letters*, 367, 52-60. <https://doi.org/10.1016/j.epsl.2013.02.020>
795
- 796 Chaussard, E., Wdowinski, S., Cabral-Cano, E., & Amelung, F. (2014). Land subsidence in central Mexico
797 detected by ALOS InSAR time-series. *Remote Sensing of Environment*, 140, 94-106.
798 <https://doi.org/10.1016/j.rse.2013.08.038>
799
- 800 Chen, M., Tomás, R., Li, Z., Motagh, M., Li, T., Hu, L., Gong, H., Li, X., Yu, J., & Gong, X. (2016). Imaging
801 land subsidence induced by groundwater extraction in Beijing (China) using satellite radar
802 interferometry. *Remote Sensing*, 8(6), Article 468. <https://doi.org/10.3390/rs8060468>
803
- 804 Crosetto, M., Monserrat, O., Cuevas-González, M., Devanthery, N., & Crippa, B. (2016). Persistent Scatterer
805 Interferometry: A review. *ISPRS Journal of Photogrammetry and Remote Sensing*, 115, 78-89.
806 <https://doi.org/10.1016/j.isprsjprs.2015.10.011>
807
- 808 Davis, J. L., Wernicke, B. P., & Tamisiea, M. E. (2012). On seasonal signals in geodetic time series. *Journal of*
809 *Geophysical Research: Solid Earth*, 117(B1), B01403. <https://doi.org/10.1029/2011JB008690>
810
- 811 Delacourt, C., Briole, P., & Achache, J. (1998). Tropospheric corrections of SAR interferograms with strong
812 topography. Application to Etna. *Geophysical Research Letters*, 25(15), 2849-2852.
813 <https://doi.org/10.1029/98GL02112>
814
- 815 Didova, O., Gunter, B., Riva, R., Klees, R., & Roesse-Koerner, L. (2016). An approach for estimating time-
816 variable rates from geodetic time series. *Journal of Geodesy*, 90(11), 1207-1221.
817 <https://doi.org/10.1007/s00190-016-0918-5>
818
- 819 Doin, M. P., Lasserre, C., Peltzer, G., Cavalié, O., & Doubre, C. (2009). Corrections of stratified tropospheric
820 delays in SAR interferometry: Validation with global atmospheric models. *Journal of Applied*
821 *Geophysics*, 69(1), 35-50. <https://doi.org/10.1016/j.jappgeo.2009.03.010>
822
- 823 Doin, M. P., Lodge, F., Guillaso, S., Jolivet, R., Lasserre, C., Ducret, G., Grandin, R., Pathier, E., & Pinel, V.
824 (2011, 19-23 Septempber). Presentation of the small baseline NSBAS processing chain on a case
825 example: the Etna deformation monitoring from 2003 to 2010 using Envisat data. *ESA Fringe*
826 *Symposium*, Frascati, Italy: ESA SP-697. 3434-3437
827 [https://earth.esa.int/documents/10174/1573056/Presentation_small_baseline_NSBAS_Etna_deformatio](https://earth.esa.int/documents/10174/1573056/Presentation_small_baseline_NSBAS_Etna_deformation_Envisat.pdf)
828 [n_Envisat.pdf](https://earth.esa.int/documents/10174/1573056/Presentation_small_baseline_NSBAS_Etna_deformation_Envisat.pdf)
829
- 830 Dzurisin, D., Lisowski, M., & Wicks, C. W. (2009). Continuing inflation at Three Sisters volcanic center, central
831 Oregon Cascade Range, USA, from GPS, leveling, and InSAR observations. *Bulletin of Volcanology*,
832 71(10), 1091-1110. <https://doi.org/10.1007/s00445-009-0296-4>
833
- 834 Elliott, J. R., Walters, R. J., England, P. C., Jackson, J. A., Li, Z., & Parsons, B. (2010). Extension on the Tibetan
835 plateau: Recent normal faulting measured by InSAR and body wave seismology. *Geophysical Journal*
836 *International*, 183(2), 503-535. <https://doi.org/10.1111/j.1365-246X.2010.04754.x>
837
- 838 Fattahi, H., & Amelung, F. (2013). DEM error correction in InSAR time series. *IEEE Transactions on*
839 *Geoscience and Remote Sensing*, 51(7), 4249-4259, Article 6423275.
840 <https://doi.org/10.1109/TGRS.2012.2227761>
841
- 842 Ferretti, A., Prati, C., & Rocca, F. (2001). Permanent scatterers in SAR interferometry. *IEEE Transactions on*
843 *Geoscience and Remote Sensing*, 39(1), 8-20. <https://doi.org/10.1109/36.898661>
844
- 845 Furuya, M., Mueller, K., & Wahr, J. (2007). Active salt tectonics in the Needles District, Canyonlands (Utah) as
846 detected by interferometric synthetic aperture radar and point target analysis: 1992-2002. *Journal of*
847 *Geophysical Research: Solid Earth*, 112(6), B06418, Article B06418.
848 <https://doi.org/10.1029/2006JB004302>
849

- 850 Gatelli, F., Guarnieri, A. M., Parizzi, F., Pasquali, P., Prati, C., & Rocca, F. (1994). The Wavenumber Shift in
851 SAR Interferometry. *IEEE Transactions on Geoscience and Remote Sensing*, 32(4), 855-865.
852 <https://doi.org/10.1109/36.298013>
853
- 854 Goel, K., & Adam, N. (2014). A distributed scatterer interferometry approach for precision monitoring of known
855 surface deformation phenomena. *IEEE Transactions on Geoscience and Remote Sensing*, 52(9), 5454-
856 5468, Article 6679273. <https://doi.org/10.1109/TGRS.2013.2289370>
857
- 858 Gong, W., Thiele, A., Hinz, S., Meyer, F. J., Hooper, A., & Agram, P. S. (2016). Comparison of small baseline
859 interferometric SAR processors for estimating ground deformation. *Remote Sensing*, 8(4), 1-26, Article
860 330. <https://doi.org/10.3390/rs8040330>
861
- 862 Grafarend, E. W., & Sansò, F. (1985). *Optimization and design of geodetic networks*. Springer-Verlag, Berlin
863 Heidelberg New York Tokyo. <https://doi.org/10.1007/978-3-642-70659-2>
864
- 865 Hetland, E. A., Musé, P., Simons, M., Lin, Y. N., Agram, P. S., & DiCaprio, C. J. (2012). Multiscale InSAR
866 Time Series (MInTS) analysis of surface deformation. *Journal of Geophysical Research: Solid Earth*,
867 117(B2), B02404. <https://doi.org/10.1029/2011JB008731>
868
- 869 Hooper, A. (2008). A multi-temporal InSAR method incorporating both persistent scatterer and small baseline
870 approaches. *Geophysical Research Letters*, 35(16), L16302, Article L16302.
871 <https://doi.org/10.1029/2008GL034654>
872
- 873 Hooper, A., Bekaert, D. P. S., Spaans, K., & Arkan, M. (2012). Recent advances in SAR interferometry time
874 series analysis for measuring crustal deformation. *Tectonophysics*, 514-517, 1-13.
875 <https://doi.org/10.1016/j.tecto.2011.10.013>
876
- 877 Hooper, A., Segall, P., & Zebker, H. A. (2007). Persistent scatterer interferometric synthetic aperture radar for
878 crustal deformation analysis, with application to Volcán Alcedo, Galápagos. *Journal of Geophysical*
879 *Research: Solid Earth*, 112(B7). <https://doi.org/10.1029/2006JB004763>
880
- 881 Hooper, A., Zebker, H. A., Segall, P., & Kampes, B. (2004). A new method for measuring deformation on
882 volcanoes and other natural terrains using InSAR persistent scatterers. *Geophysical Research Letters*,
883 31(23), L23611. <https://doi.org/10.1029/2004GL021737>
884
- 885 Jiang, L., Lin, H., Ma, J., Kong, B., & Wang, Y. (2011). Potential of small-baseline SAR interferometry for
886 monitoring land subsidence related to underground coal fires: Wuda (Northern China) case study.
887 *Remote Sensing of Environment*, 115(2), 257-268. <https://doi.org/10.1016/j.rse.2010.08.008>
888
- 889 Jolivet, R., Grandin, R., Lasserre, C., Doin, M. P., & Peltzer, G. (2011). Systematic InSAR tropospheric phase
890 delay corrections from global meteorological reanalysis data. *Geophysical Research Letters*, 38(17),
891 L17311, Article L17311. <https://doi.org/10.1029/2011GL048757>
892
- 893 Jolivet, R., Lasserre, C., Doin, M. P., Guillaso, S., Peltzer, G., Dailu, R., Sun, J., Shen, Z. K., & Xu, X. (2012).
894 Shallow creep on the Haiyuan fault (Gansu, China) revealed by SAR interferometry. *Journal of*
895 *Geophysical Research: Solid Earth*, 117(B6), B06401. <https://doi.org/10.1029/2011JB008732>
896
- 897 Kim, J. W., Lu, Z., Jia, Y., & Shum, C. K. (2015). Ground subsidence in Tucson, Arizona, monitored by time-
898 series analysis using multi-sensor InSAR datasets from 1993 to 2011. *ISPRS Journal of*
899 *Photogrammetry and Remote Sensing*, 107, 126-141. <https://doi.org/10.1016/j.isprsjprs.2015.03.013>
900
- 901 Koch, K. R. (1985). First order design: Optimization of the configuration of a network by introducing small
902 position changes. In E. W. Grafarend & F. Sansò (Eds.), *Optimization and design of geodetic networks*
903 (pp. 56-73). Springer-Verlag, Berlin Heidelberg New York Tokyo. <https://doi.org/10.1007/978-3-642-70659-2>
904
- 905 Kohlhasse, A. O., Feigl, K. L., & Massonnet, D. (2003). Applying differential InSAR to orbital dynamics: A new
906 approach for estimating ERS trajectories. *Journal of Geodesy*, 77(9), 493-502.
907 <https://doi.org/10.1007/s00190-003-0336-3>
908

- 909
910 Kroese, D. P., Brereton, T., Taimre, T., & Botev, Z. I. (2014). Why the Monte Carlo method is so important
911 today. *Wiley Interdisciplinary Reviews: Computational Statistics*, 6(6), 386-392.
912 <https://doi.org/10.1002/wics.1314>
913
- 914 Kuang, S. (1993). On optimal design of levelling networks. *Australian Surveyor*, 38(4), 257-273.
915 <https://doi.org/10.1080/00050326.1993.10438874>
916
- 917 Kuang, S. (1996). *Geodetic network analysis and optimal design: concepts and applications* (p. 368). [Book].
918 Ann Arbor Press, Chelsea, Michigan.
919
- 920 Lanari, R., Casu, F., Manzo, M., Zeni, G., Berardino, P., Manunta, M., & Pepe, A. (2007). An overview of the
921 Small Baseline Subset algorithm: A DInSAR technique for surface deformation analysis. *Pure and*
922 *Applied Geophysics*, 164(4), 637-661. <https://doi.org/10.1007/s00024-007-0192-9>
923
- 924 Lanari, R., Mora, O., Manunta, M., Mallorquí, J. J., Berardino, P., & Sansosti, E. (2004). A small-baseline
925 approach for investigating deformations on full-resolution differential SAR interferograms. *IEEE*
926 *Transactions on Geoscience and Remote Sensing*, 42(7), 1377-1386.
927 <https://doi.org/10.1109/TGRS.2004.828196>
928
- 929 Lauknes, T. R., Zebker, H. A., & Larsen, Y. (2011). InSAR deformation time series using an L1-norm small-
930 baseline approach. *IEEE Transactions on Geoscience and Remote Sensing*, 49(1 PART 2), 536-546,
931 Article 5512638. <https://doi.org/10.1109/TGRS.2010.2051951>
932
- 933 Lee, C. W., Lu, Z., & Jung, H. S. (2012). Simulation of time-series surface deformation to validate a multi-
934 interferogram InSAR processing technique. *International Journal of Remote Sensing*, 33(22), 7075-
935 7087. <https://doi.org/10.1080/01431161.2012.700137>
936
- 937 Lin, Y. N. N., Simons, M., Hetland, E. A., Muse, P., & Dicaprio, C. (2010). A multiscale approach to estimating
938 topographically correlated propagation delays in radar interferograms. *Geochemistry, Geophysics,*
939 *Geosystems*, 11(9), 1-17, Article Q09002. <https://doi.org/10.1029/2010GC003228>
940
- 941 López-Quiroz, P., Doin, M.-P., Tupin, F., Briole, P., & Nicolas, J.-M. (2009). Time series analysis of Mexico
942 City subsidence constrained by radar interferometry. *Journal of Applied Geophysics*, 69(1), 1-15.
943 <https://doi.org/10.1016/j.jappgeo.2009.02.006>
944
- 945 Lubitz, C., Motagh, M., Wetzel, H. U., & Kaufmann, H. (2013). Remarkable urban uplift in Staufen im
946 Breisgau, Germany: Observations from terraSAR-X InSAR and leveling from 2008 to 2011. *Remote*
947 *Sensing*, 5(6), 3082-3100. <https://doi.org/10.3390/rs5063082>
948
- 949 Lundgren, P., Usai, S., Sansosti, E., Lanari, R., Tesauro, M., Fornaro, G., & Berardino, P. (2001). Modeling
950 surface deformation observed with synthetic aperture radar interferometry at Campi Flegrei caldera.
951 *Journal of Geophysical Research: Solid Earth*, 106(B9), 19355-19366, Article 2001jb000194.
952 <https://doi.org/10.1029/2001JB000194>
953
- 954 Mora, O., Mallorquí, J. J., & Broquetas, A. (2003). Linear and nonlinear terrain deformation maps from a
955 reduced set of interferometric SAR images. *IEEE Transactions on Geoscience and Remote Sensing*,
956 41(10 PART I), 2243-2253. <https://doi.org/10.1109/TGRS.2003.814657>
957
- 958 Motagh, M., Djamour, Y., Walter, T. R., Wetzel, H. U., Zschau, J., & Arabi, S. (2007). Land subsidence in
959 Mashhad Valley, northeast Iran: Results from InSAR, levelling and GPS. *Geophysical Journal*
960 *International*, 168(2), 518-526. <https://doi.org/10.1111/j.1365-246X.2006.03246.x>
961
- 962 Murray, K. D., Bekaert, D. P. S., & Lohman, R. B. (2019). Tropospheric corrections for InSAR: Statistical
963 assessments and applications to the Central United States and Mexico. *Remote Sensing of Environment*,
964 232, Article 111326. <https://doi.org/10.1016/j.rse.2019.111326>
965
- 966 Murray, K. D., & Lohman, R. B. (2018). Short-lived pause in Central California subsidence after heavy winter
967 precipitation of 2017. *Science Advances*, 4(8), Article eaar8144. <https://doi.org/10.1126/sciadv.aar8144>

- 968
969 Neely, W. R., Borsa, A. A., & Silverii, F. (2020). GInSAR: A cGPS Correction for Enhanced InSAR Time
970 Series. *IEEE Transactions on Geoscience and Remote Sensing*, 58(1), 136-146, Article 8839742.
971 <https://doi.org/10.1109/TGRS.2019.2934118>
972
- 973 Osmanoglu, B., Dixon, T. H., Wdowinski, S., Cabral-Cano, E., & Jiang, Y. (2011). Mexico City subsidence
974 observed with persistent scatterer InSAR. *International journal of applied earth observation and*
975 *geoinformation*, 13(1), 1-12. <https://doi.org/10.1016/j.jag.2010.05.009>
976
- 977 Pepe, A., Ortiz, A. B., Lundgren, P. R., Rosen, P. A., & Lanari, R. (2011). The Stripmap–ScanSAR SBAS
978 Approach to Fill Gaps in Stripmap Deformation Time Series With ScanSAR Data. *IEEE Transactions*
979 *on Geoscience and Remote Sensing*, 49(12 PART 1), 4788-4804, Article 6045335.
980 <https://doi.org/10.1109/TGRS.2011.2167979>
981
- 982 Schmidt, D. A., & Bürgmann, R. (2003). Time-dependent land uplift and subsidence in the Santa Clara valley,
983 California, from a large interferometric synthetic aperture radar data set. *Journal of Geophysical*
984 *Research: Solid Earth*, 108(B9). <https://doi.org/10.1029/2002JB002267>
985
- 986 Schmitt, G. (1985a). Review of network designs: criteria, risk functions, design ordering. In E. W. Grafarend &
987 F. Sansò (Eds.), *Optimization and design of geodetic networks* (pp. 6-10). Springer-Verlag, Berlin
988 Heidelberg New York Tokyo. <https://doi.org/10.1007/978-3-642-70659-2>
989
- 990 Schmitt, G. (1985b). Third order design. In E. W. Grafarend & F. Sansò (Eds.), *Optimization and design of*
991 *geodetic networks* (pp. 122-131). Springer-Verlag, Berlin Heidelberg New York Tokyo.
992 <https://doi.org/10.1007/978-3-642-70659-2>
993
- 994 Shanker, P., Casu, F., Zebker, H. A., & Lanari, R. (2011). Comparison of persistent scatterers and small baseline
995 time-series InSAR results: A case study of the San Francisco bay area. *IEEE Geoscience and Remote*
996 *Sensing Letters*, 8(4), 592-596, Article 5692806. <https://doi.org/10.1109/LGRS.2010.2095829>
997
- 998 Short, N., Brisco, B., Couture, N., Pollard, W., Murnaghan, K., & Budkewitsch, P. (2011). A comparison of
999 TerraSAR-X, RADARSAT-2 and ALOS-PALSAR interferometry for monitoring permafrost
1000 environments, case study from Herschel Island, Canada. *Remote Sensing of Environment*, 115(12),
1001 3491-3506. <https://doi.org/10.1016/j.rse.2011.08.012>
1002
- 1003 Sowter, A., Bateson, L., Strange, P., Ambrose, K., & Fifiksyafiudin, M. (2013). DInSAR estimation of land
1004 motion using intermittent coherence with application to the south derbyshire and leicestershire
1005 coalfields. *Remote Sensing Letters*, 4(10), 979-987. <https://doi.org/10.1080/2150704X.2013.823673>
1006
- 1007 Teunissen, P. J. G. (1985). Zero Order Design: Generalized Inverses, Adjustment, the Datum Problem and S-
1008 Transformations. In E. W. Grafarend & F. Sansò (Eds.), *Optimization and design of geodetic networks*
1009 (pp. 11-55). Springer-Verlag, Berlin Heidelberg New York Tokyo. [https://doi.org/10.1007/978-3-642-](https://doi.org/10.1007/978-3-642-70659-2)
1010 [70659-2](https://doi.org/10.1007/978-3-642-70659-2)
1011
- 1012 Tong, X., Sandwell, D. T., & Smith-Konter, B. (2013). High-resolution interseismic velocity data along the San
1013 Andreas Fault from GPS and InSAR. *Journal of Geophysical Research: Solid Earth*, 118(1), 369-389.
1014 <https://doi.org/10.1029/2012JB009442>
1015
- 1016 Tymofyeyeva, E., & Fialko, Y. (2015). Mitigation of atmospheric phase delays in InSAR data, with application
1017 to the eastern California shear zone. *Journal of Geophysical Research: Solid Earth*, 120(8), 5952-5963.
1018 <https://doi.org/10.1002/2015JB011886>
1019
- 1020 Usai, S. (2003). A Least Squares Database Approach for SAR Interferometric Data. *IEEE Transactions on*
1021 *Geoscience and Remote Sensing*, 41(4 PART I), 753-760. <https://doi.org/10.1109/TGRS.2003.810675>
1022
- 1023 Vaniček, P., & Krakiwsky, E. J. (1986). *Geodesy: the concepts*. (2nd ed.). Elsevier Science Publisher,
1024 Amsterdam, The Netherlands.
1025

- 1026 Williams, S., Bock, Y., & Fang, P. (1998). Integrated satellite interferometry: Tropospheric noise, GPS estimates
1027 and implications for interferometric synthetic aperture radar products. *Journal of Geophysical*
1028 *Research: Solid Earth*, 103(B11), 27051-27067, Article 98jb02794. <https://doi.org/10.1029/98jb02794>
1029
- 1030 Yague-Martinez, N., Prats-Iraola, P., Gonzalez, F. R., Brcic, R., Shau, R., Geudtner, D., Eineder, M., & Bamler,
1031 R. (2016). Interferometric Processing of Sentinel-1 TOPS Data. *IEEE Transactions on Geoscience and*
1032 *Remote Sensing*, 54(4), 2220-2234, Article 7390052. <https://doi.org/10.1109/TGRS.2015.2497902>
1033
- 1034 Yu, C., Li, Z., & Penna, N. T. (2018). Interferometric synthetic aperture radar atmospheric correction using a
1035 GPS-based iterative tropospheric decomposition model. *Remote Sensing of Environment*, 204, 109-121.
1036 <https://doi.org/10.1016/j.rse.2017.10.038>
1037
- 1038 Yu, C., Li, Z., Penna, N. T., & Crippa, P. (2018). Generic Atmospheric Correction Model for Interferometric
1039 Synthetic Aperture Radar Observations. *Journal of Geophysical Research: Solid Earth*, 123(10), 9202-
1040 9222. <https://doi.org/10.1029/2017JB015305>
1041
- 1042 Yu, C., Penna, N. T., & Li, Z. (2017). Generation of real-time mode high-resolution water vapor fields from GPS
1043 observations. *Journal of Geophysical Research*, 122(3), 2008-2025.
1044 <https://doi.org/10.1002/2016JD025753>
1045
- 1046 Zebker, H. A., Rosen, P. A., & Hensley, S. (1997). Atmospheric effects in interferometric synthetic aperture
1047 radar surface deformation and topographic maps. *Journal of Geophysical Research B: Solid Earth*,
1048 102(B4), 7547-7563, Article 96JB03804. <https://doi.org/10.1029/96JB03804>
1049
- 1050 Zebker, H. A., & Villasenor, J. (1992). Decorrelation in interferometric radar echoes. *IEEE Transactions on*
1051 *Geoscience and Remote Sensing*, 30(5), 950-959. <https://doi.org/10.1109/36.175330>
1052
1053
- 1054

1055
1056
1057
1058
1059
1060
1061
1062
1063
1064
1065
1066
1067
1068
1069
1070
1071
1072
1073
1074
1075
1076
1077
1078

List of Figure Captions

Figure 1. Scatter plot of simulated perpendicular baseline history. Each black dot represents a SAR scene acquisition.

Figure 2. Comparison of rates computed by unweighted linear fit from combinations of different deformation signals. From left to right are simulated linear rate cases 1 to 3 (Table 2) contaminated by simulated residual interferogram noise. Black, green and blue polylines are SBAS derived rates computed from simulated data with simulated noise cases A to C, respectively (Table 3). Red horizontal lines represent the simulated rates. The black dashed box in the left panel is used to contrast between positive and negative rates that indicates contradictory trends.

Figure 3. Comparison of rate errors computed from different networks for four example pixels. The top, center and bottom rows correspond to simulated noise cases A, B and C (Table 3). Note the different scale on the y-axis for each noise case.

Figure 4. Comparison of unweighted linear-fit rates from linear plus annual signals. From left to right are simulated signal cases 1 to 3 (Table 2) contaminated by various simulated residual interferogram noise. Black, green and blue polylines indicate the results computed from simulated data with noise cases A, B and C (Table 3). Red horizontal lines represent the simulated rates. The black dashed box in the left panel used to contrast between positive and negative rates that indicates contradictory trends in some cases.

1079 Figure 5. Comparison of unweighted LS rates (top panel) and annual amplitudes (bottom
1080 panel) computed from linear plus annual signals. From left to right correspond to simulated
1081 signal cases 1 to 3 (Table 2) contaminated by various simulated residual interferogram noise.
1082 Black, green and blue polylines indicate the results computed from simulated data with noise
1083 cases A, B and C (Table 3). Red horizontal lines represent the simulated rates or annual
1084 amplitudes. The black dashed box in the top-left panel used to contrast between positive and
1085 negative rates that indicates contradictory trends in some cases.

1086

1087 Figure 6. Comparison of the interferogram network gaps in percentage. Gray lines indicate
1088 InSAR interferograms connecting images denoted by black dots. Red dots indicate missing
1089 images (i.e., gaps). The number under each network refers to the number of interferograms.

1090

1091 Figure 7. Comparison of unweighted LS rates computed from linear plus annual signals
1092 between the interferogram network of no gaps and those with randomly chosen gaps of
1093 various percentages. From left to right correspond to simulated signal cases 1 to 3 (Table 2)
1094 contaminated by various simulated residual interferogram noise. Black, green and blue
1095 polylines indicate the results computed from simulated data with noise cases A, B and C
1096 (Table 3). Red horizontal lines represent the simulated rates. The black dashed box in the left
1097 panel used to contrast between positive and negative rates that indicates contradictory trends
1098 in some cases.

1099

1100 Figure 8. Comparison of the RMSs of the difference between simulated and SBAS-derived
1101 deformation time series of all pixels between the SBAS interferogram network of no gap and
1102 those with random gaps. Black, green and blue polylines indicate the results computed from
1103 simulated data with noise cases A, B and C (Table 3).

1104

1105 Figure 9. The influence of the change in number of interferograms chosen by various
1106 temporal baseline thresholds (blue) and due to random data gaps (green) on SBAS-derived
1107 unweighted LS rates. From top to bottom: simulated signal cases A, B and C (Table 2). From
1108 left to right: residual interferogram noise cases 1 to 3 (Table 3). Black dashed boxes in the top
1109 panel used to contrast between positive and negative rates that indicates contradictory trends
1110 in some cases.

1111

1112 Figure 10. Interferogram networks without (left) and with (middle, right) missing images. The
1113 number of missing images is 34 out of 133 corresponding to about 25%, which are selected
1114 randomly (middle) and in the northern winter season (right). The networks are formed using a
1115 temporal baseline threshold of ~ 4 months and a perpendicular baseline threshold of 200
1116 meters. Gray lines indicate interferograms, with images denoted by black dots. Red dots
1117 indicate missing images (i.e., gaps). The number under each network refers to the number of
1118 interferograms.

1119

1120 Figure 11. Comparison of unweighted LS rates computed from linear plus annual signals
1121 according to interferogram networks with no gaps, random gaps and “winter” gaps. The
1122 networks adopt a temporal baseline threshold of ~ 4 months and a perpendicular baseline
1123 threshold of 200 meters. From left to right are simulated signal cases 1 to 3 (Table 2). Black,
1124 green and blue polylines indicate the results computed from simulated data with noise cases
1125 A, B, and C (Table 3). Red horizontal lines represent the simulated rate.

1126

1127 Figure 12. Comparison of the RMSs of the difference between simulated and SBAS-derived
1128 deformation time series for all pixels between the interferogram networks of no gaps and

1129 those with randomly chosen gaps and “winter” gaps corresponding to ~25% missing images.
1130 Black, green and blue polylines indicate the results computed from simulated data with noise
1131 cases A, B and C (Table 3).

1132

1133 Figure 13. The dependence of the r -numbers and SBAS-derived unweighted LS rates for
1134 1,000 pixels on the number of chosen interferograms with various linear plus annual signals.
1135 From left to right are simulated signal cases 1 to 3 (Table 2). Black, green and blue polylines
1136 show the results for noise cases A, B and C (Table 3). Red horizontal lines represent the
1137 simulated rates.

1138

1139 Figure 14. (left): Comparison of the change in the r -numbers and the RMSs of the difference
1140 between simulated and SBAS-derived deformation time series. Black, green and blue
1141 polylines indicate the results computed from simulated data with noise cases A, B and C
1142 (Table 3). Dashed lines indicate the “optimal” r -numbers corresponding to the networks with
1143 smallest amount of interferograms chosen with all RMSs being smaller than the chosen trade-
1144 off values of 1/10 of the residual interferogram noise (Table 3). (right): Comparison of the
1145 change in the r -numbers according to SBAS network interferogram numbers and computation
1146 time.

1147

1148 Figure 15. An example of simulated and SBAS-derived deformation time series of simulated
1149 signal of linear rate plus annual amplitude of -2 mm/yr plus 2 mm (simulated signal case 1 in
1150 Table 2) contaminated by residual interferogram noise cases A (left), B (middle) and C (right)
1151 as listed in Table 3. The results are computed from applying various SBAS interferogram
1152 networks corresponding to computed r -numbers of ~ 0.2 (251 interferograms, blue polylines),

1153 ~0.8 (1,571 interferograms, green polylines), and ~0.9 (2,330 interferograms, black
1154 polylines), respectively. Red polylines indicate the simulated deformation time series. The
1155 blue line in the right panel shows the extreme case where spurious uplift is indicated, whereas
1156 subsidence is simulated.